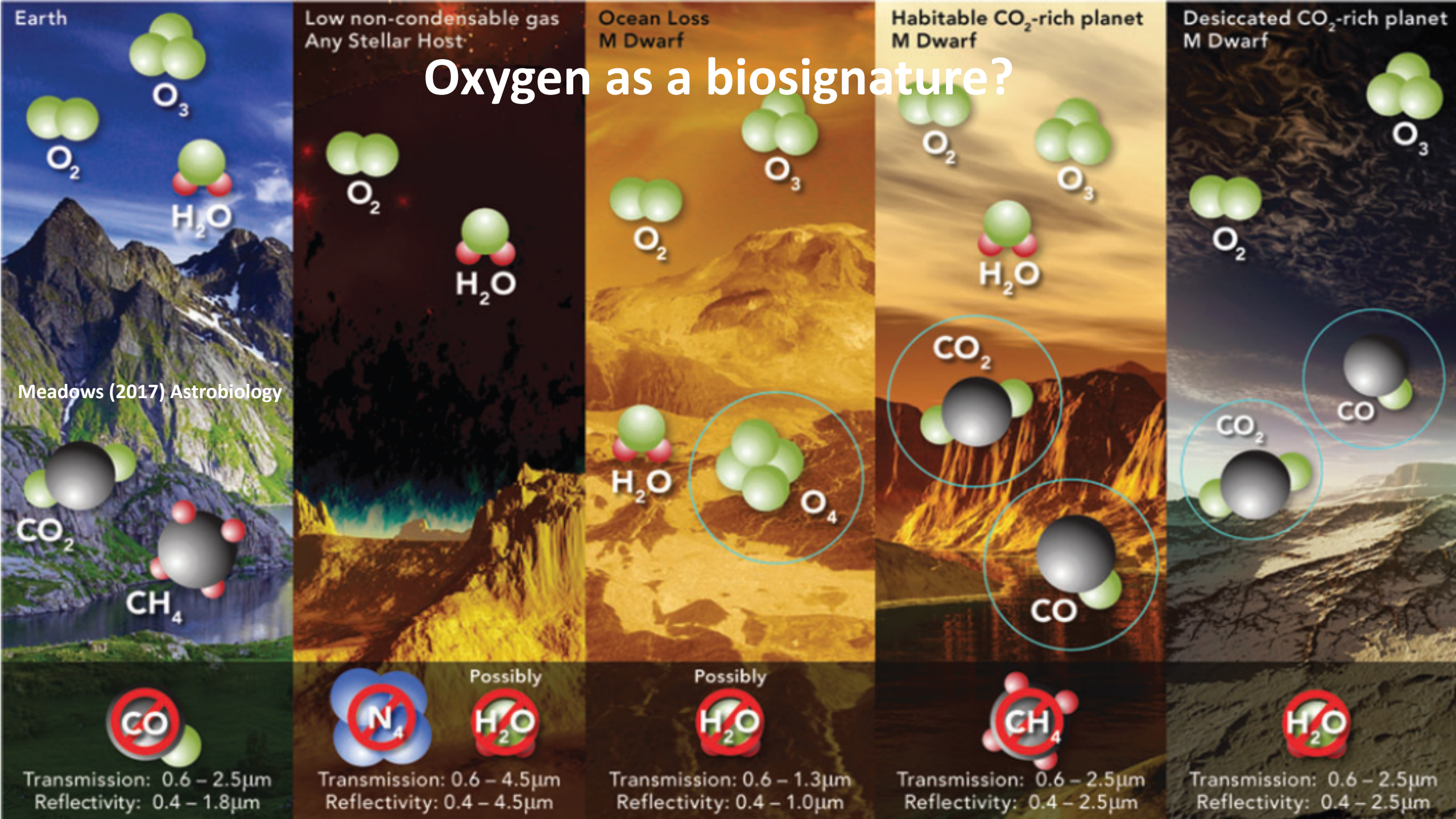


Disequilibrium biosignatures over Earth history and detecting life in anoxic atmospheres

Joshua Krissansen-Totton
University of Washington (soon UCSC)

Oxygen as a biosignature?



Meadows (2017) Astrobiology

Transmission: 0.6 – 2.5μm
Reflectivity: 0.4 – 1.8μm

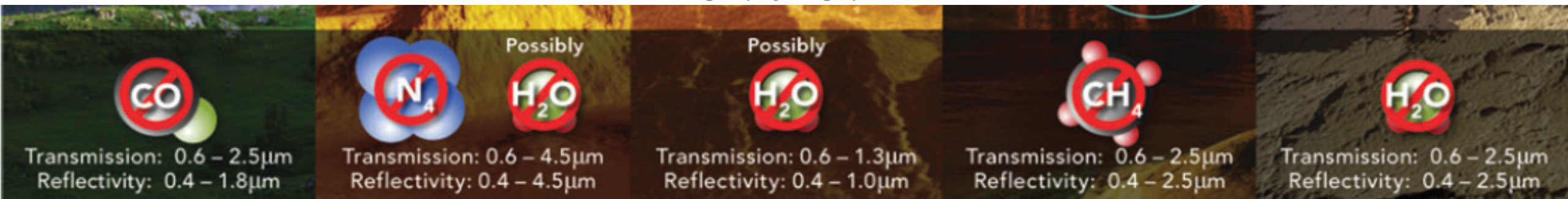
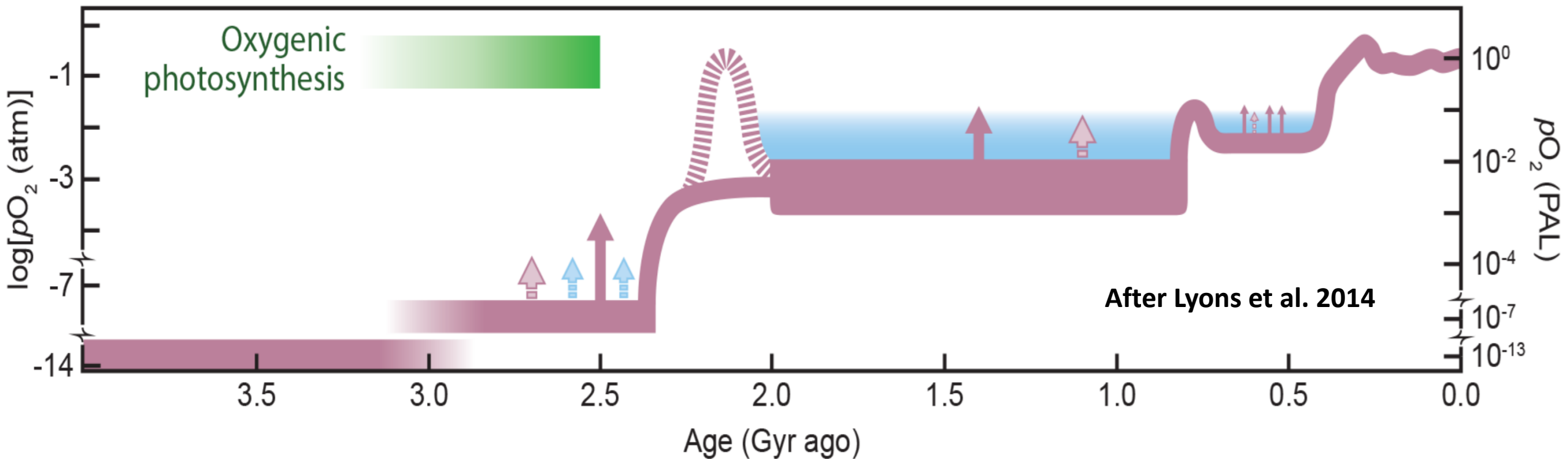
Transmission: 0.6 – 4.5μm
Reflectivity: 0.4 – 4.5μm

Transmission: 0.6 – 1.3μm
Reflectivity: 0.4 – 1.0μm

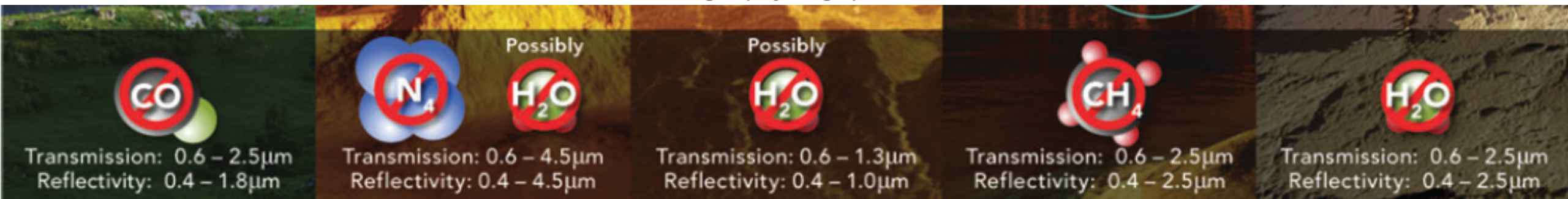
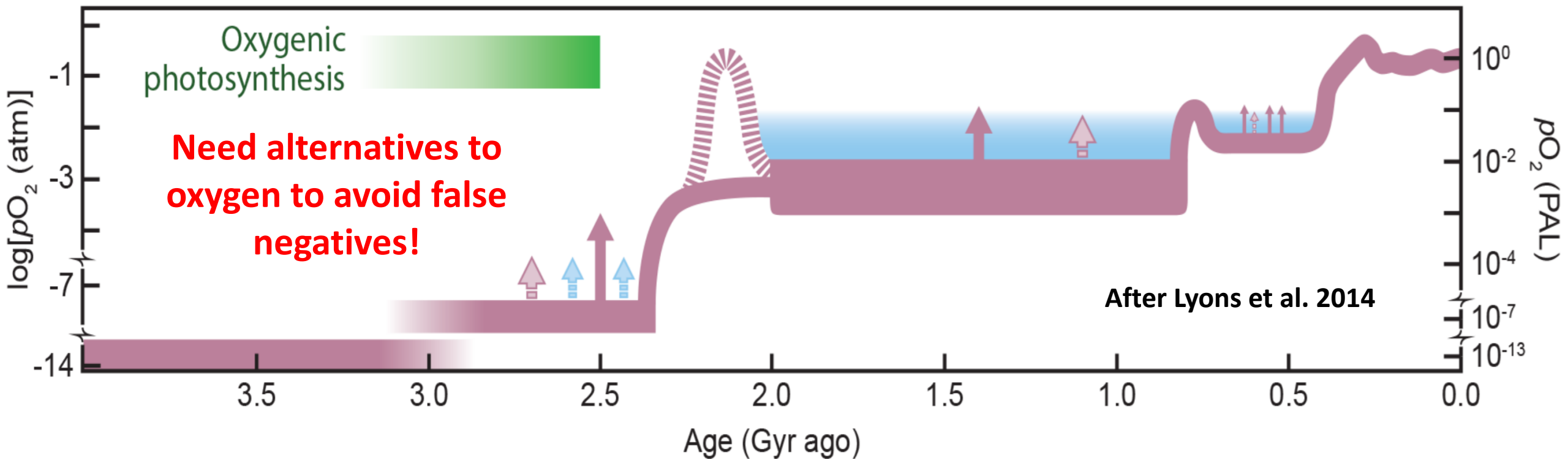
Transmission: 0.6 – 2.5μm
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Transmission: 0.6 – 2.5μm
Reflectivity: 0.4 – 2.5μm

Oxygen as a biosignature?



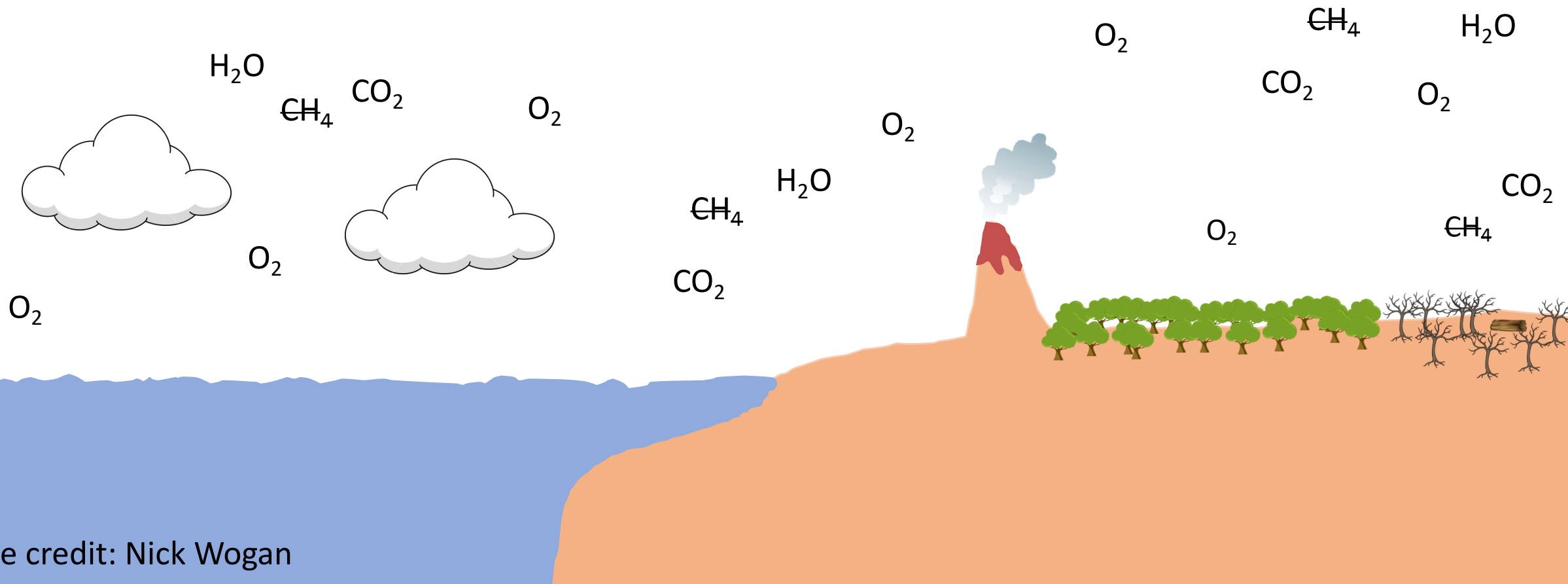
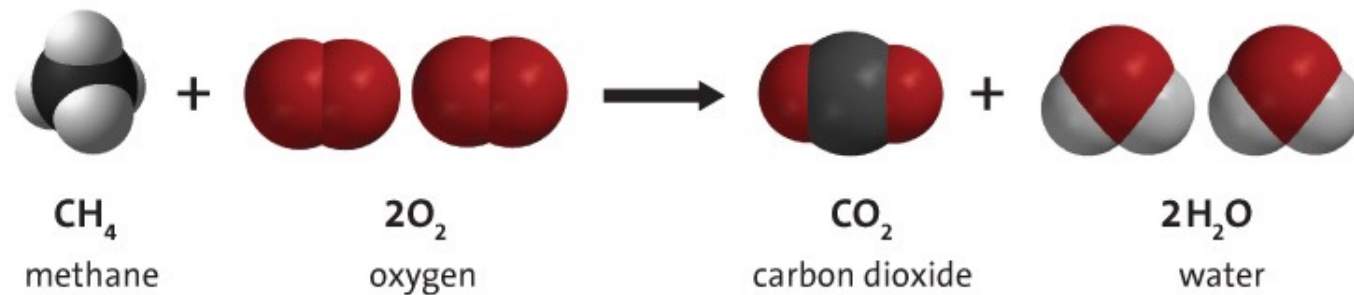
Oxygen as a biosignature?



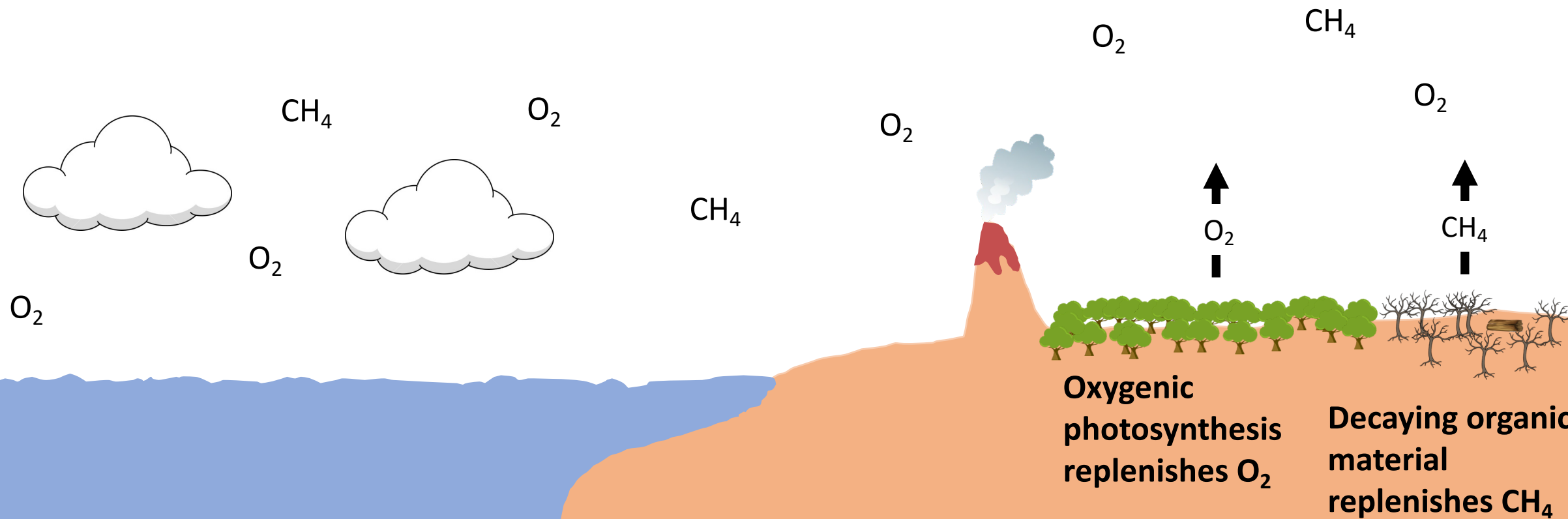
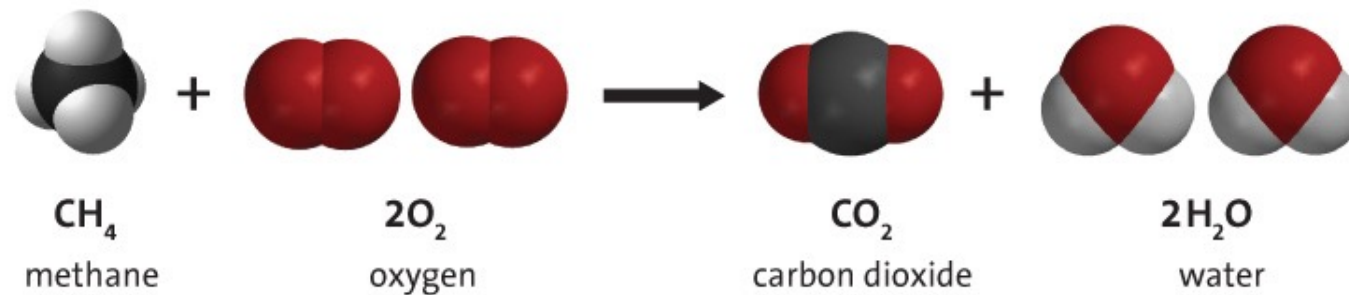


Disequilibrium
The co-existence of incompatible species

Chemical disequilibrium as an exoplanet biosignature

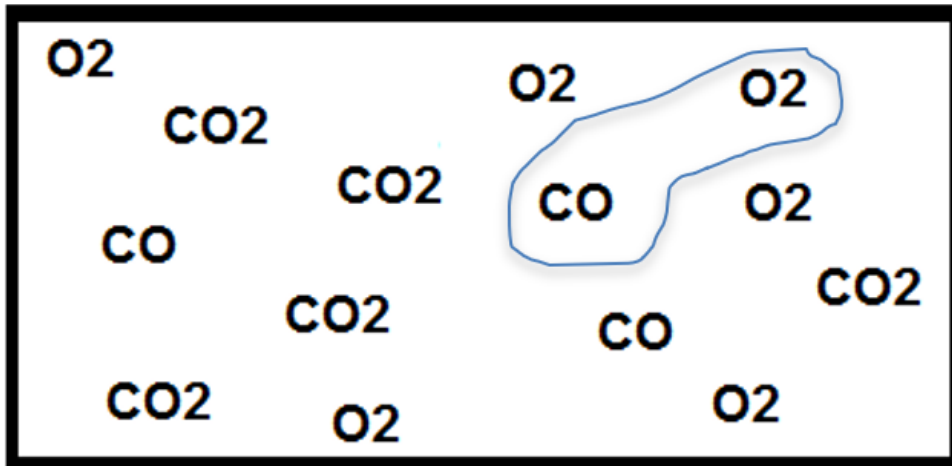


Chemical disequilibrium as an exoplanet biosignature



Quantify chemical disequilibrium

Initial/Observed State

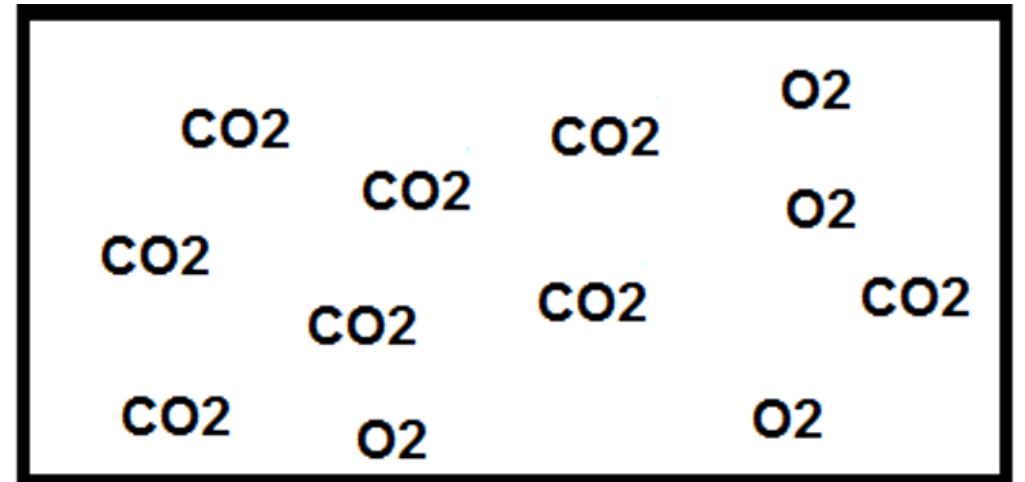


React to
equilibrium via
Gibbs energy
minimization



Mixing ratios
and molalities
change, but
atoms and
charge
conserved.

Equilibrium State

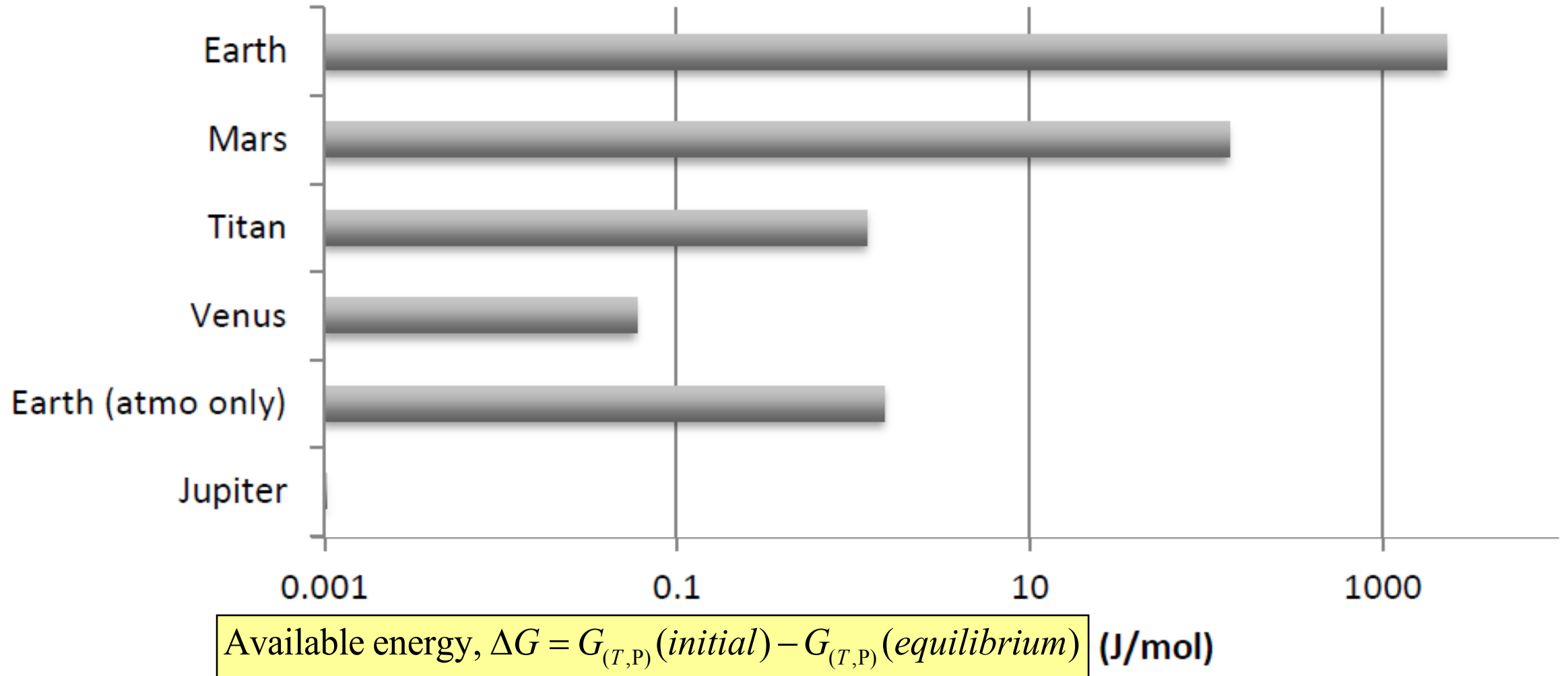
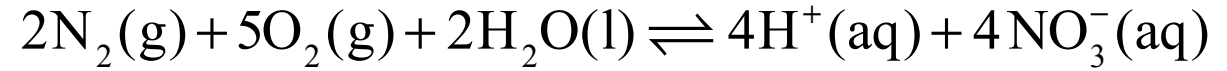


Gibbs energy of initial state, $G_{(T,P)}(\mathbf{n}_{\text{initial}})$

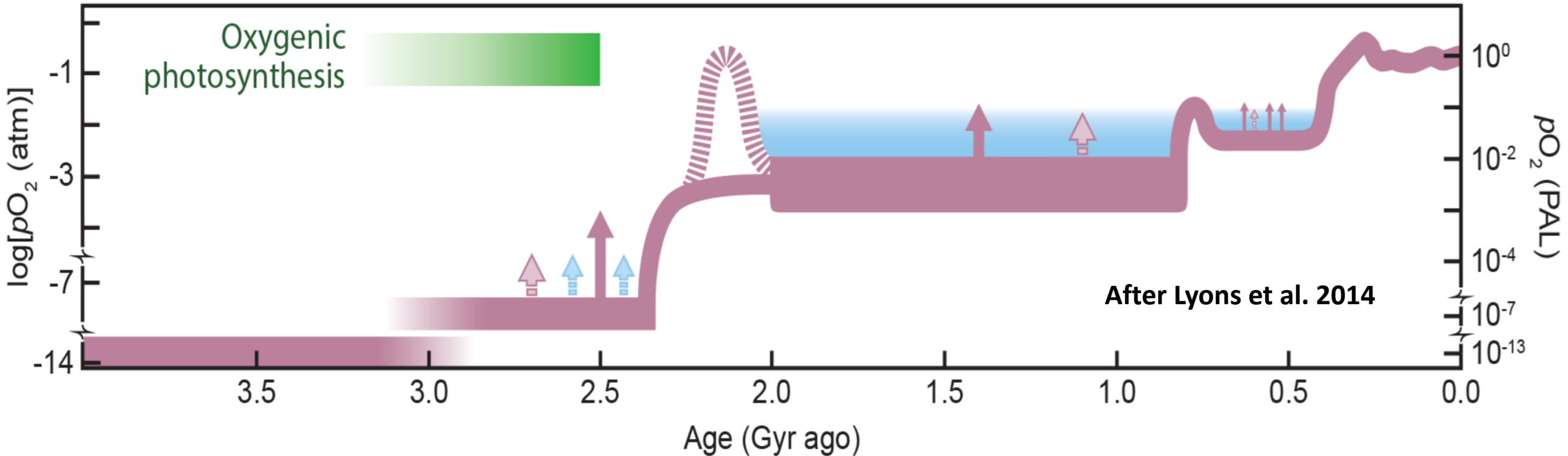
Gibbs energy of equilibrium state, $G_{(T,P)}(\mathbf{n}_{\text{final}})$

$$\text{Available energy, } \Delta G = G_{(T,P)}(\text{initial}) - G_{(T,P)}(\text{equilibrium})$$

Earth has largest disequilibrium in the solar system



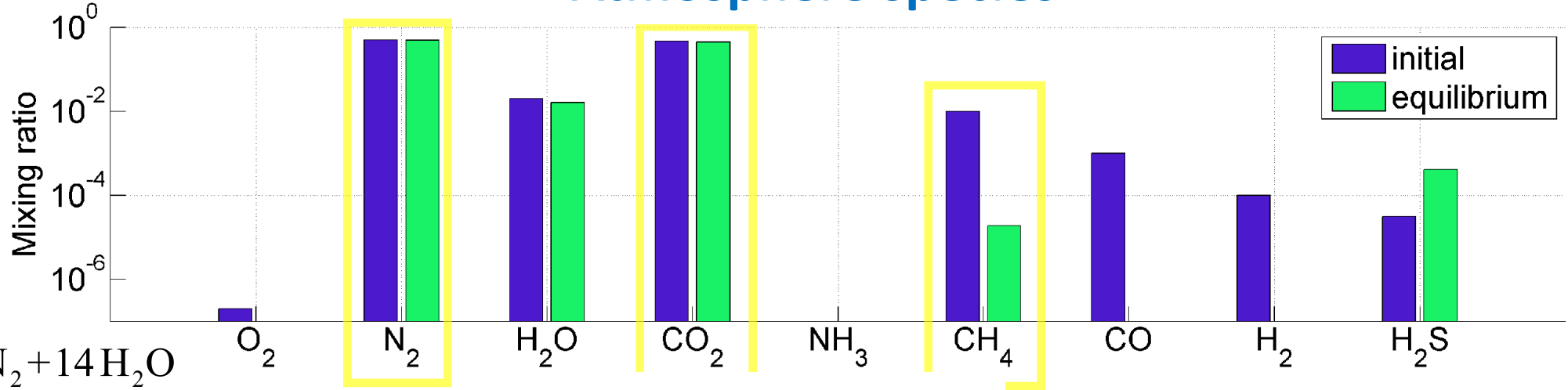
BUT O₂-N₂-H₂O and O₂/O₃-CH₄ are variations on oxygen as a biosignature!



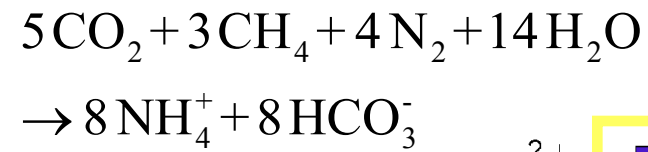
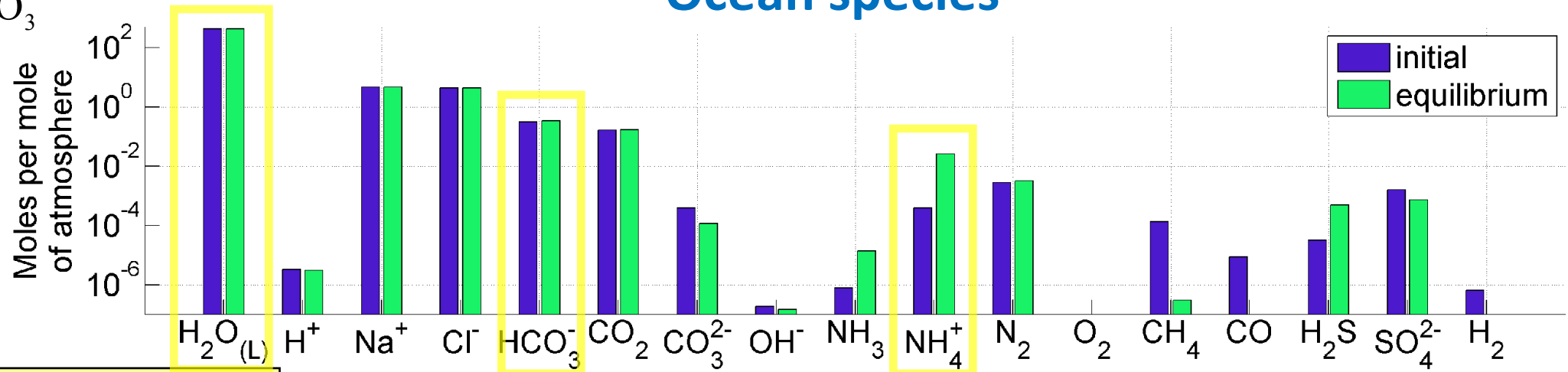
- The early Earth is the only example we have of an anoxic biosphere!

Archean atmosphere-ocean disequilibrium

Atmosphere species

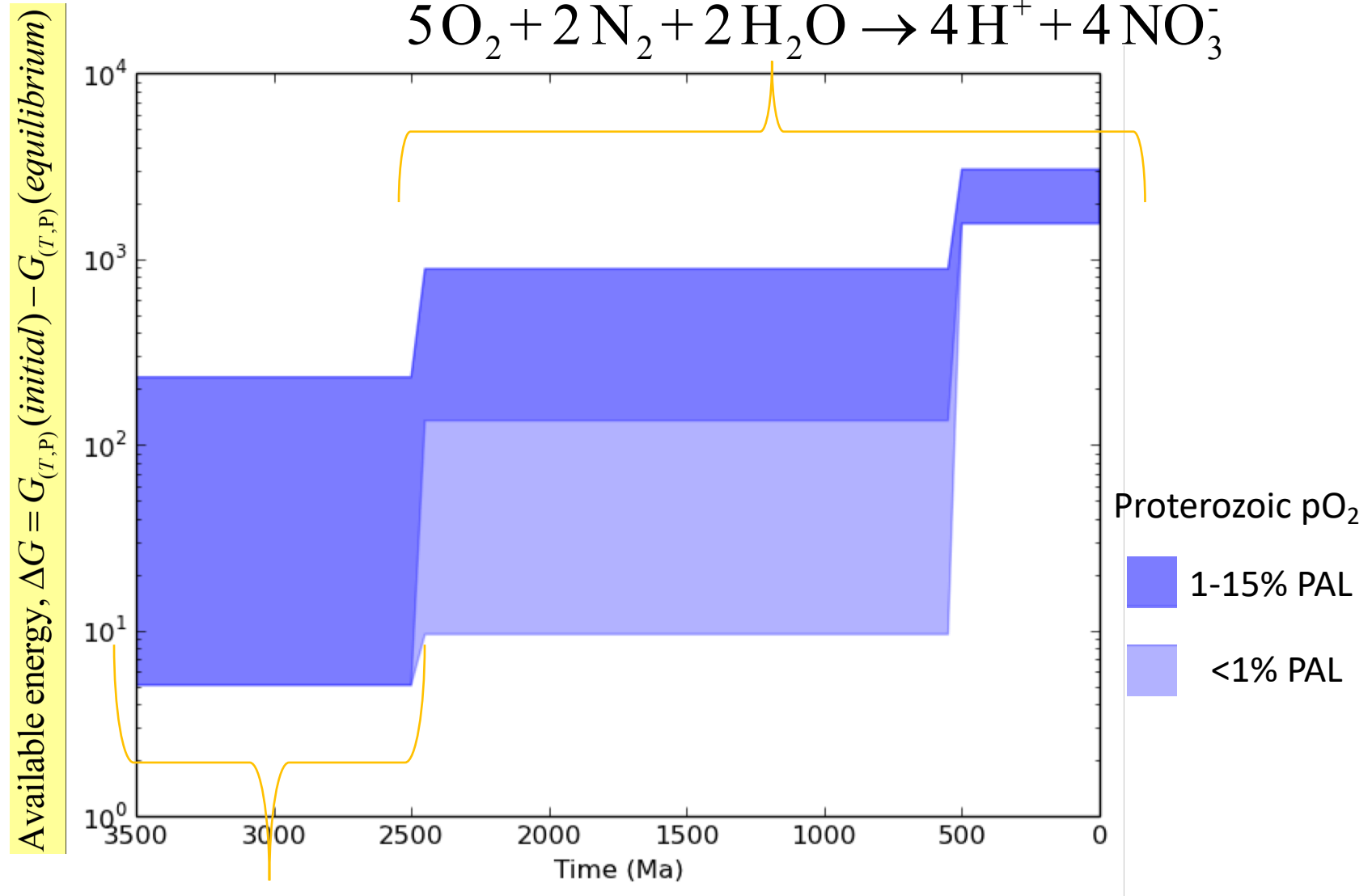


Ocean species

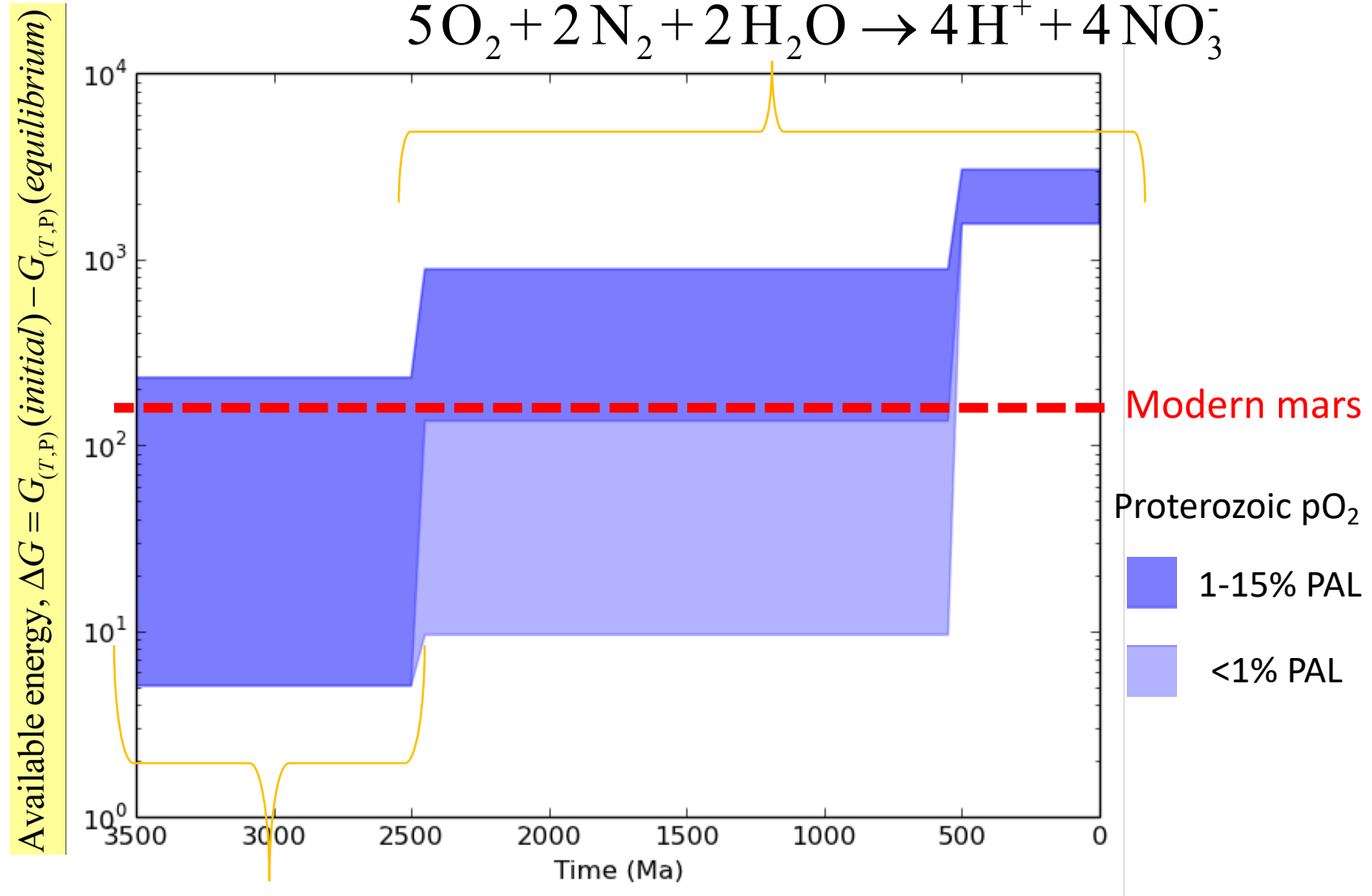


Available energy,
 $\Delta G = G_{(T,P)}(\text{obser}) - G_{(T,P)}(\text{equil})$
 $\Delta G < 231 \text{ J/mol}$

The Evolution of Earth's thermodynamic disequilibrium



The Evolution of Earth's thermodynamic disequilibrium



Would a $\text{CH}_4 + \text{CO}_2$ disequilibrium combination be a good exoplanet biosignature?

- Methanogenesis is an ancient and “primitive” metabolism



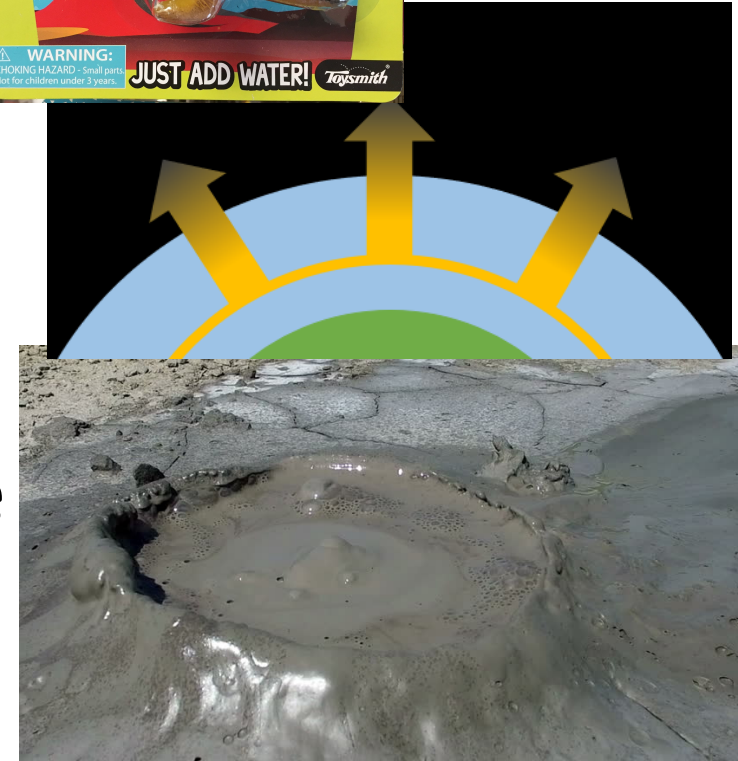
Would a $\text{CH}_4 + \text{CO}_2$ disequilibrium combination be a good exoplanet biosignature?

- Methanogenesis is an ancient and “primitive” metabolism
- CH_4 has short lifetime, perhaps $\sim 30,000$ yrs for diffusion limited escape.

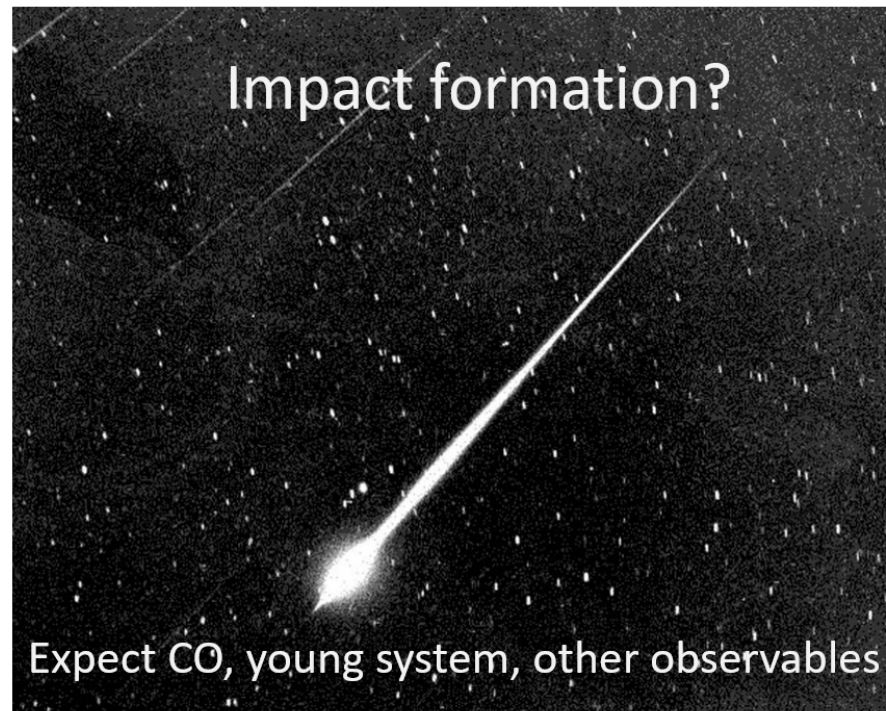
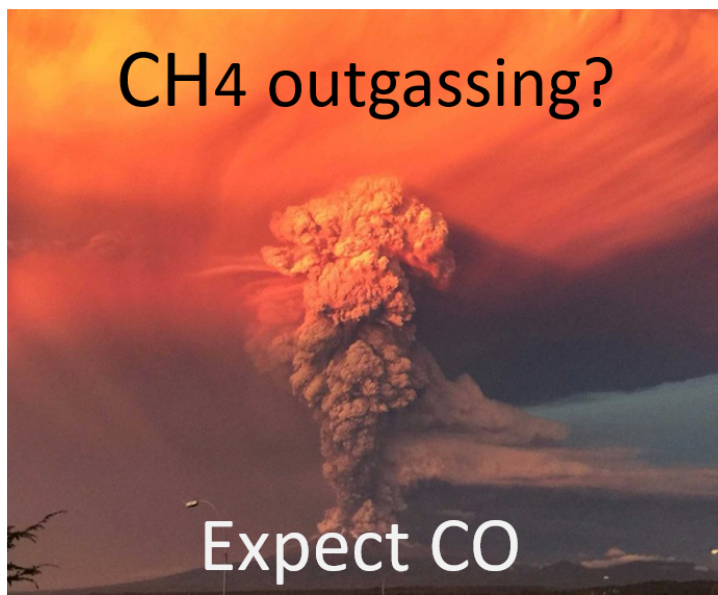


Would a $\text{CH}_4 + \text{CO}_2$ disequilibrium combination be a good exoplanet biosignature?

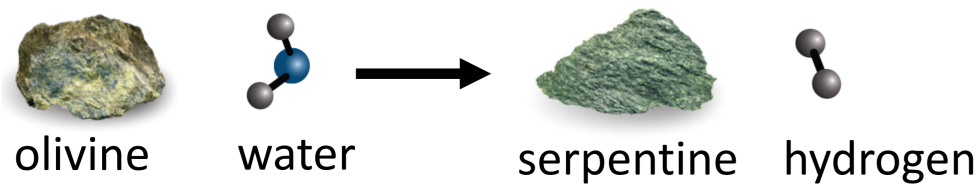
- Methanogenesis is an ancient and “primitive” metabolism
- CH_4 has short lifetime, perhaps $\sim 30,000$ yrs for diffusion limited escape.
- Would not expect $\text{CH}_4 + \text{CO}_2$ atmospheres to persist without replenishing flux. Could there be large abiotic fluxes of CH_4 ?



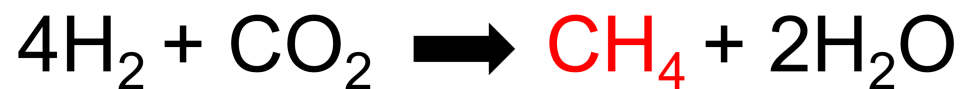
Possible false positives for biogenic CH₄ + CO₂?



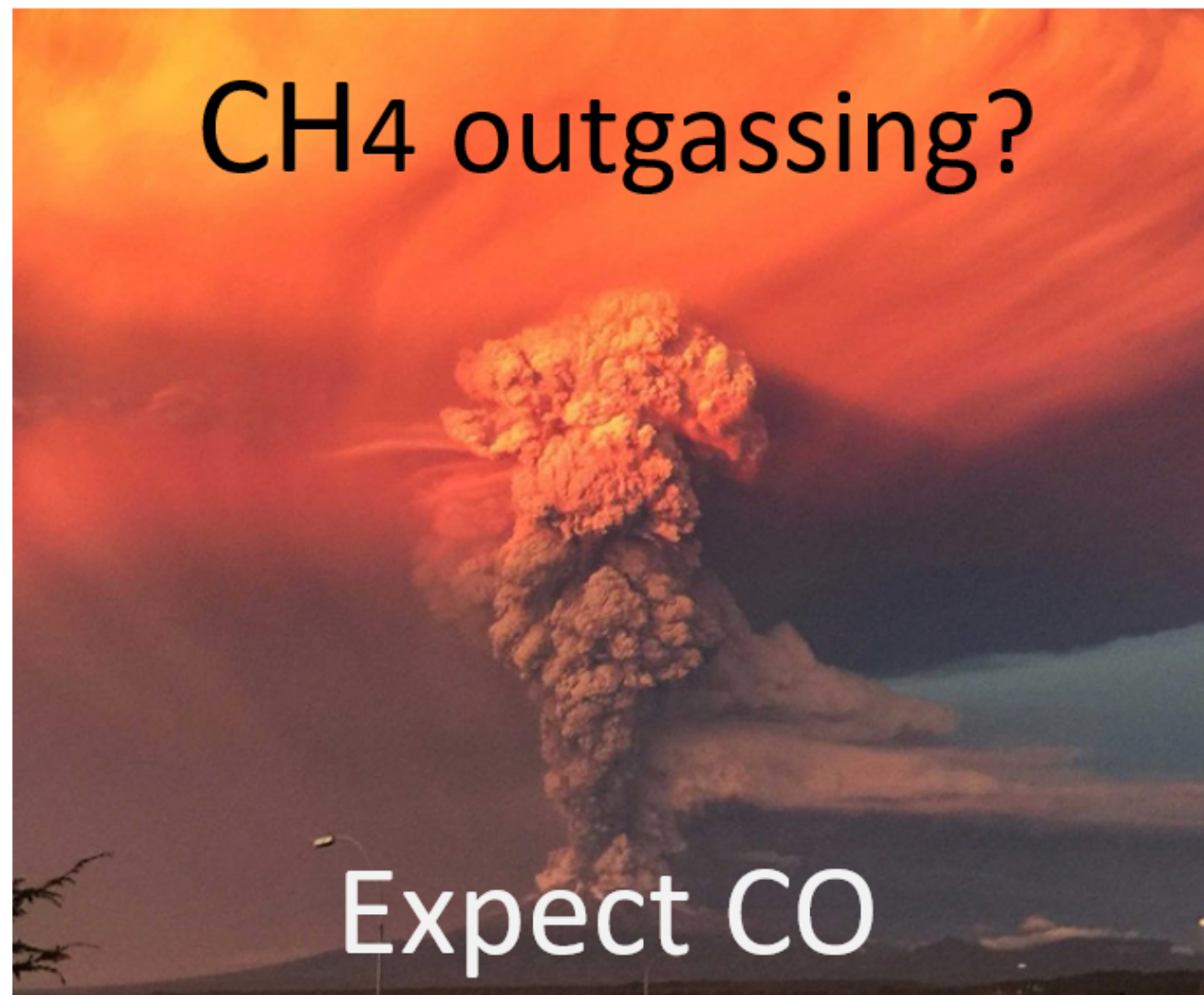
Serpentinization: Source of Hydrogen



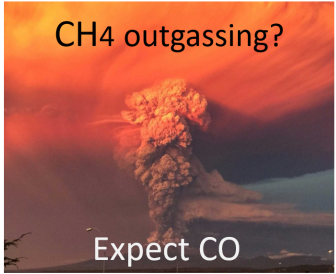
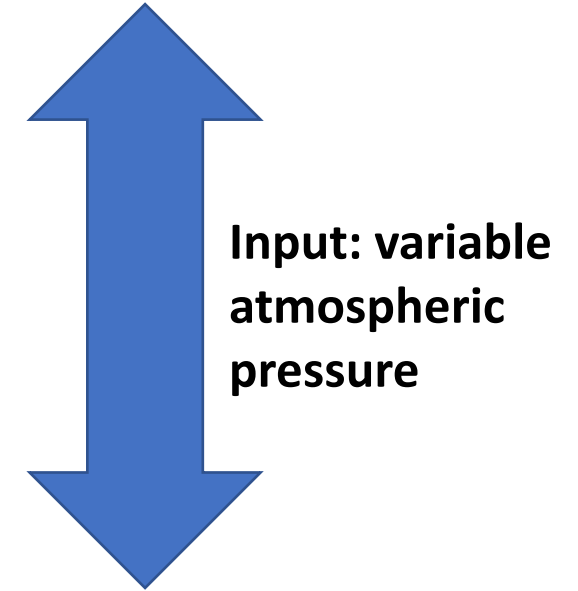
Credit: Jennifer Glass



Could a planet outgas CH_4 and CO_2 ?



Output: CO, CH₄, CO₂, H₂O, H₂ abundances



Inputs: variable temperature, pressure, composition

Continental magma chamber

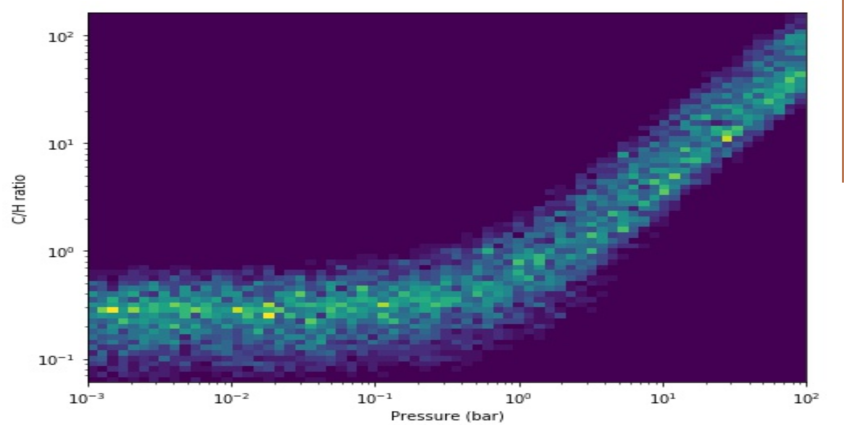
Input: variable ocean depth

Output: CO, CH₄, CO₂, H₂O, H₂ abundances



Inputs: variable temperature, pressure, composition

Submarine magma chamber



Output: CO, CH₄, CO₂, H₂O, H₂ abundances

Key outputs: CO/CO₂ and CO/CH₄ outgassing ratios

Inputs: variable temperature, pressure, composition

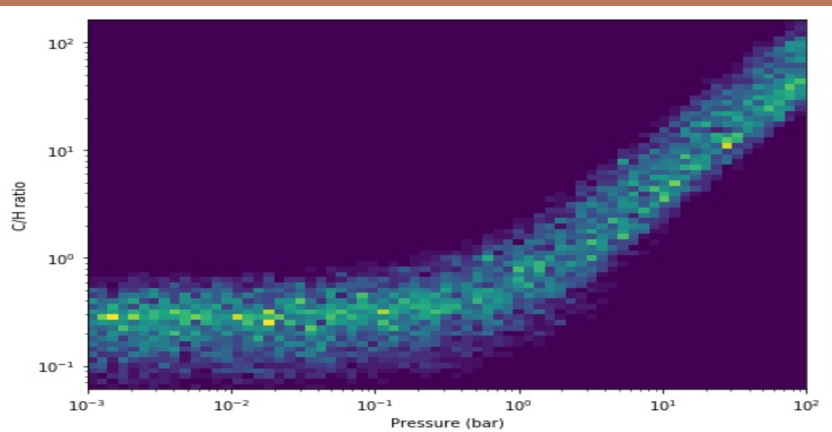
Continental magma chamber

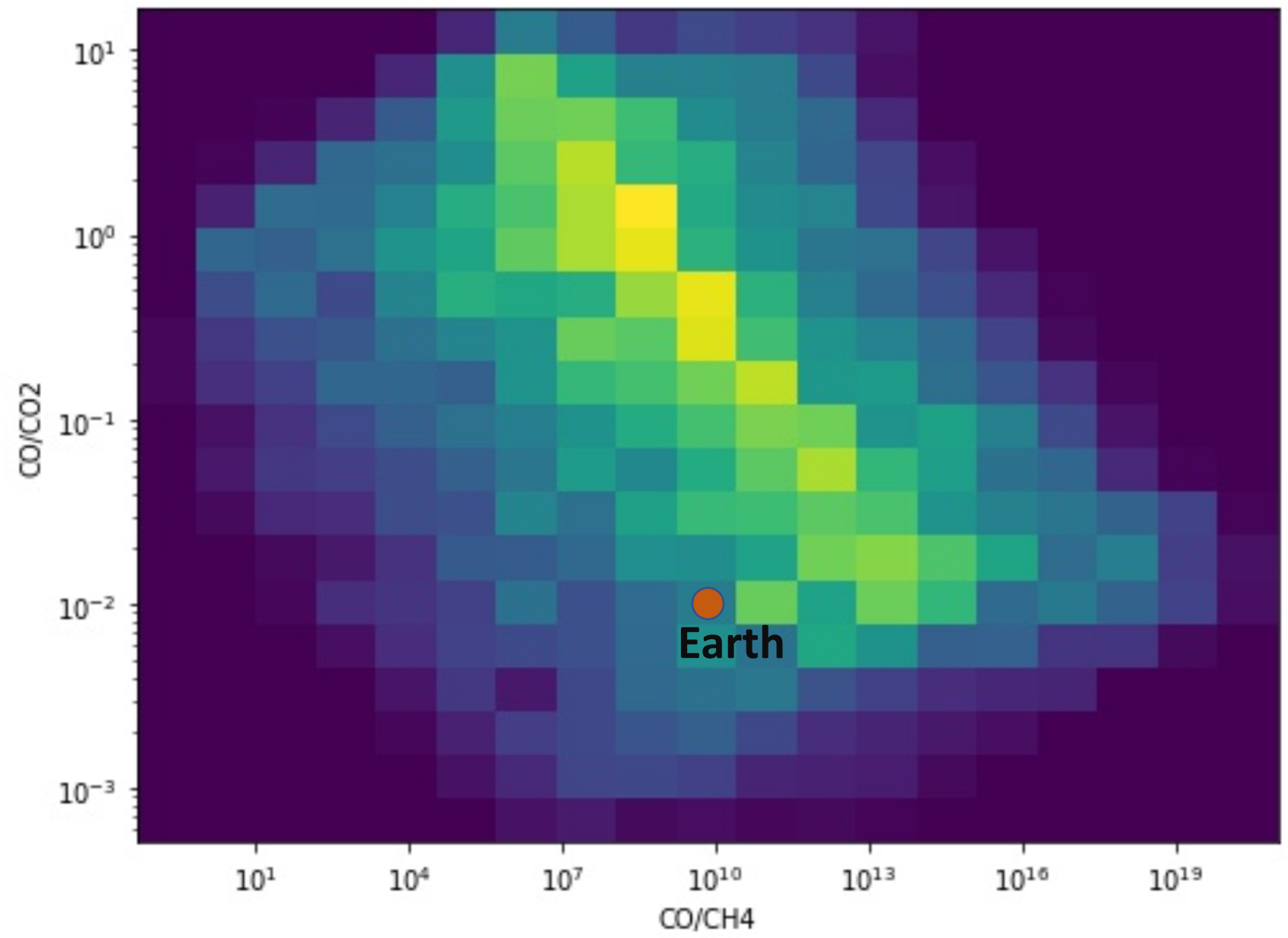
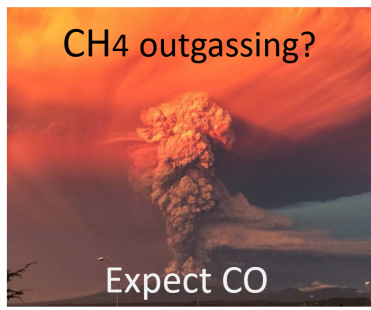
Input: variable ocean depth

Output: CO, CH₄, CO₂, H₂O, H₂ abundances

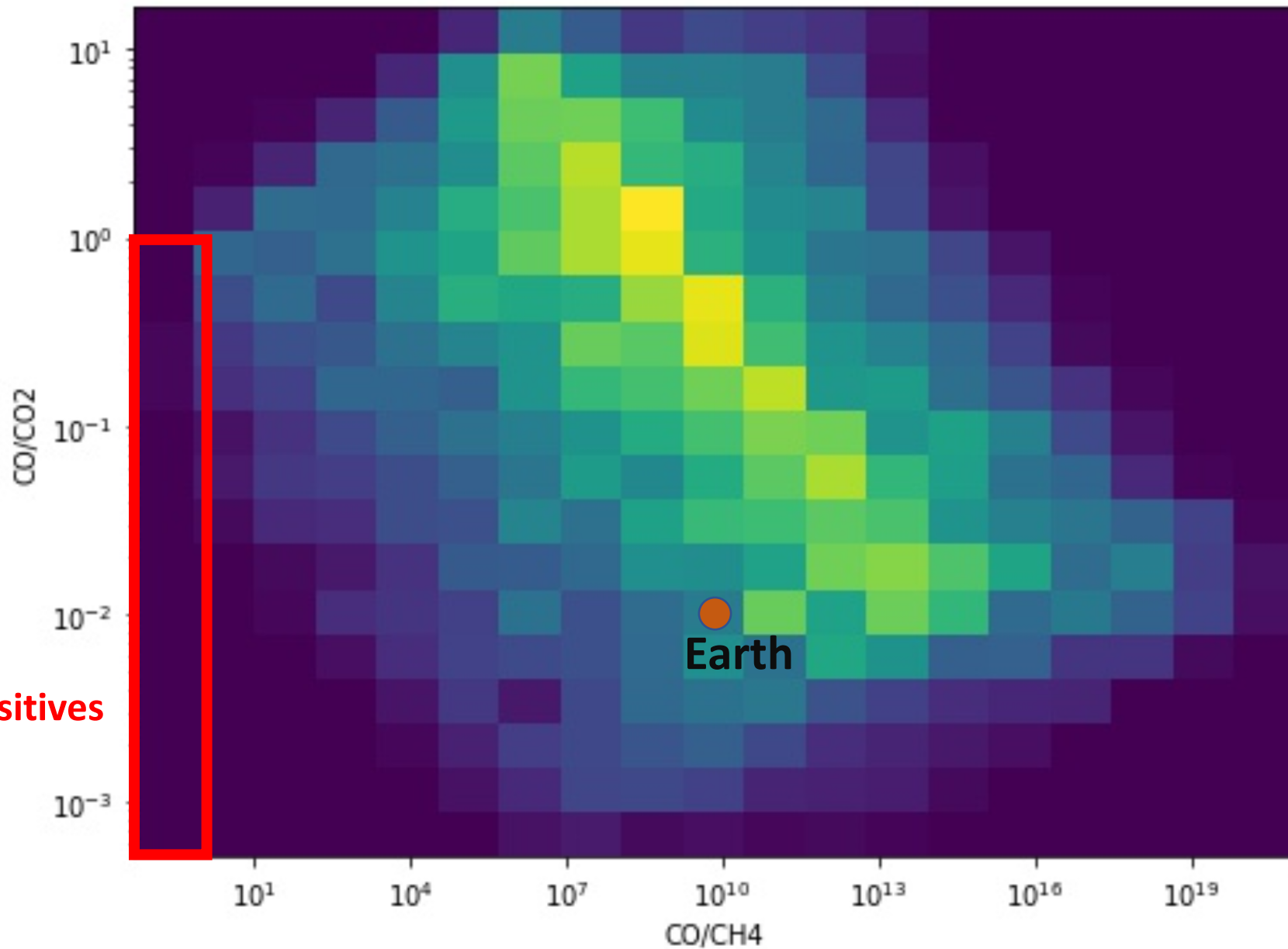
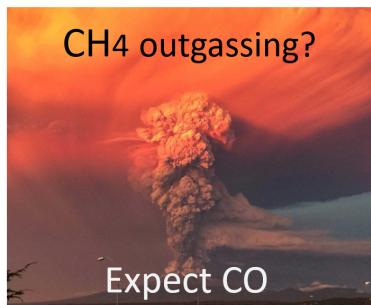
Inputs: variable temperature, pressure, composition

Submarine magma chamber





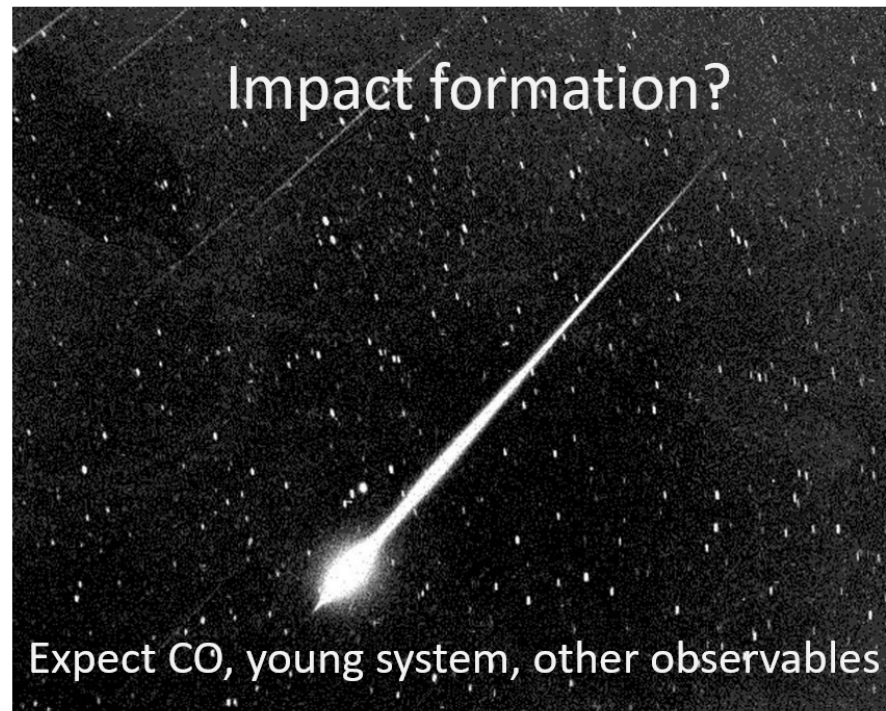
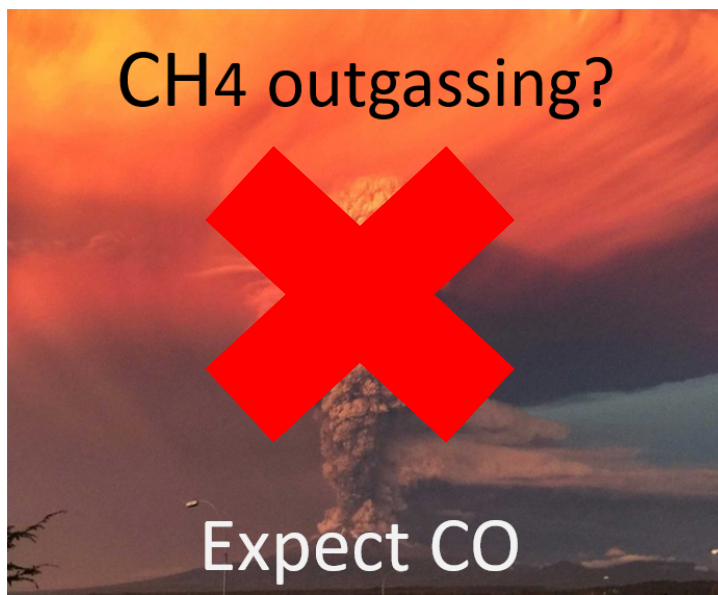
Key point: simultaneous outgassing of CO₂ and CH₄ without CO is very unlikely



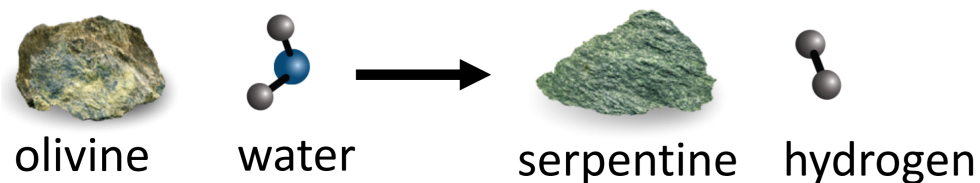
Possible false positives

Key point: simultaneous outgassing of CO₂ and CH₄ without CO is very unlikely

Possible false positives for biogenic CH₄ + CO₂?



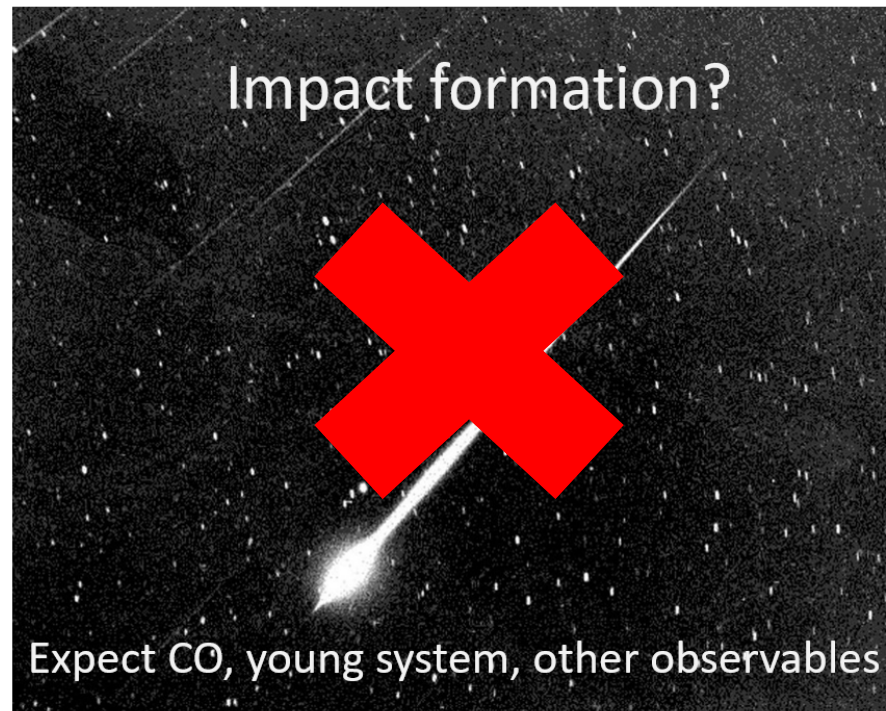
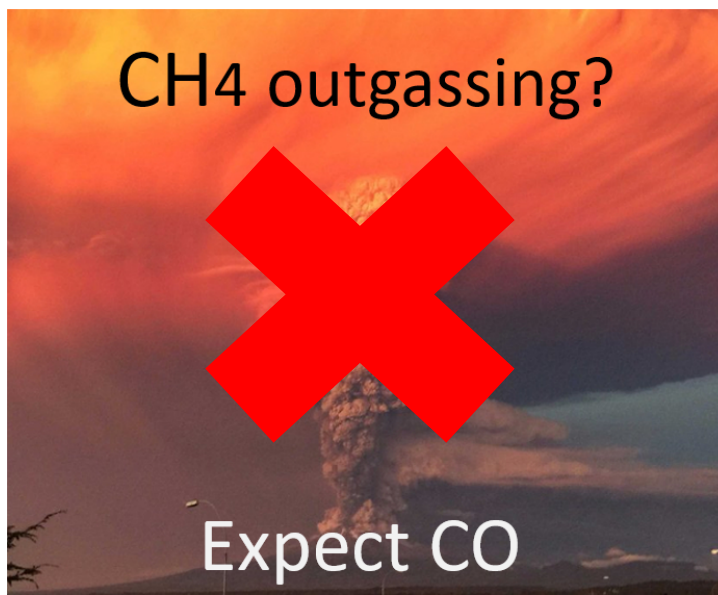
Serpentinization: Source of Hydrogen



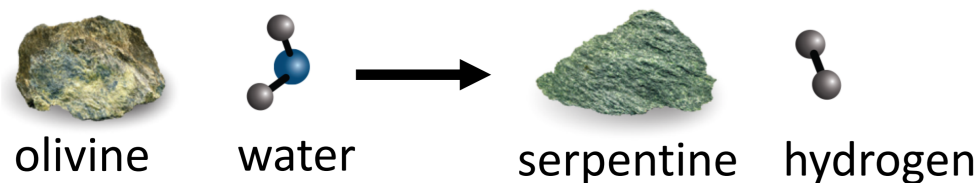
Credit: Jennifer Glass



Possible false positives for biogenic CH₄ + CO₂?



Serpentinization: Source of Hydrogen



Credit: Jennifer Glass



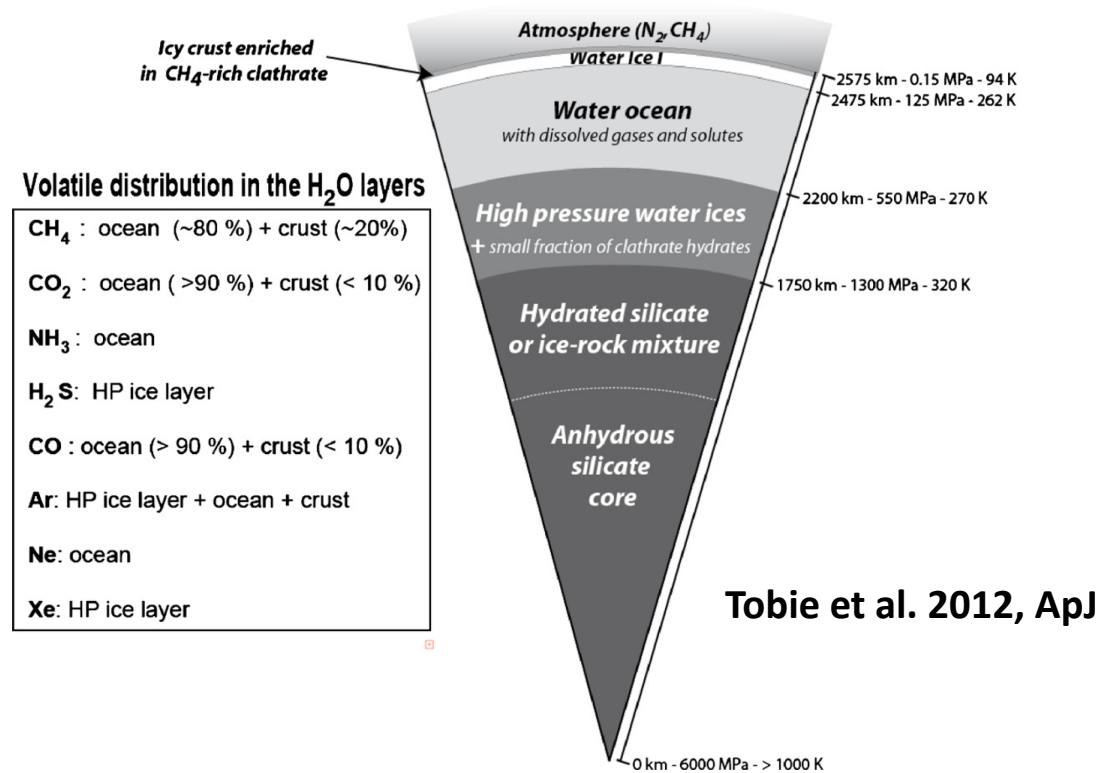
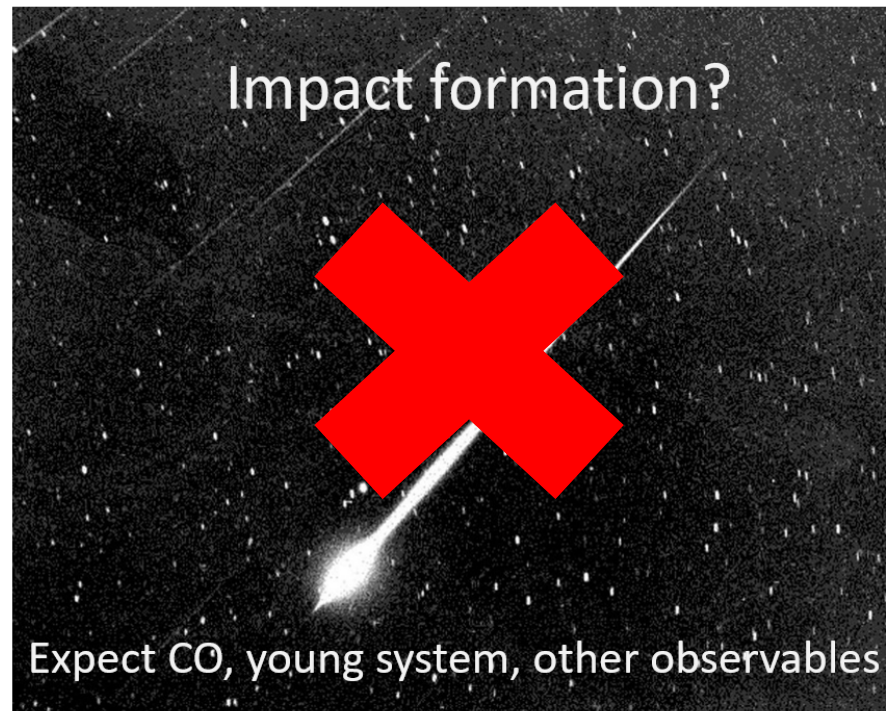
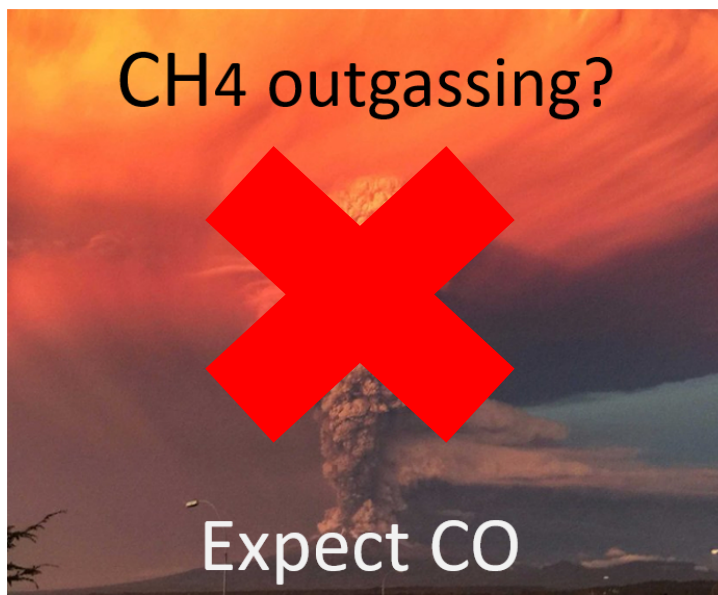


Figure 2. Possible present-day structure of Titan's interior and estimated volatile distribution within the thick H₂O mantle. The radius, temperature, and pressure of each interface are only indicative.

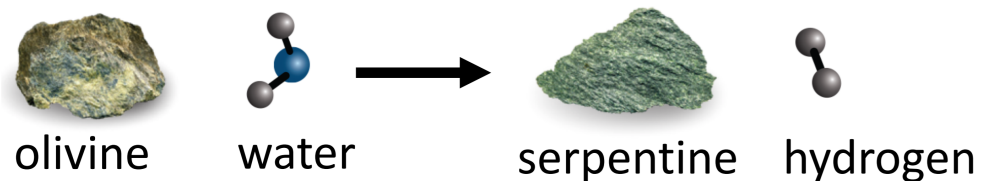
Table: Estimated photochemical lifetimes of CH₄ for Earth-sized planets with surface oceans of varying size, and Titan-like initial volatile inventories.

Mass of surface ocean (wt% planet mass)	0.1 wt% water	1.0 wt% water	10 wt% water	50 wt% water
Lifetime of CH ₄ (Myr)	2	20	200	1000

Possible false positives for biogenic CH₄ + CO₂?



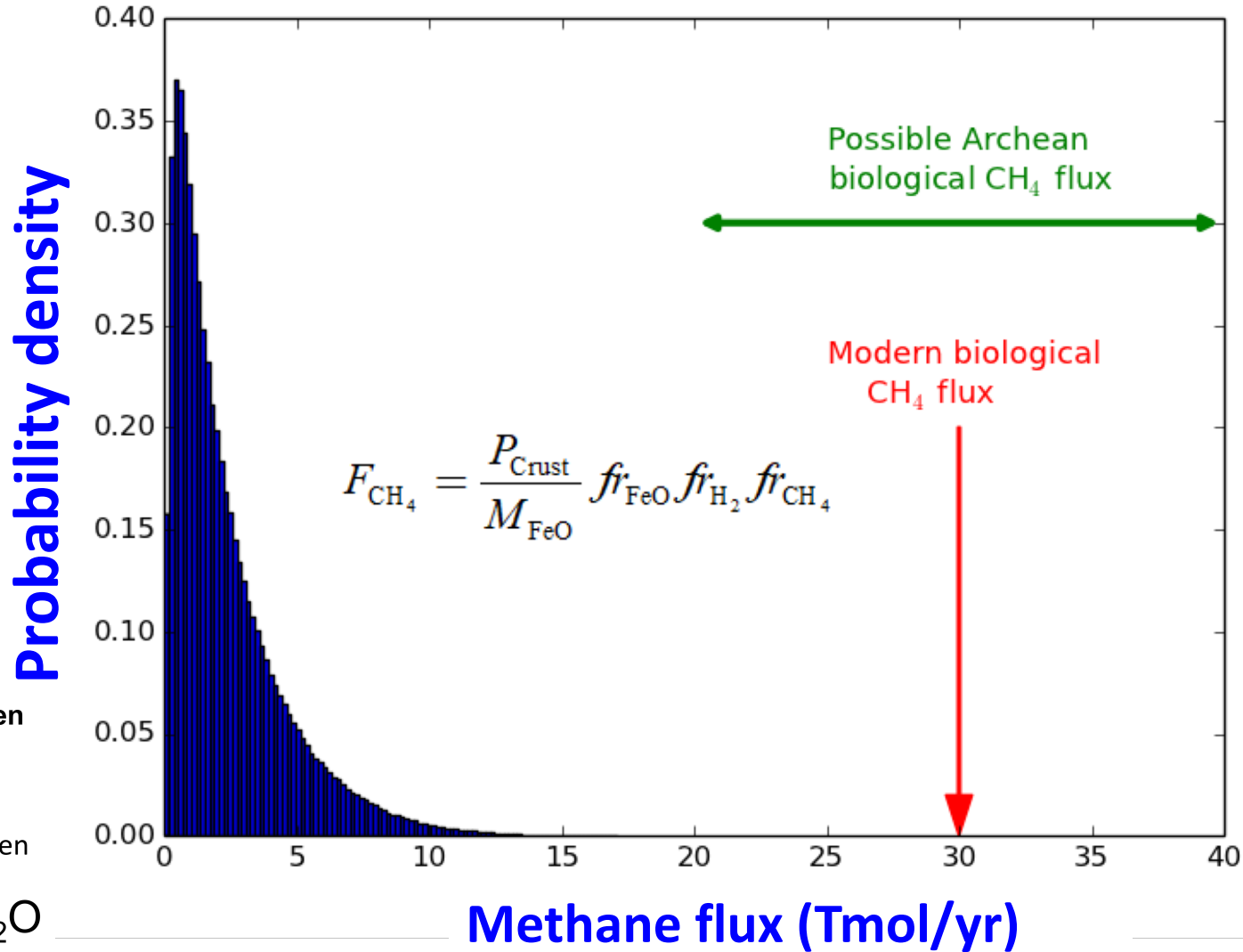
Serpentinization: Source of Hydrogen



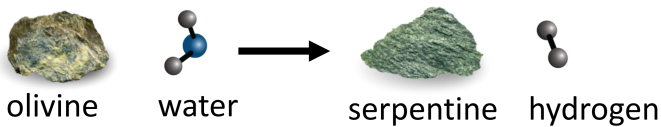
Credit: Jennifer Glass



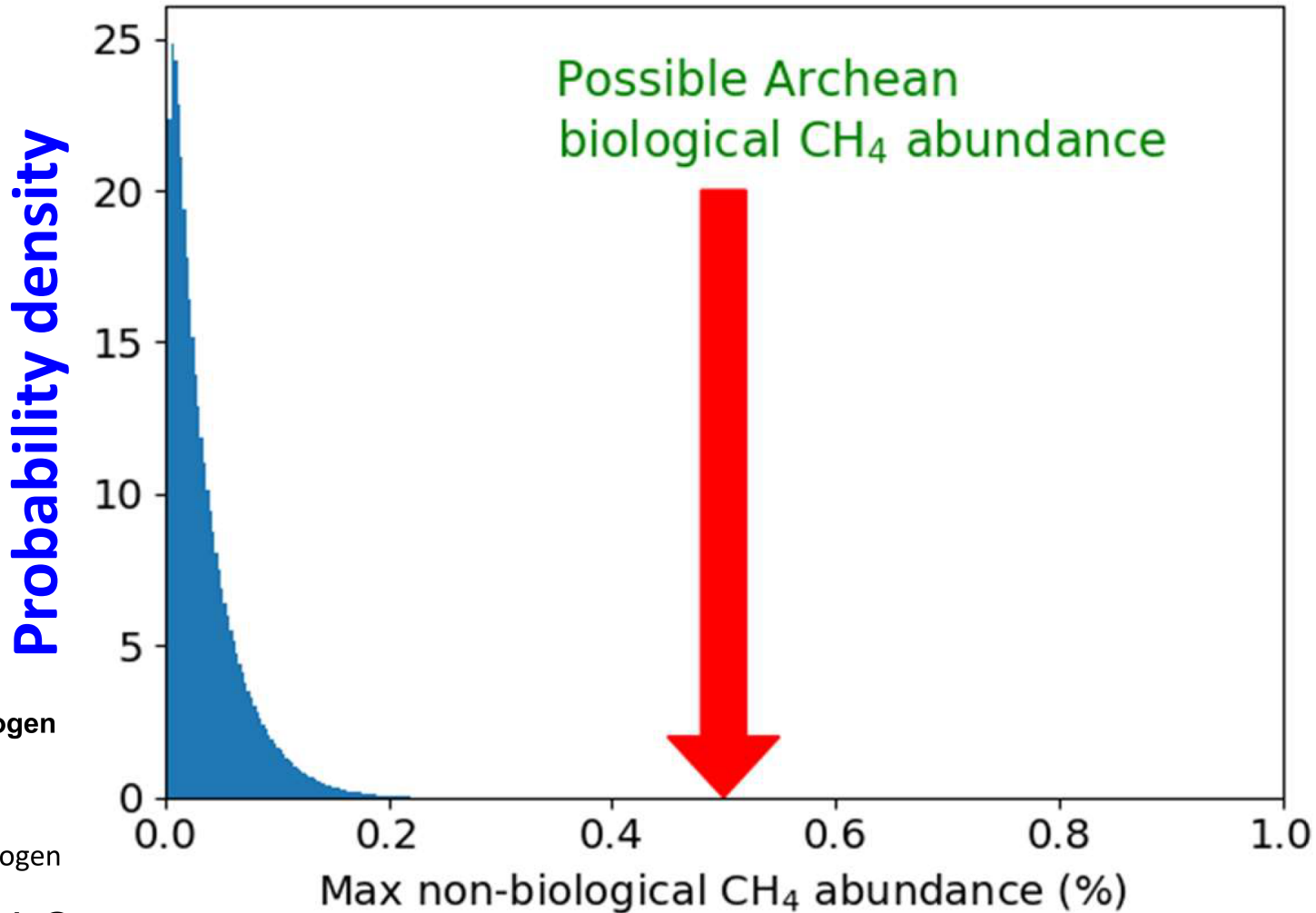
Maximum abiotic methane from serpentinization << biological fluxes



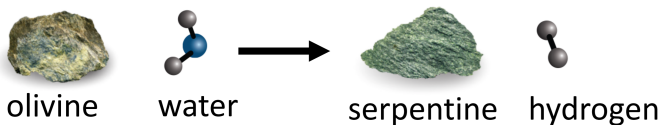
Serpentinization: Source of Hydrogen



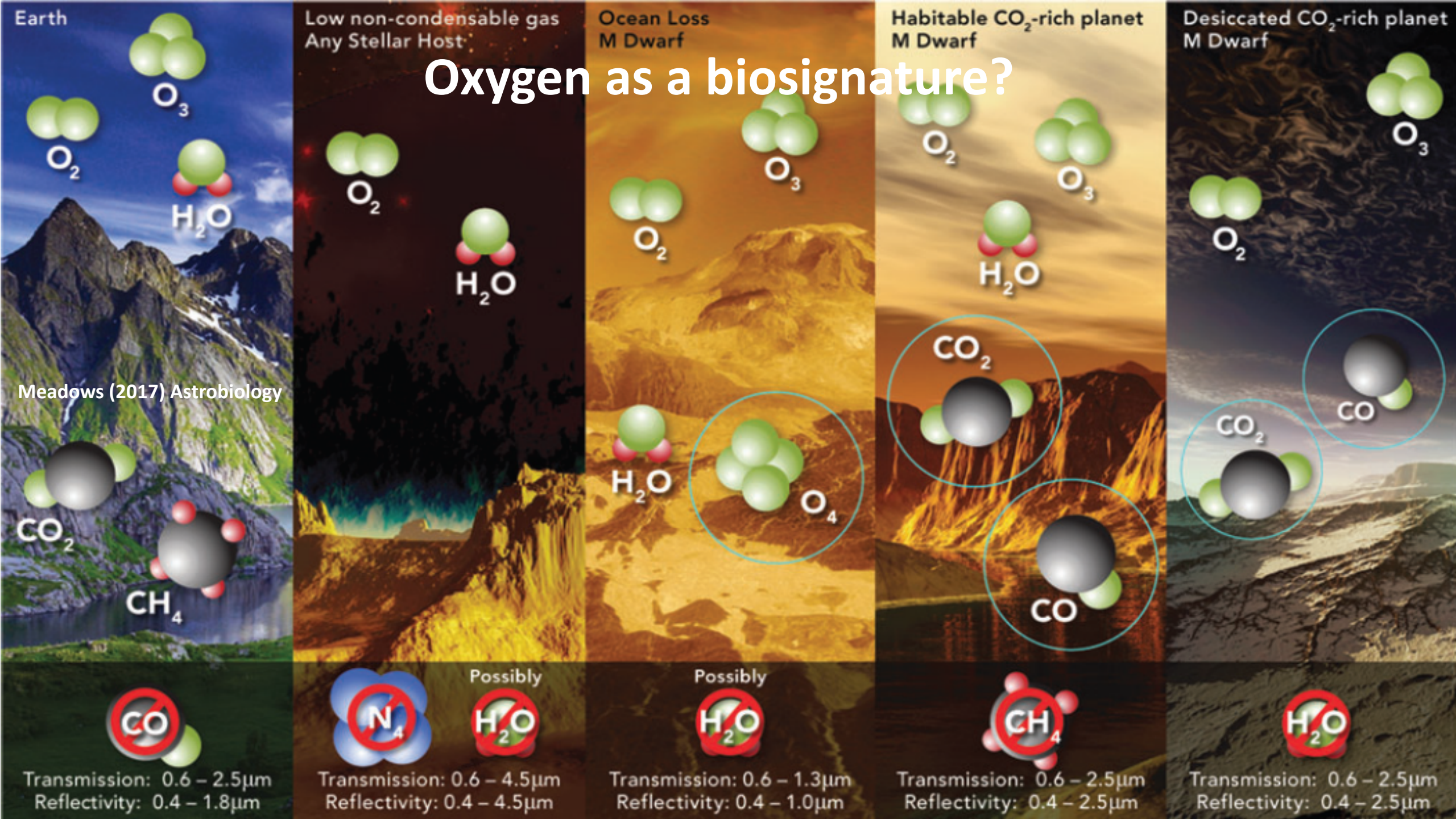
Maximum abiotic methane from serpentinization << biological fluxes



Serpentinization: Source of Hydrogen



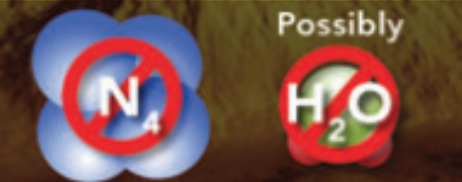
Oxygen as a biosignature?



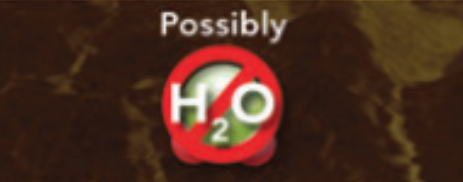
Meadows (2017) Astrobiology



Transmission: 0.6 – 2.5μm
Reflectivity: 0.4 – 1.8μm



Possibly
Transmission: 0.6 – 4.5μm
Reflectivity: 0.4 – 4.5μm



Transmission: 0.6 – 1.3μm
Reflectivity: 0.4 – 1.0μm



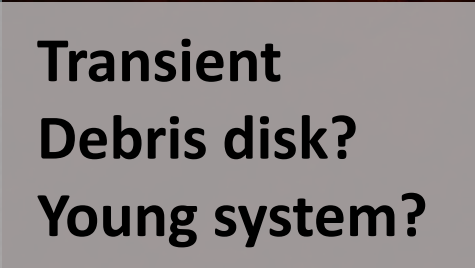
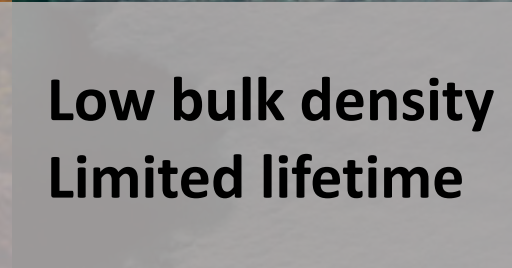
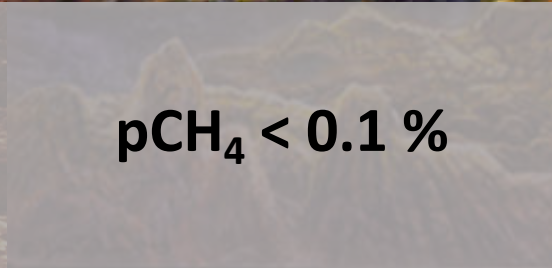
Transmission: 0.6 – 2.5μm
Reflectivity: 0.4 – 2.5μm



Transmission: 0.6 – 2.5μm
Reflectivity: 0.4 – 2.5μm



Methane as a biosignature?





The bottom line:

The simultaneous detection of atmospheric CO₂ and >1 part per thousand CH₄ (and little/no CO) on a habitable planet is a promising disequilibrium biosignature.

Also: CO₂/CH₄ pair more common O₂/O₃ biosignatures?

Is this detectable with the James Webb Space Telescope?

The inverse problem

Archean Earth-like
composition (no haze)



Mass = $0.772 M_{\text{Earth}}$
 Radius = $0.91 R_{\text{Earth}}$
 $p_{\text{CO}_2} = 5\%$
 $p_{\text{CH}_4} = 0.5\%$
 $p_{\text{CO}} = 0\%$
 $p_{\text{H}_2\text{O}} = 1\%$
 $T_{\text{strat}} = 200 \text{ K}$
 $P_{\text{surf}} = 1 \text{ bar}$

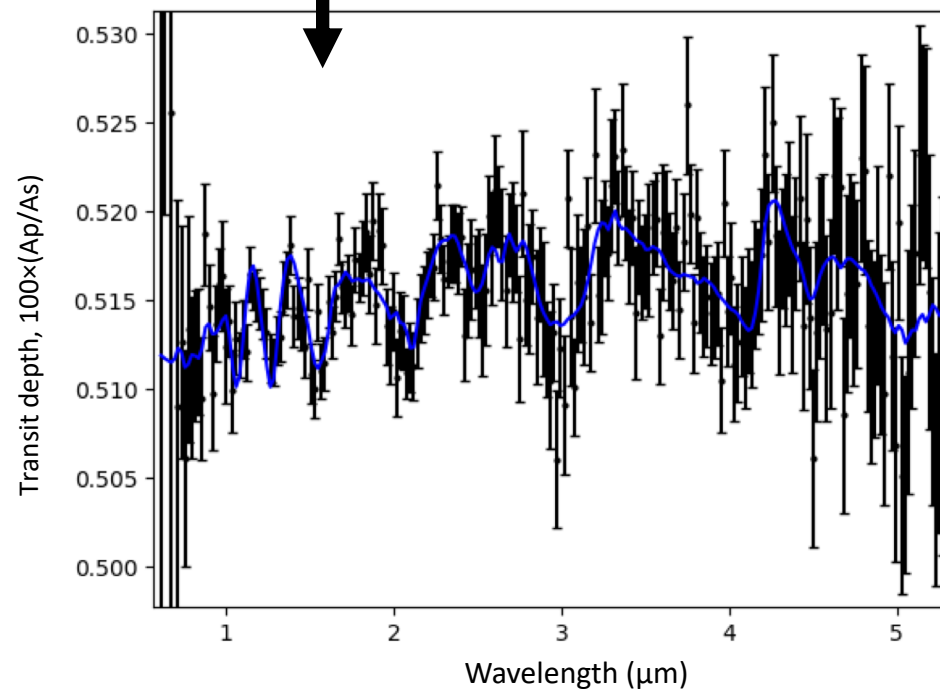
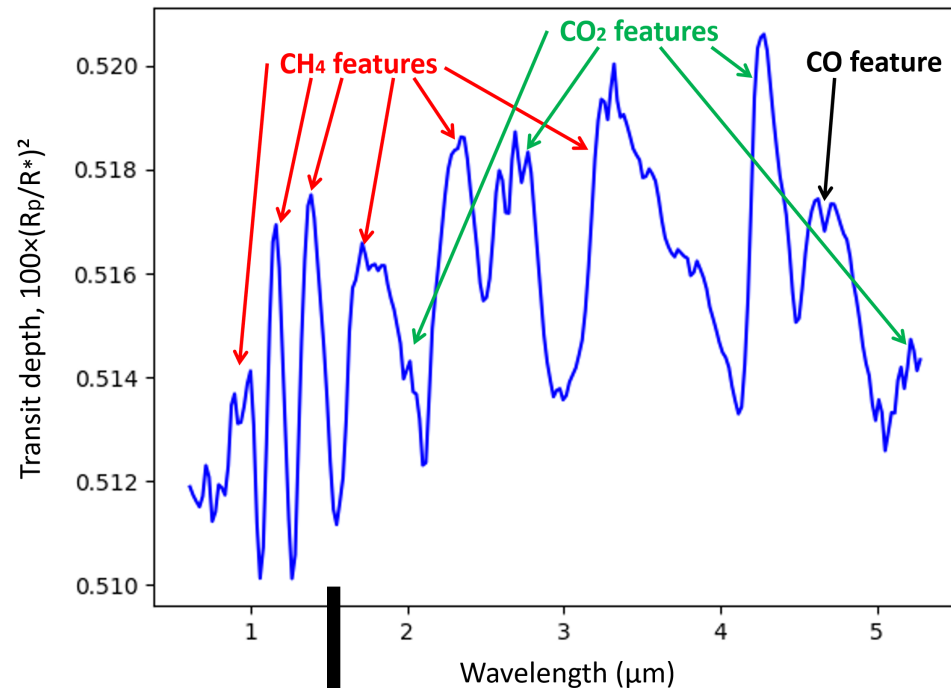
Forward model
NEMESIS

Compare to “true”
 properties



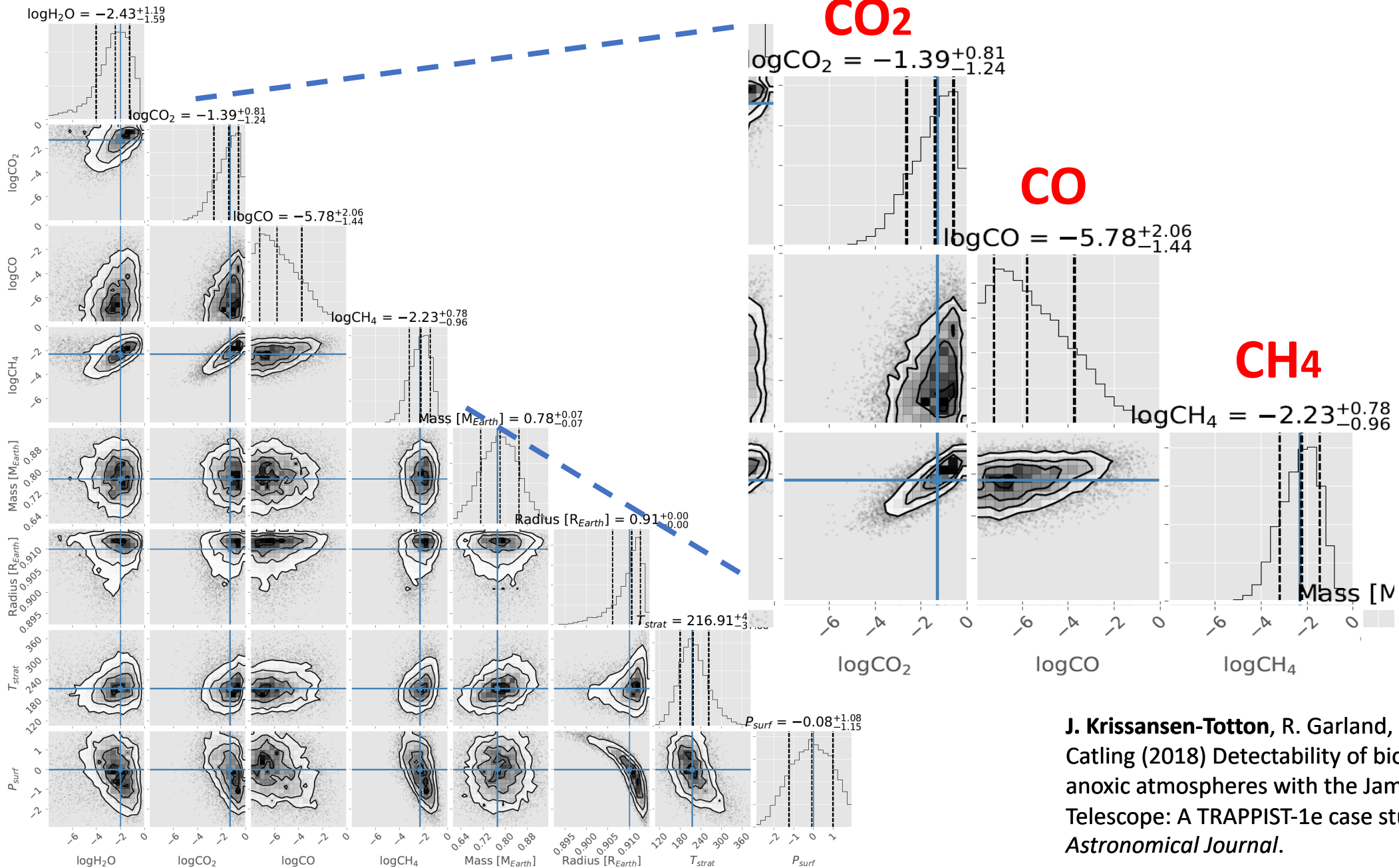
Mass = $0.81 \pm 0.2 M_{\text{Earth}}$
 Radius = $0.901 \pm 0.003 R_{\text{Earth}}$
 $p_{\text{CO}_2} = 3 \pm 2\%$
 $p_{\text{CH}_4} = 0.3 \pm 0.29\%$
 $p_{\text{CO}} = 0.001 \pm 0.0002\%$
 $p_{\text{H}_2\text{O}} = 0.5 \pm 0.48\%$
 $T_{\text{strat}} = 188 \pm 54 \text{ K}$
 $P_{\text{surf}} = 2.3 \pm 1.3 \text{ bar}$

JWST simulator
Solve inverse problem



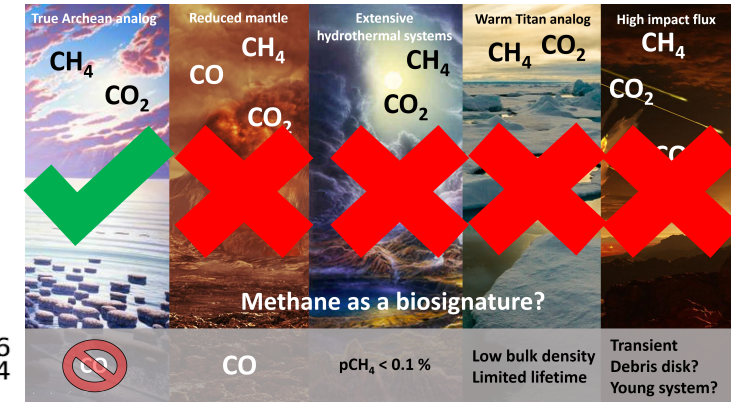
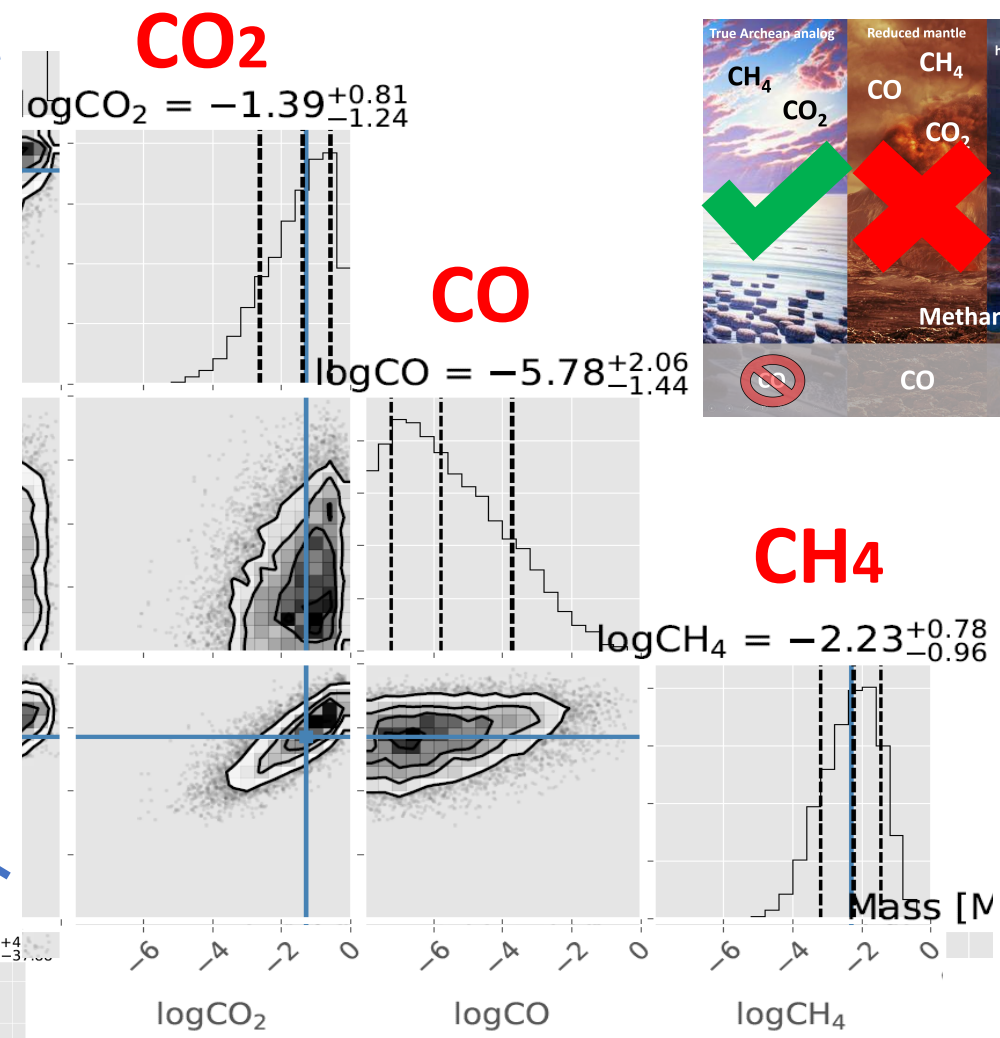
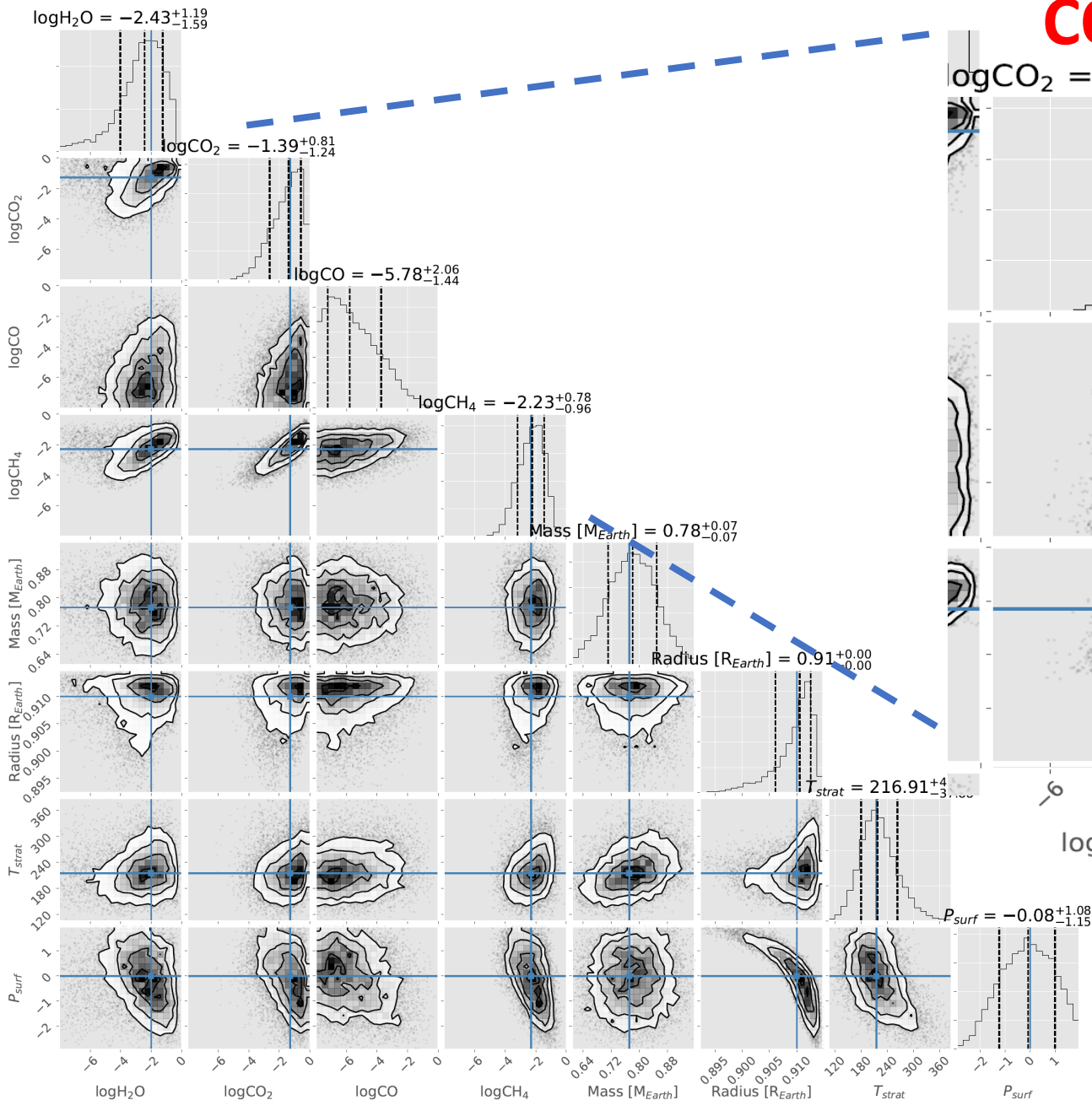
Instrument simulator: PandExo described
 in Batalha et al. (2017) PASP

Results: Observing TRAPPIST-1e with James Webb for 10 transits



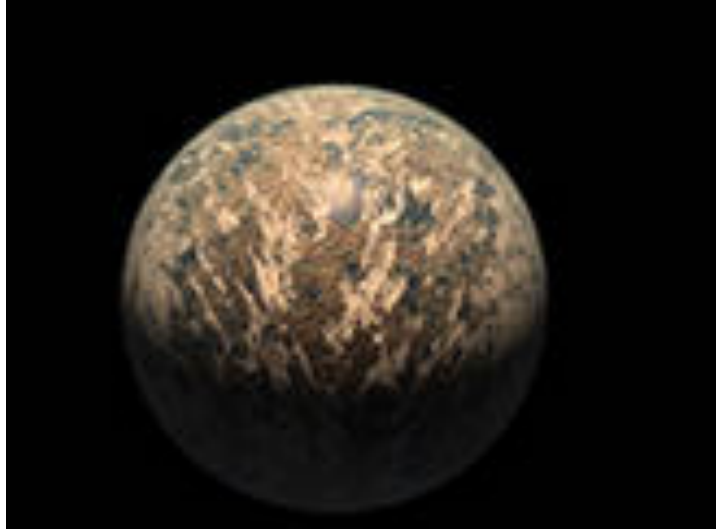
J. Krissansen-Totton, R. Garland, P. Irwin, D. C. Catling (2018) Detectability of biosignatures in anoxic atmospheres with the James Webb Space Telescope: A TRAPPIST-1e case study, *The Astronomical Journal*.

Results: Observing TRAPPIST-1e with James Webb for 10 transits

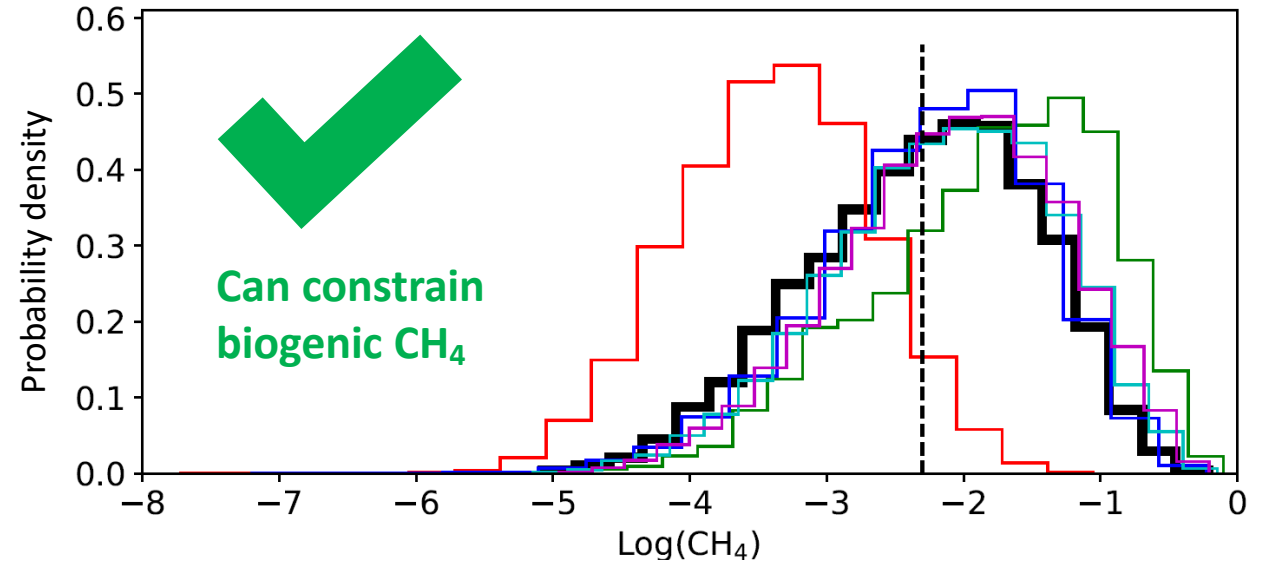


J. Krissansen-Totton, R. Garland, P. Irwin, D. C. Catling (2018) Detectability of biosignatures in anoxic atmospheres with the James Webb Space Telescope: A TRAPPIST-1e case study, *The Astronomical Journal*.

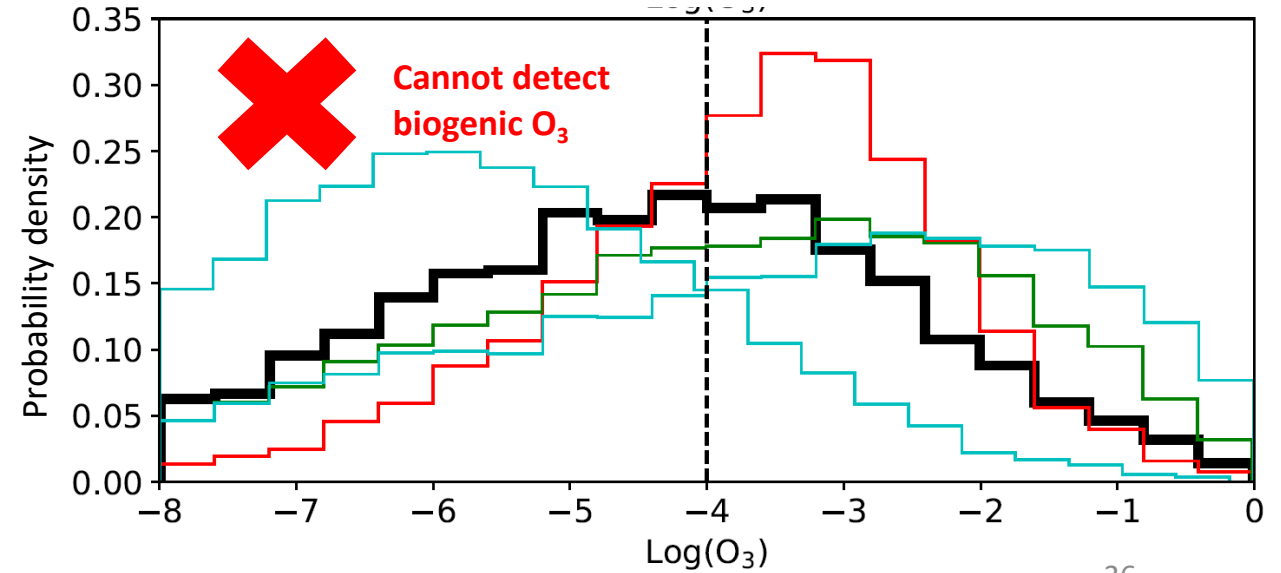
Is $\text{CH}_4 + \text{CO}_2$ biosignature easier to observe than O_3 ?



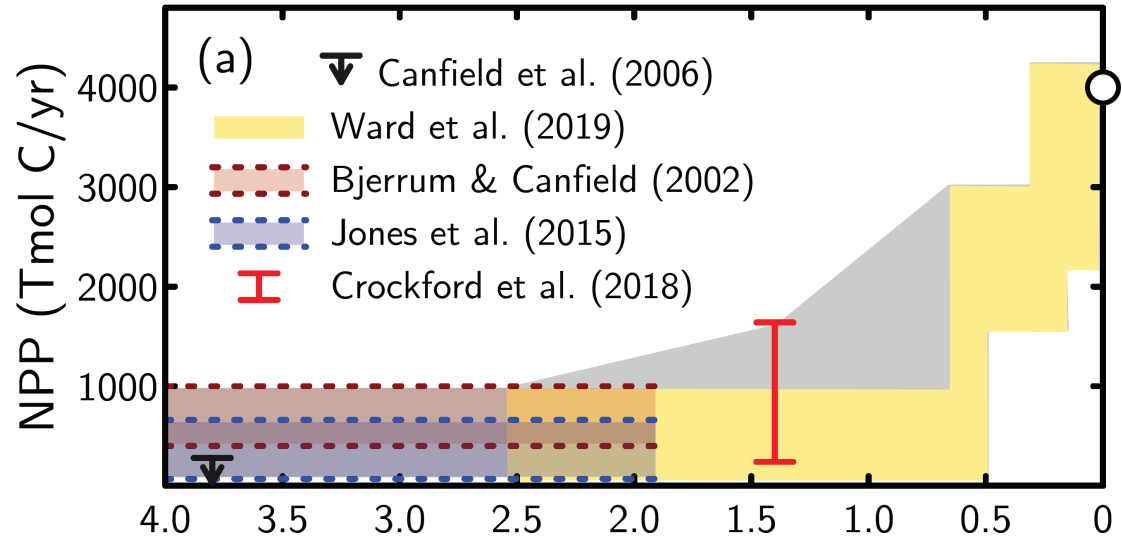
JWST
observing an
early Earth
analog



JWST
observing a
modern Earth
analog in
thermal IR

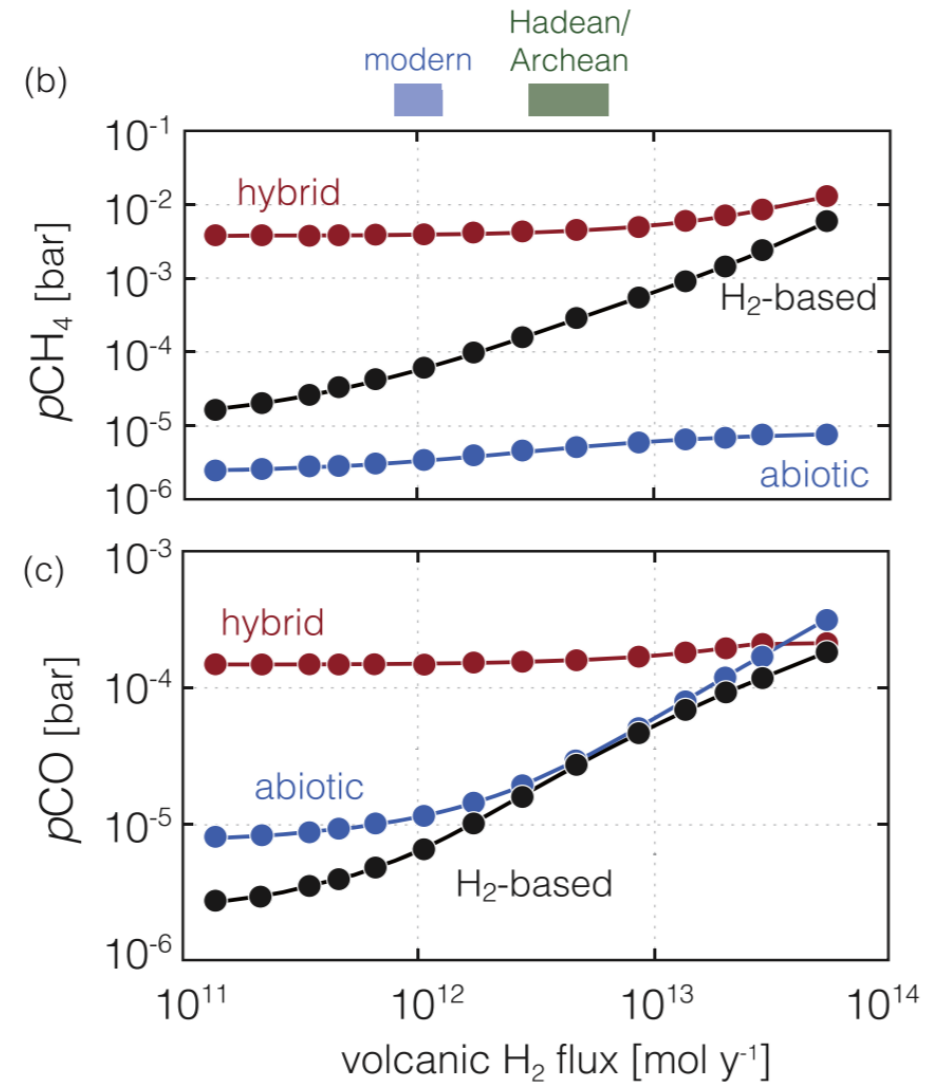
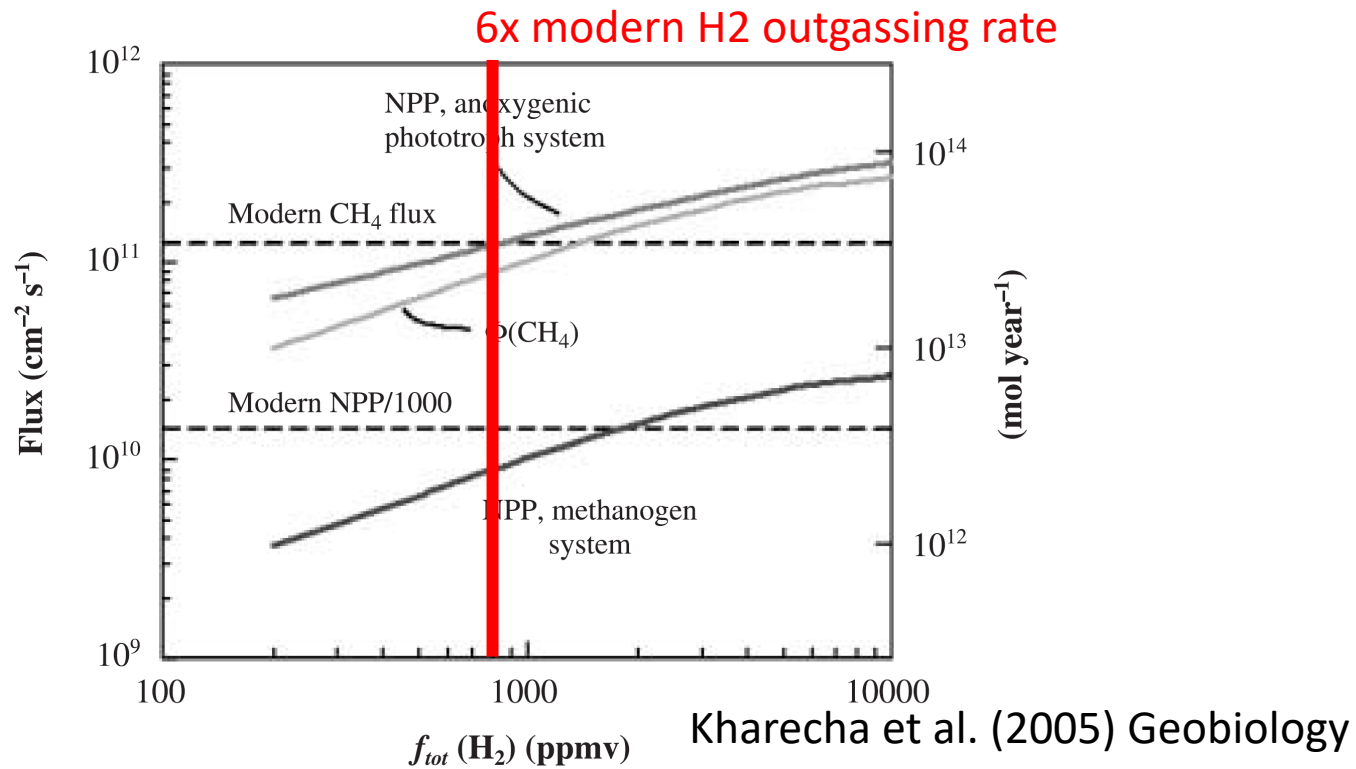


Reasons to be cautious...



Would low productivity preclude Archean-like CH₄+CO₂ biosignatures?

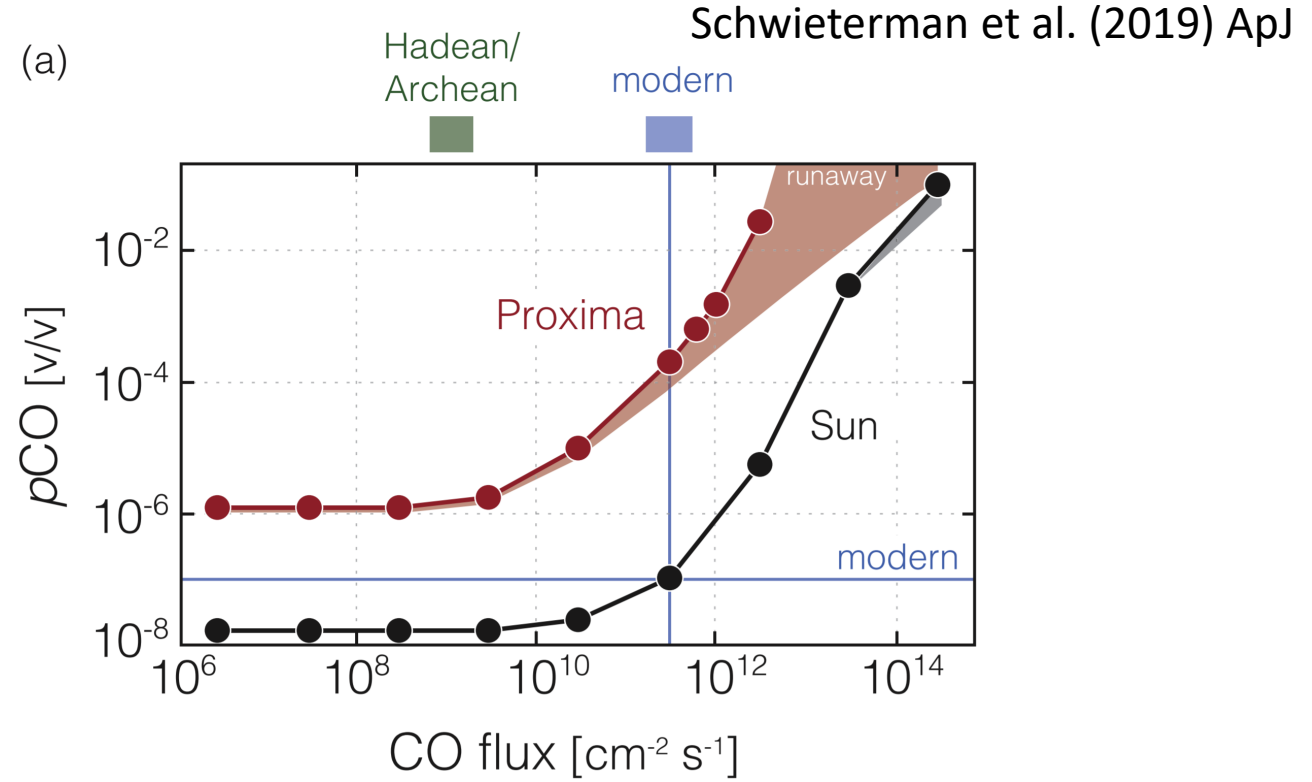
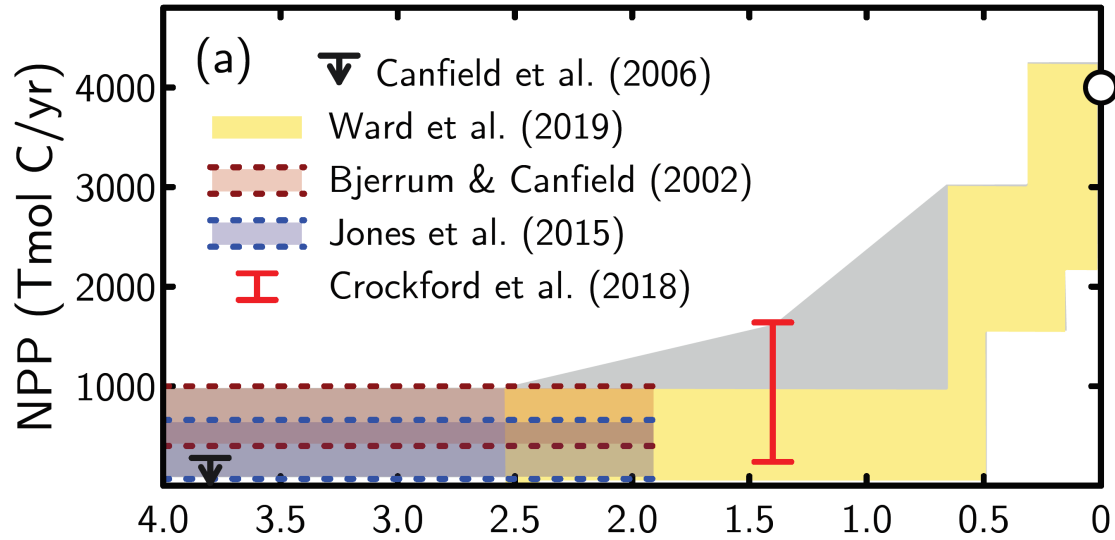
- If oxygenic photosynthesis has evolved -> **definitely not**
- If anoxygenic photosynthesis has evolved -> **probably not**
- If purely chemoautotrophic -> **depends on H₂ flux**



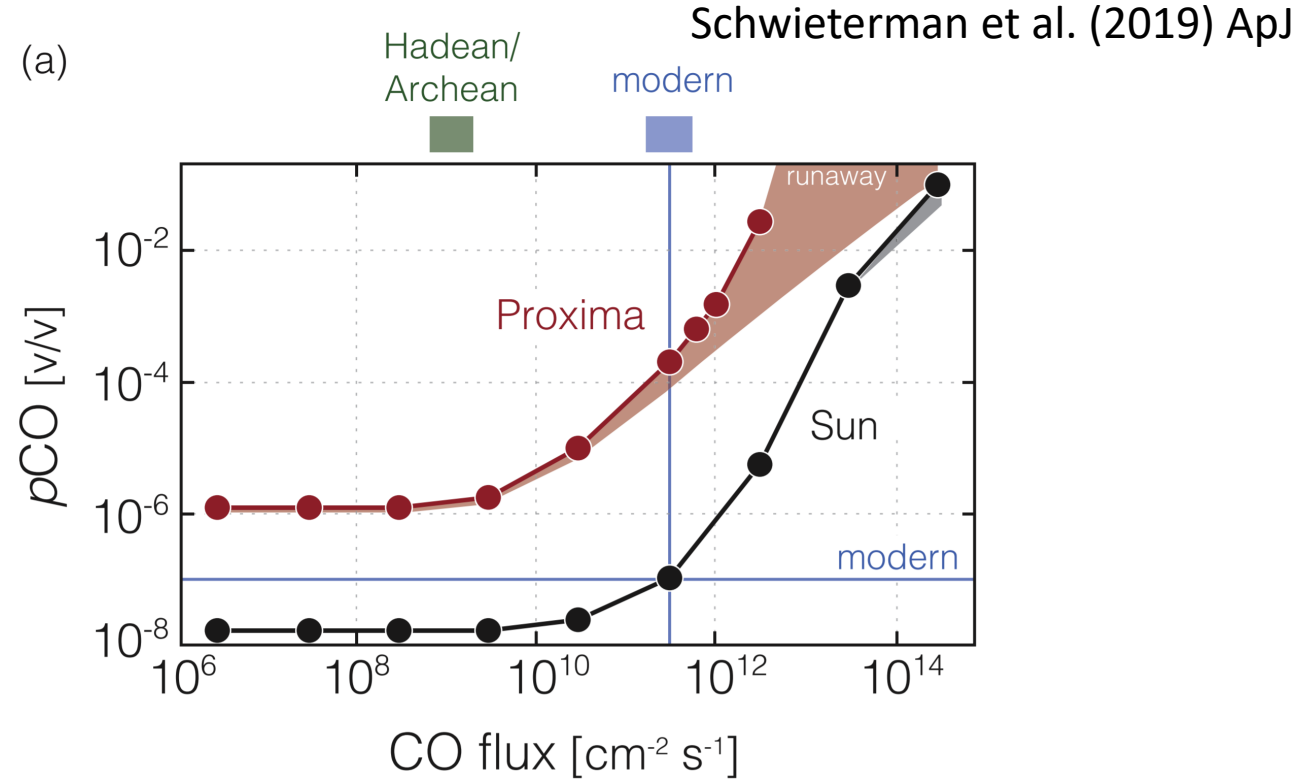
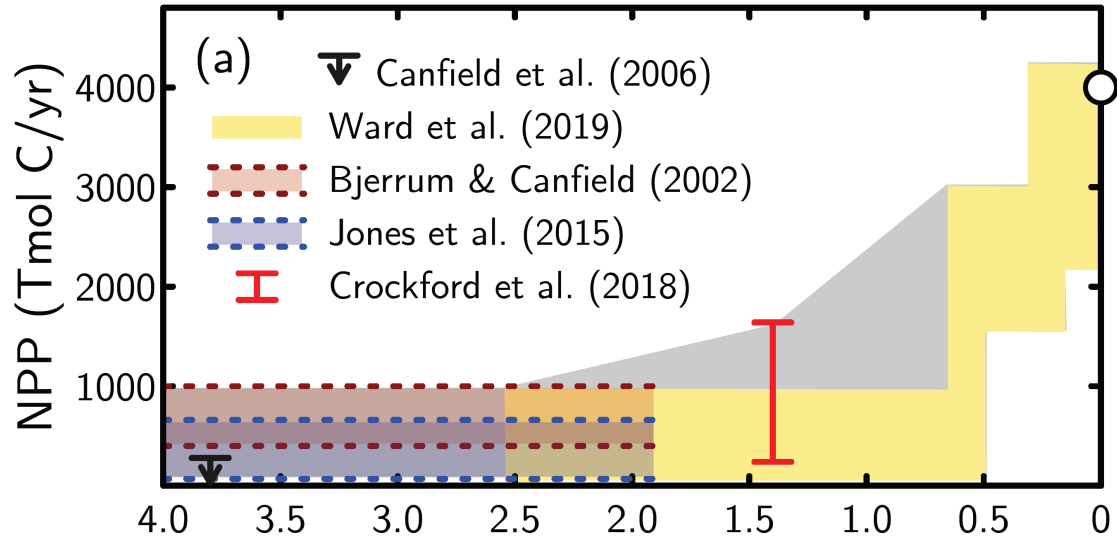
- Empirical evidence from Xe isotopes (Zahnle et al. 2019)

Schwieterman et al. (2019) ApJ

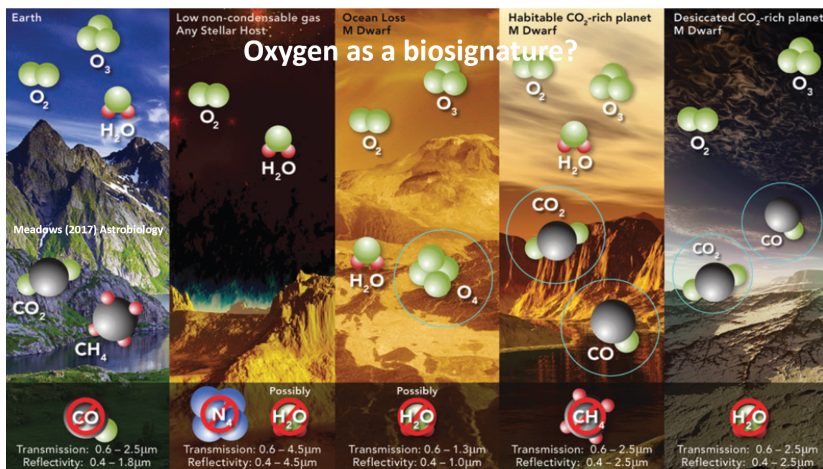
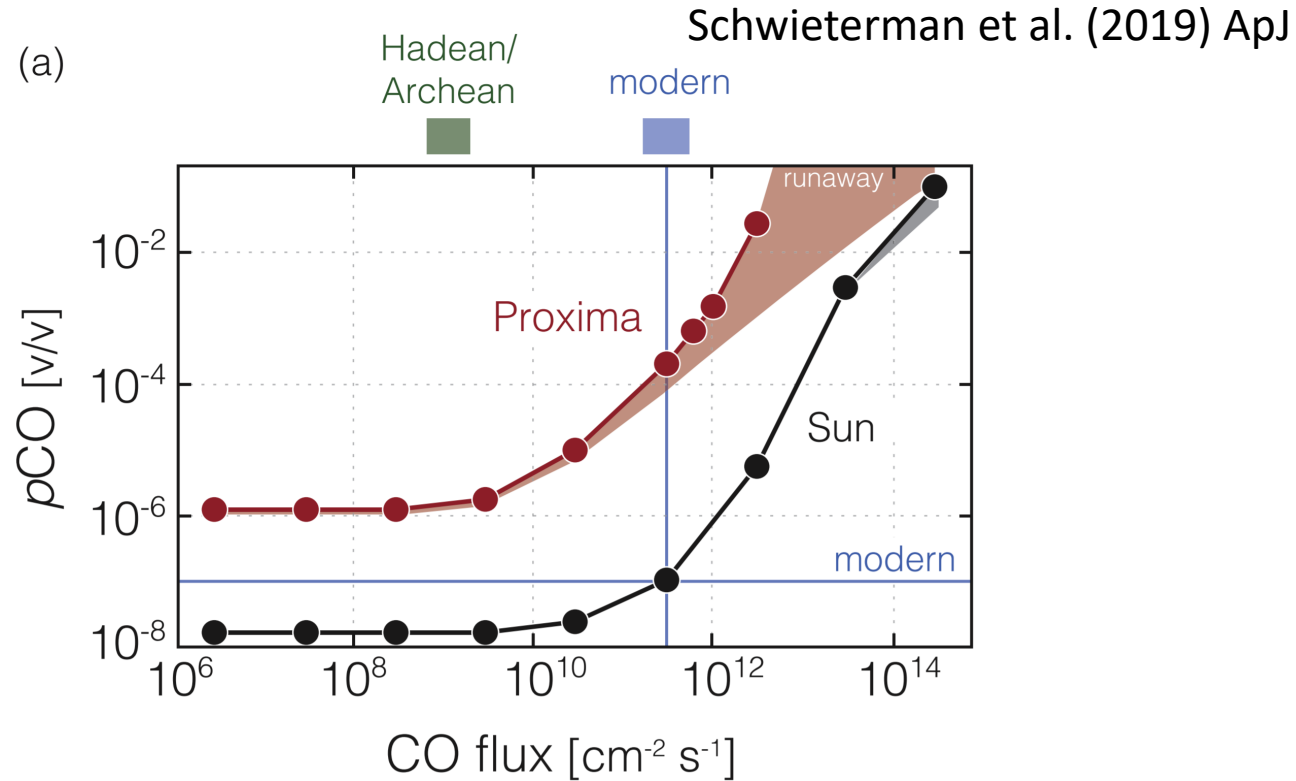
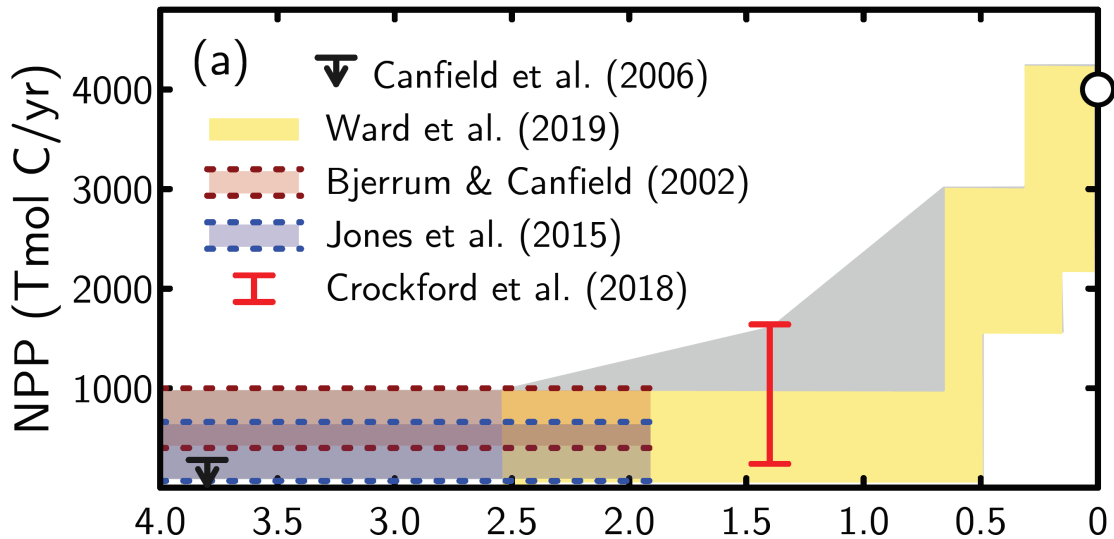
Reasons to be cautious...



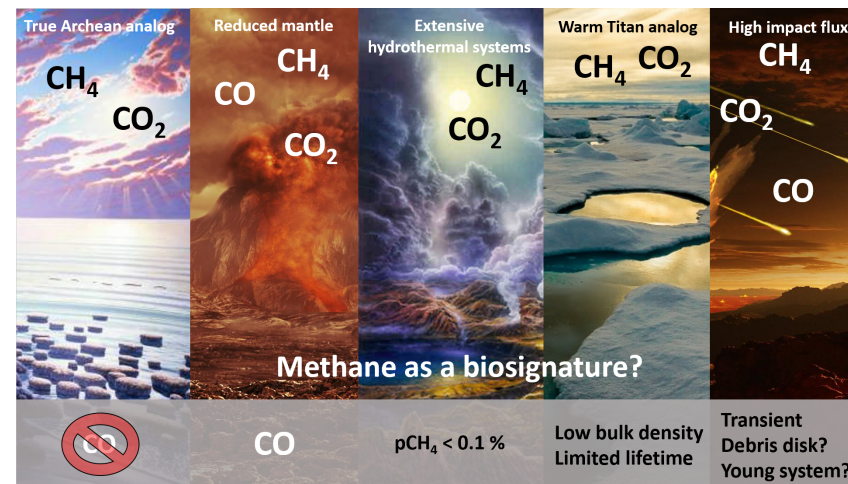
Reasons to be cautious...



Reasons to be cautious...

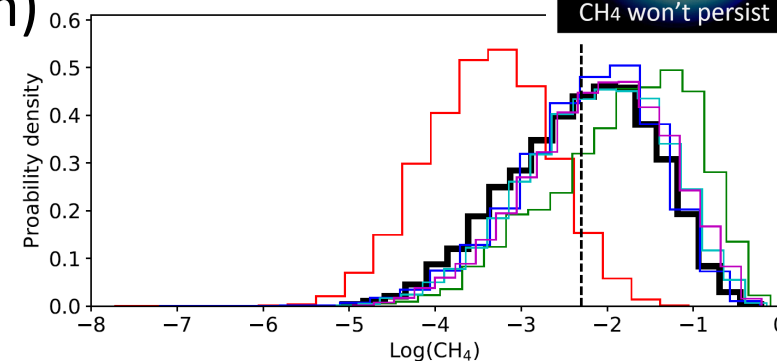
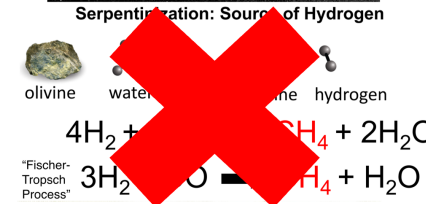
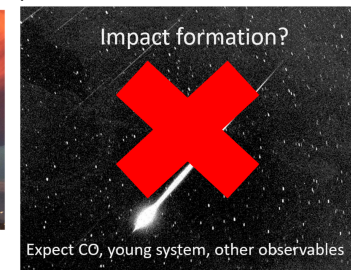
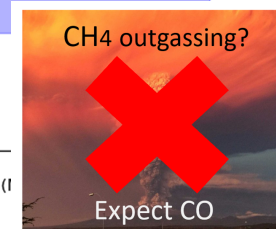
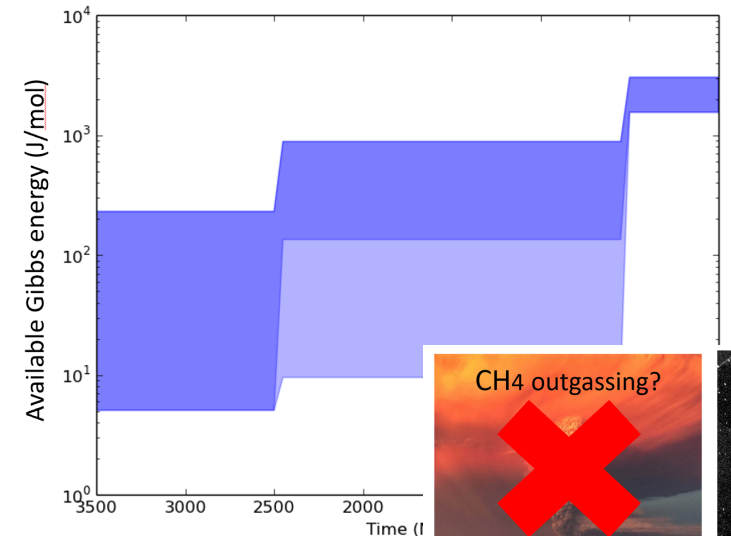


Vs.

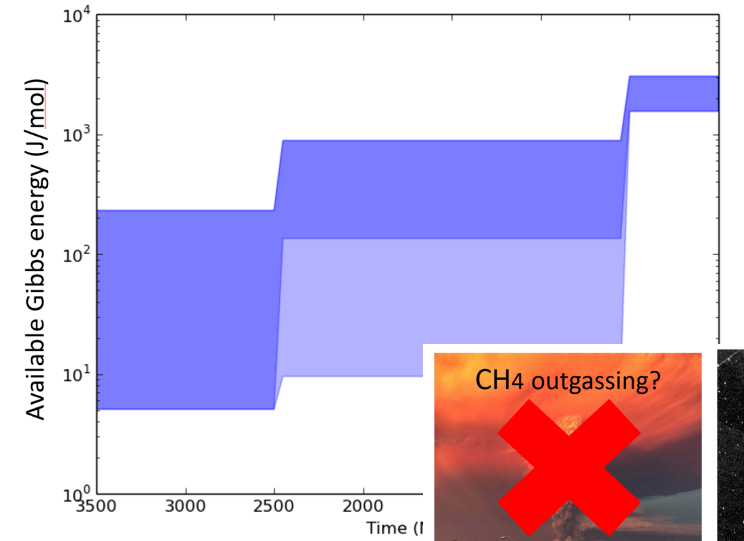
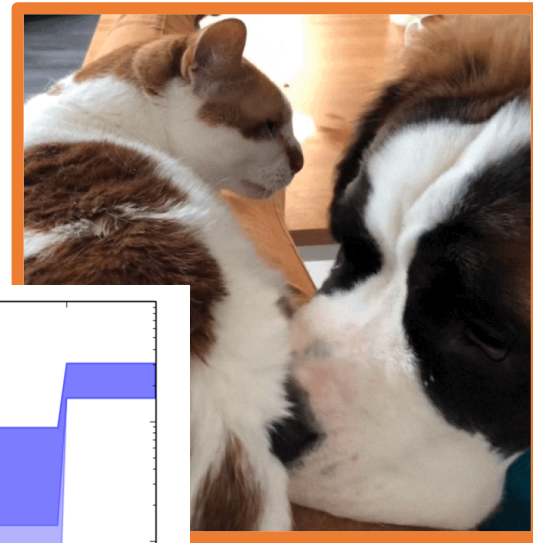


Conclusions

- Chemical disequilibrium is a useful conceptual framework for thinking about exoplanet biosignatures.
- The Earth has had a sizeable biogenic disequilibrium in its atmosphere-ocean system throughout its history.
- The coexistence of carbon dioxide and abundant methane is difficult to explain with known non-biological processes and could be a compelling exoplanet biosignature (as it was in the Archean)
- This biosignature combination is potentially detectable with JWST



Thanks!



CH₄ outgassing?

 Expect CO

Impact formation?

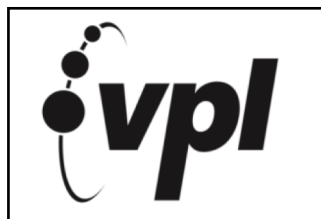
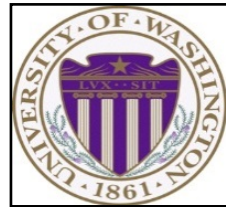
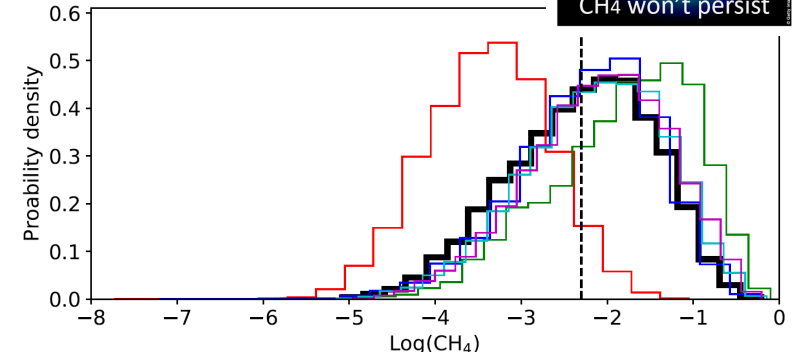
 Expect CO, young system, other observables

Warm Titan?

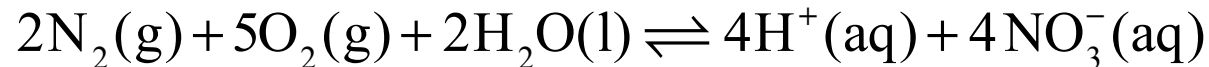
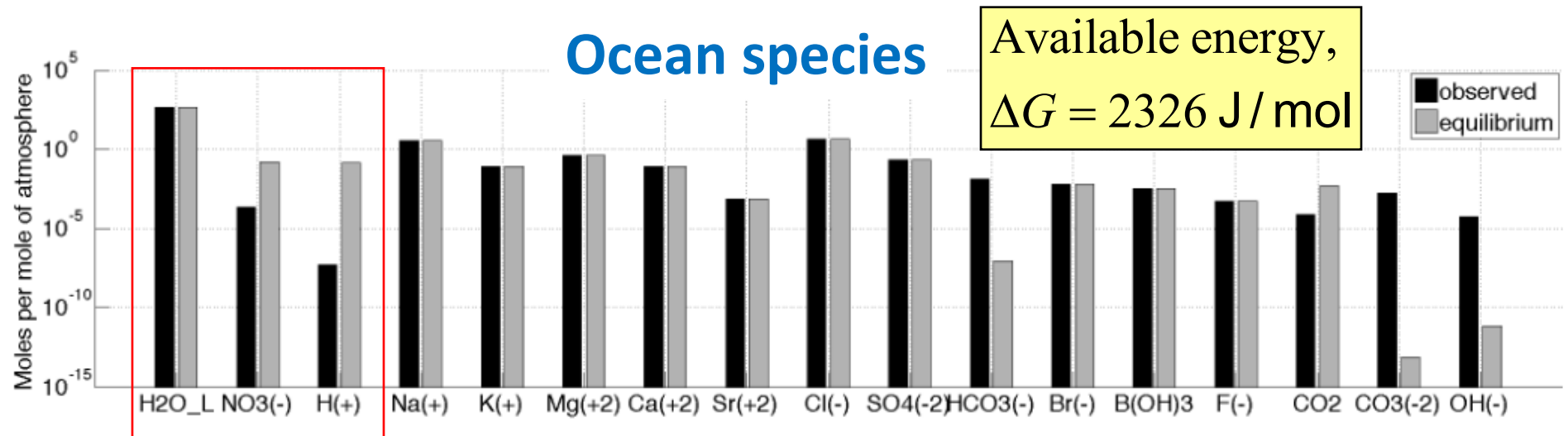
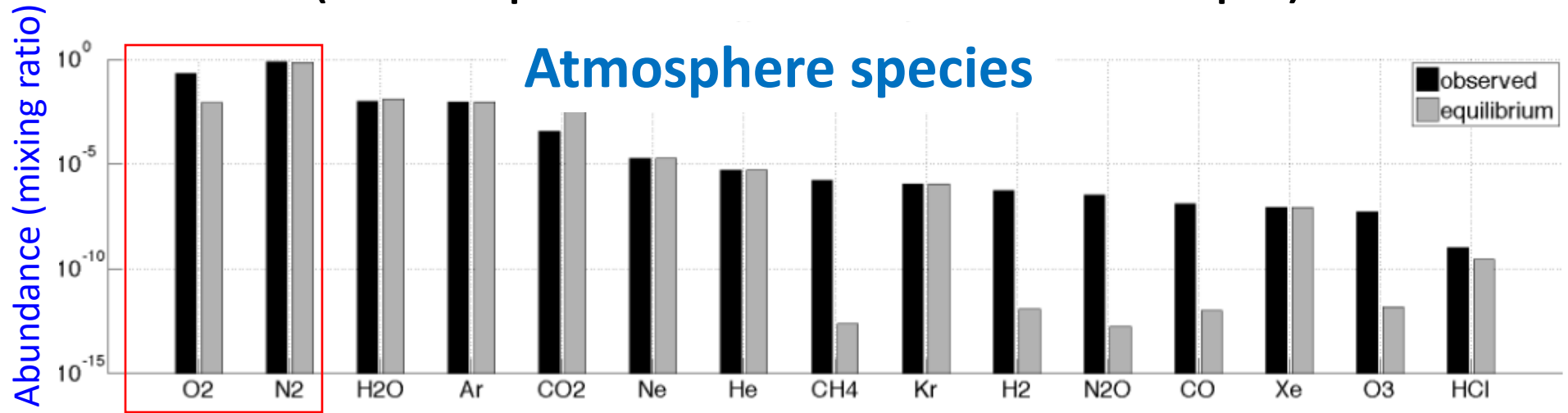
 CH₄ won't persist

Serpentinization: Source of Hydrogen

 $4\text{H}_2 + \text{CO}_2 \rightarrow \text{CH}_4 + 2\text{H}_2\text{O}$
 $3\text{H}_2 + \text{CO} \rightarrow \text{CH}_4 + \text{H}_2\text{O}$
Fischer-Tropsch Process



Earth (atmosphere-ocean fluid envelope)

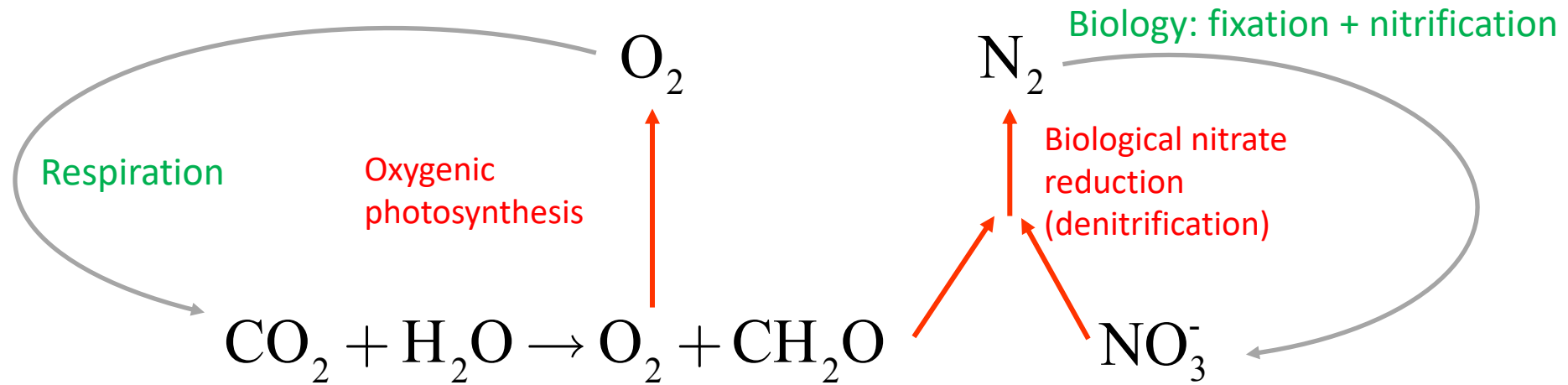


Gilbert Lewis (1923): “starting with air and water...nitric acid should form. It is to be hoped that nature will not discover a catalyst for this reaction, which would... turn the oceans into dilute nitric acid”.

Sensitivity test

		Available energy, Φ (J/mol)
Temperature	T= 273.15 K	1634.78
	T= 288.15 K	2325.76
	T= 298.15 K	2824.48
Pressure	0.1 bar	1354.20
	1.013 bar	2325.76
	10 bar	3891.96
	1000 bar	6878.35
Ocean pH	2	1983.28
	4	2314.26
	6	2325.71
	8.187 (Earth)	2325.76
	12	2325.65
Salinity	0 mol/kg	2290.01
	1.1 mol/kg	2325.76
	11.1 mol/kg	2276.40
Ocean volume	0.1 Earth ocean	413.62
	0.5 Earth ocean	1442.95
	1 Earth ocean	2325.76
	2 Earth oceans	4188.27
	10 Earth oceans	8956.34
	50 Earth oceans	12626.22

What does disequilibrium mean?



- In the absence of life, lightning would convert all oxygen to nitrate within a few million years.
- Thus the coexistence of N_2 , O_2 and liquid water is a more compelling biosignature than O_2 alone.

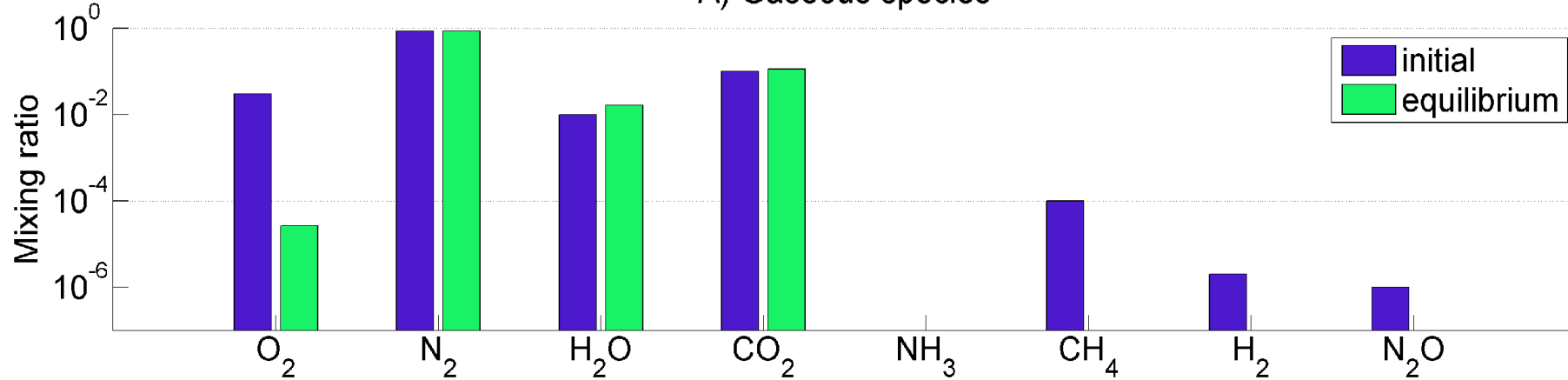
Krissansen-Totton, J., D. S. Bergsman, and D. C. Catling (2016), On detecting biospheres from chemical thermodynamic disequilibrium in planetary atmospheres, *Astrobiology*, 16(1), 39-67.

	Archean range			Proterozoic range		
Atmospheric species	Mixing ratio		Reference/Explanation	Mixing ratio		Reference/Explanation
	Min. diseq.	Max diseq.		Min. diseq.	Max diseq.	
N ₂ (g)	0.98	0.5	Mixing ratios sum to 1. But see also Som et al. 2016; Marty et al. 2013	0.99	0.86	Mixing ratios sum to 1
O ₂ (g)	1×10 ⁻¹⁰	2×10 ⁻⁷	Zahnle et al. 2006	0.0001	0.03	Lyons et al. 2014
CH ₄ (g)	0.0001	0.01	Kasting and Brown 1998	3×10 ⁻⁶	1×10 ⁻⁴	Olson et al. 2016
CO ₂ (g)	0.001	0.5	Sheldon 2006; Kanzaki 2015	0.0001	0.1	Sheldon 2006; Kanzaki 2015
H ₂ (g)	0	0.0001	Kharecha et al. 2005	0	2×10 ⁻⁶	Segura et al. 2003
N ₂ O(g)	0	0	No denitrification	0	1×10 ⁻⁶	Roberson et al. 2011
NH ₃ (g)	0	1×10 ⁻⁹	Pavlov et al. 2001	0	0	NA
O ₃ (g)	0	0	Negligible	0	0	Negligible
CO(g)	0	0.001	Kharecha et al. 2005	0	2×10 ⁻⁷	Segura et al. 2003
Ocean species	Molality (mmoles/kg)		Reference/Explanation	Molality (mmoles/kg)		Reference/Explanation
	Min. diseq.	Max diseq.		Min. diseq.	Max diseq.	
Na(+)	550	586	Charge balance	547	549	Charge balance
Cl(-)	546	546	Modern value	546	546	Modern value
SO ₄ (2-)	0	0.2	Olson	0.25	5	Olson
H ₂ S	0	0.004	Euxinic oceans like Black Sea	0	0.004	Euxinic oceans like Black Sea
NH ₄ (+)	0	0.050	Set by P (N fixation evolved early)	0	0.050	Set by P
NO ₃ (-)	0	0	Anoxic bulk ocean	0	0	Anoxic bulk ocean
ALK	4	40	Krissansen-Totton*	1.0	3.0	Krissansen-Totton*
pH	5.0	8.0	Carbon chemistry equilibrium**	6.0	8.0	Carbon chemistry equilibrium**

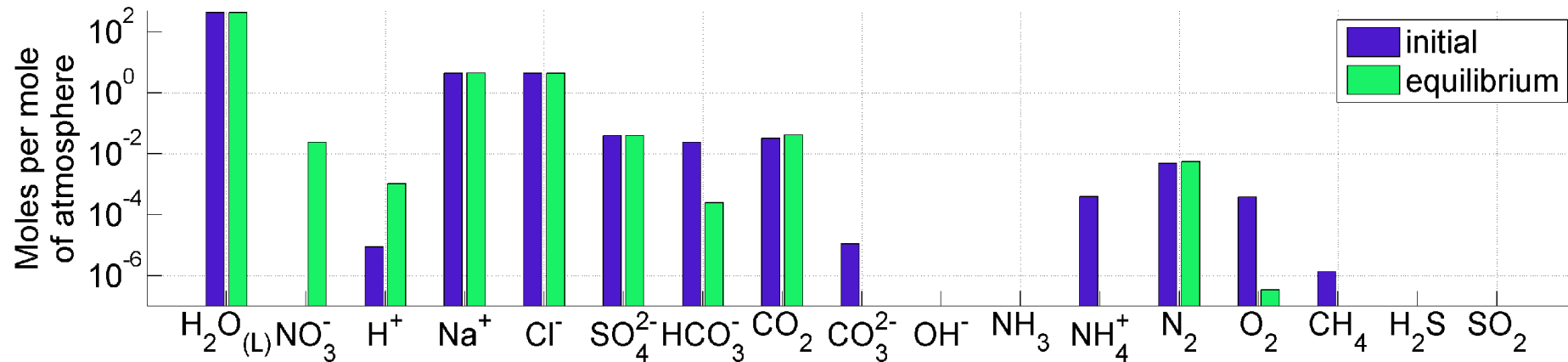
Proterozoic atmosphere-ocean disequilibrium



A) Gaseous species



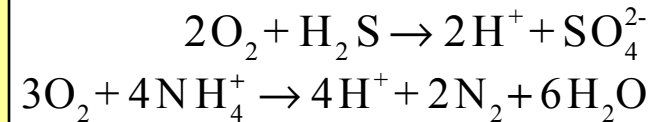
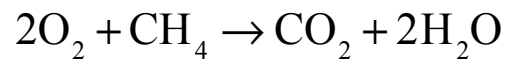
B) Aqueous species



Available energy,

$$\Delta G = G_{(T,P)}(\text{obser}) - G_{(T,P)}(\text{equil})$$

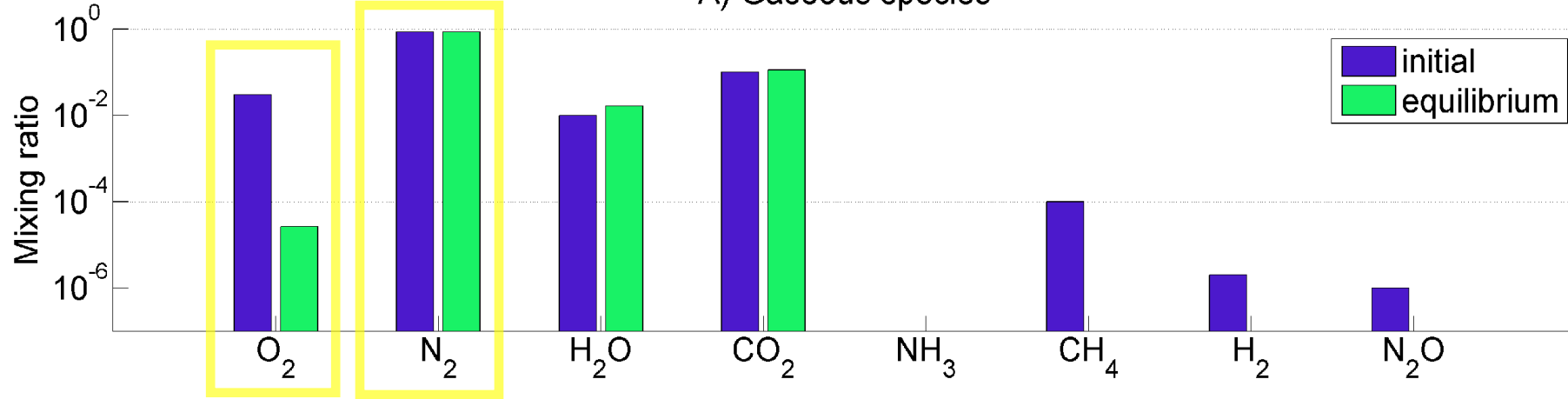
$$\Delta G = 884 \text{ J/mol}$$



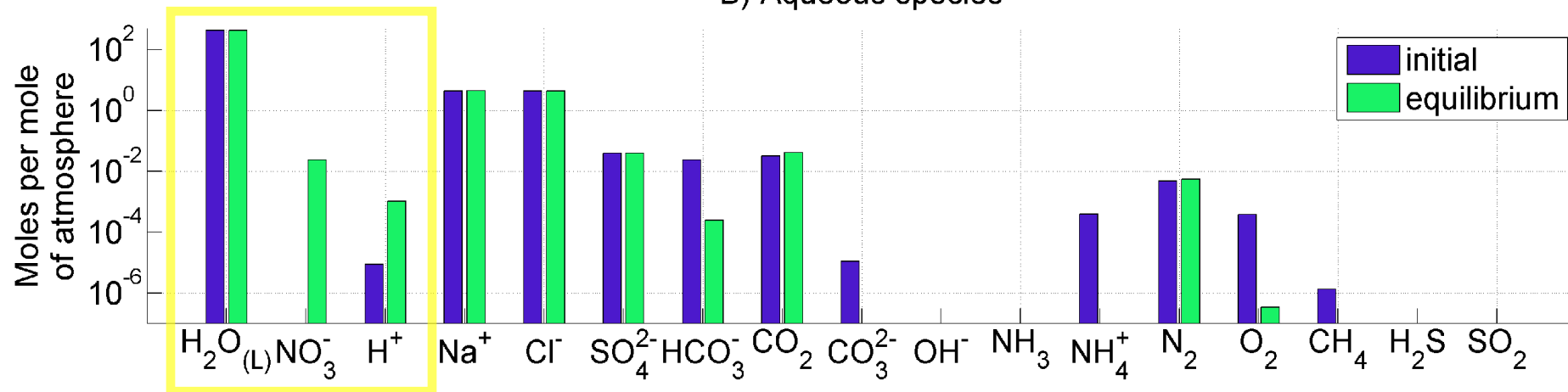
Proterozoic atmosphere-ocean disequilibrium



A) Gaseous species



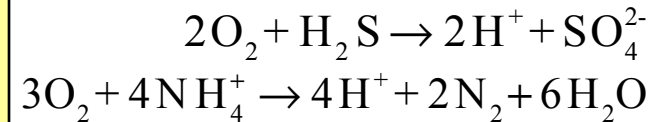
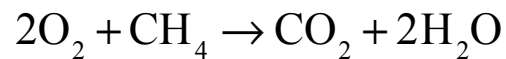
B) Aqueous species



Available energy,

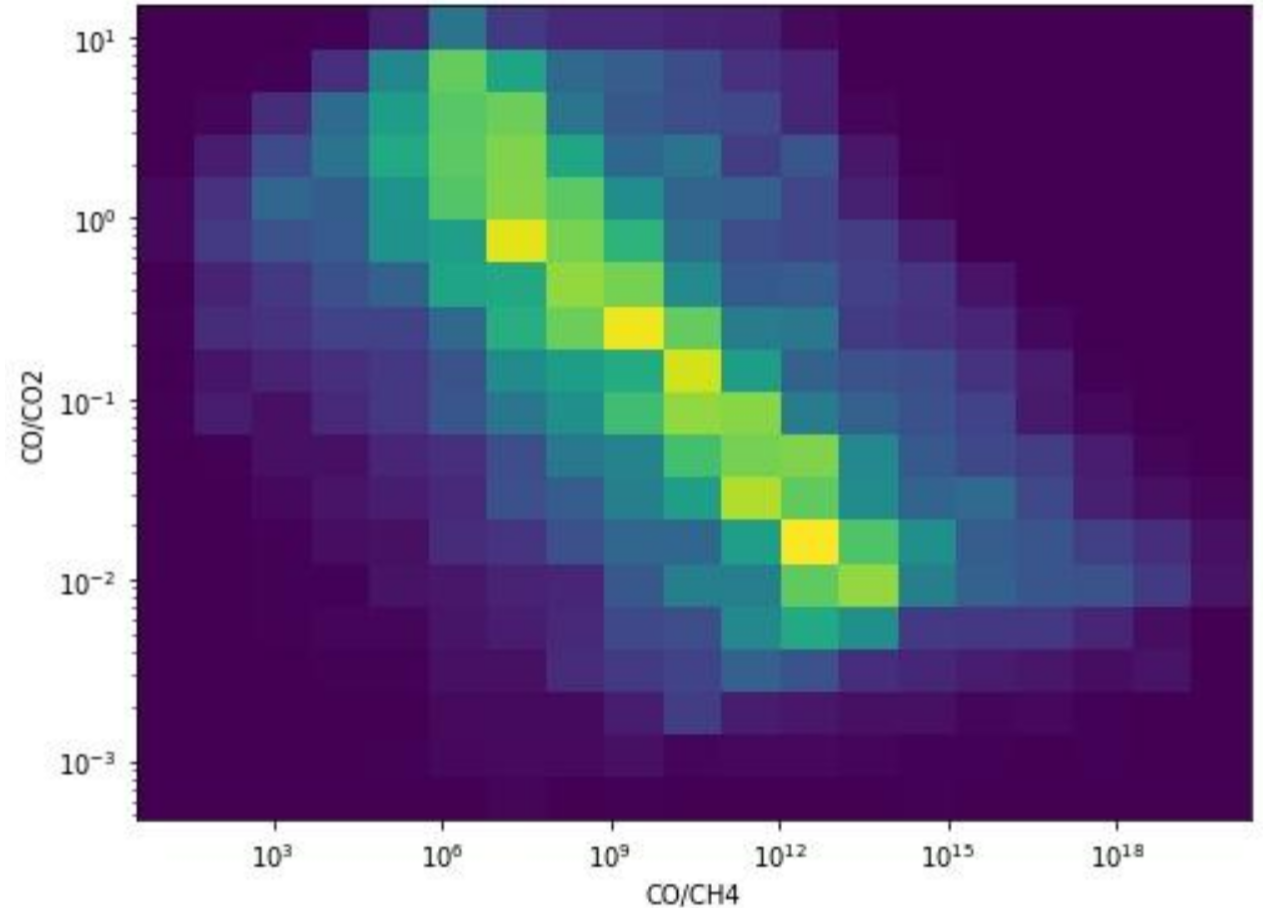
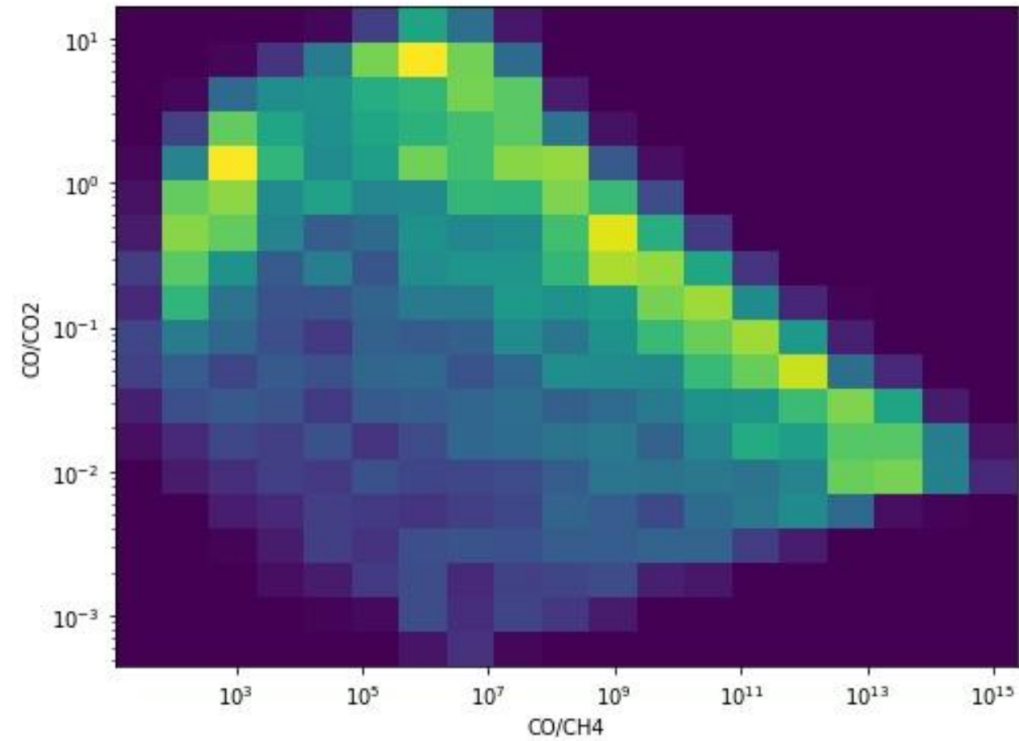
$$\Delta G = G_{(T,P)}(\text{obser}) - G_{(T,P)}(\text{equil})$$

$$\Delta G = 884 \text{ J/mol}$$



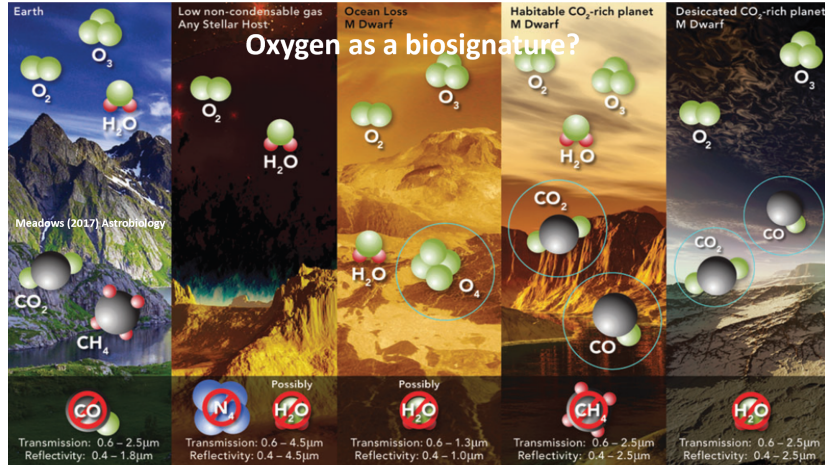
Ocean worlds, 100-1000 bar

Submarine + subaerial outgassing,
0.01-100 bar

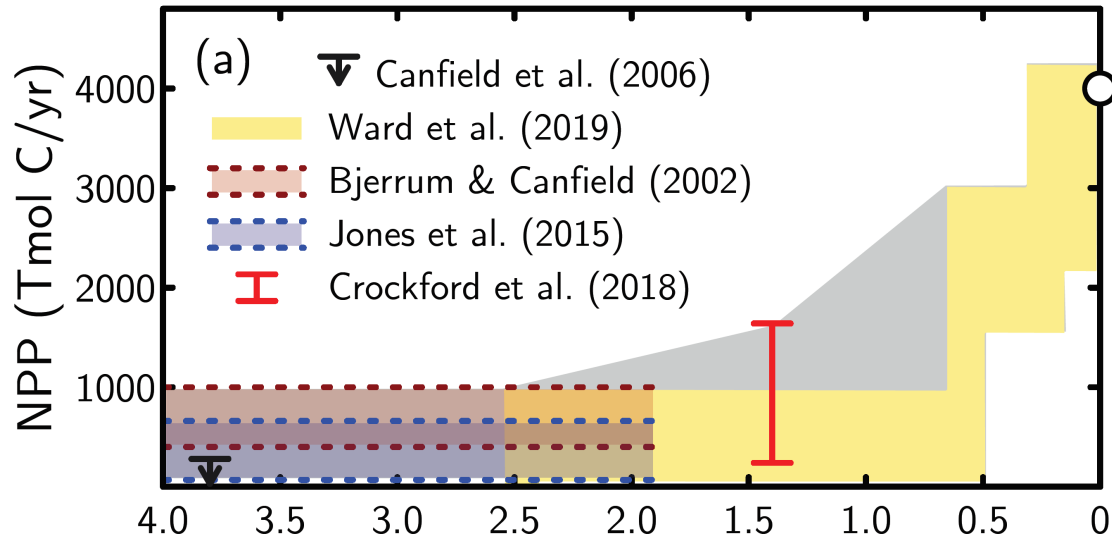
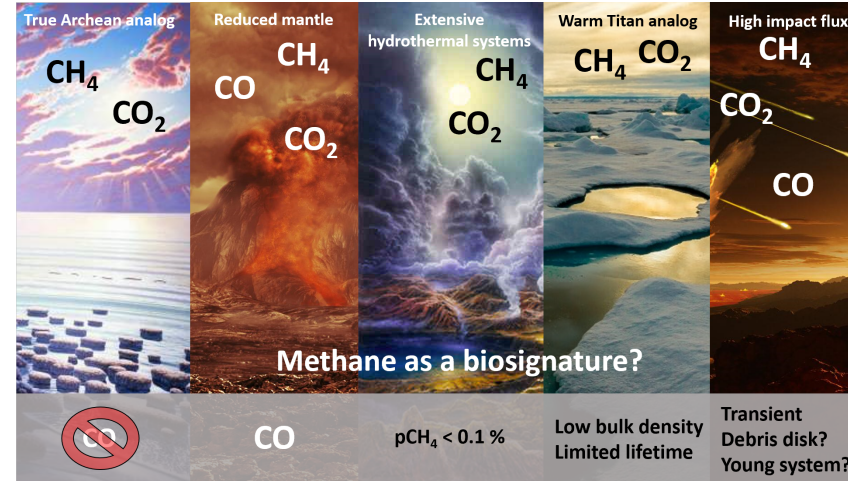


fO_2 : -4.5 to 2.0 FMQ, magma chamber 800-1200 bar, also varied CO_2 and H_2O composition host rock but secondary effect.

Reasons to be cautious...



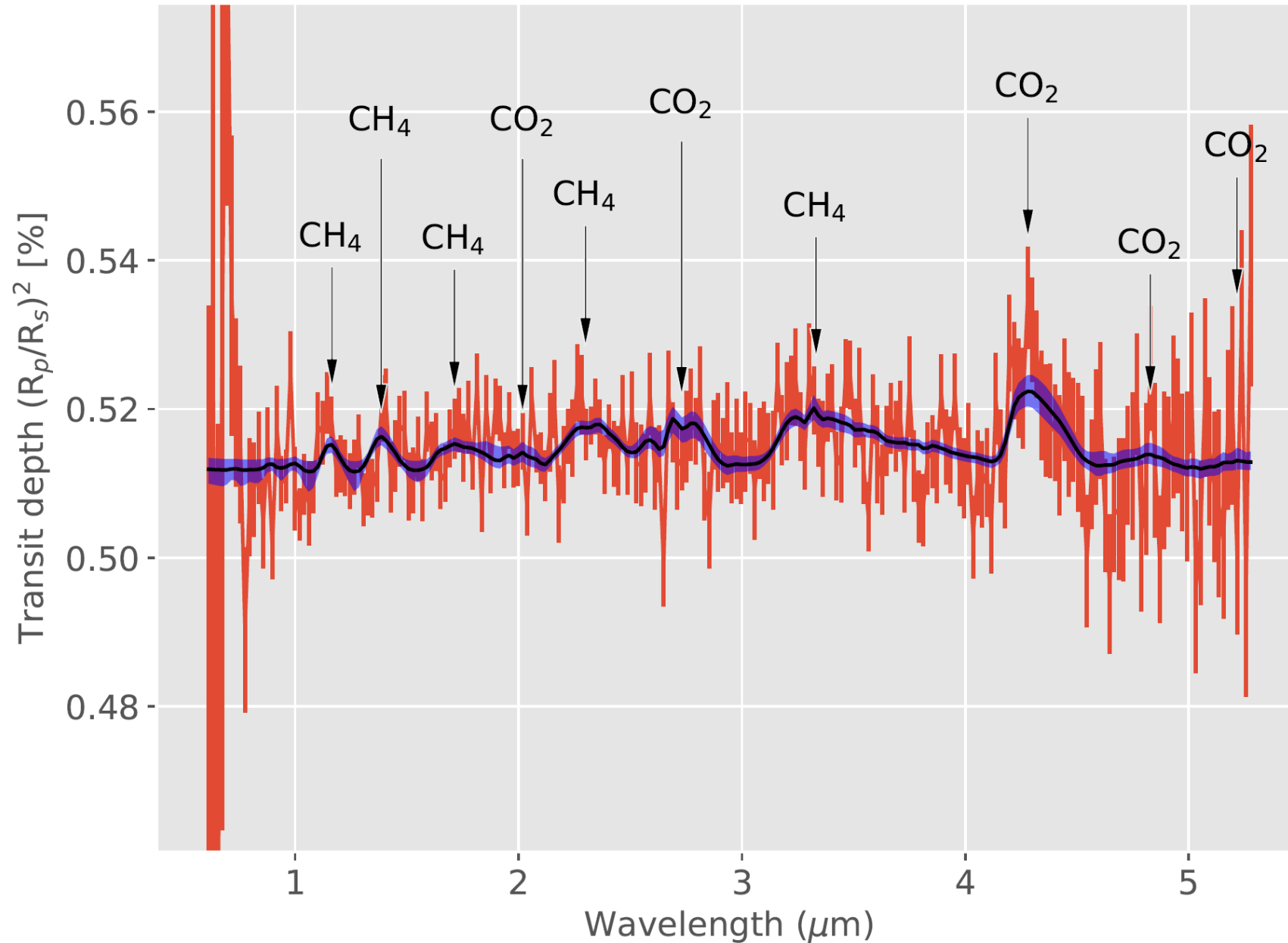
Vs.



Failures of imagination

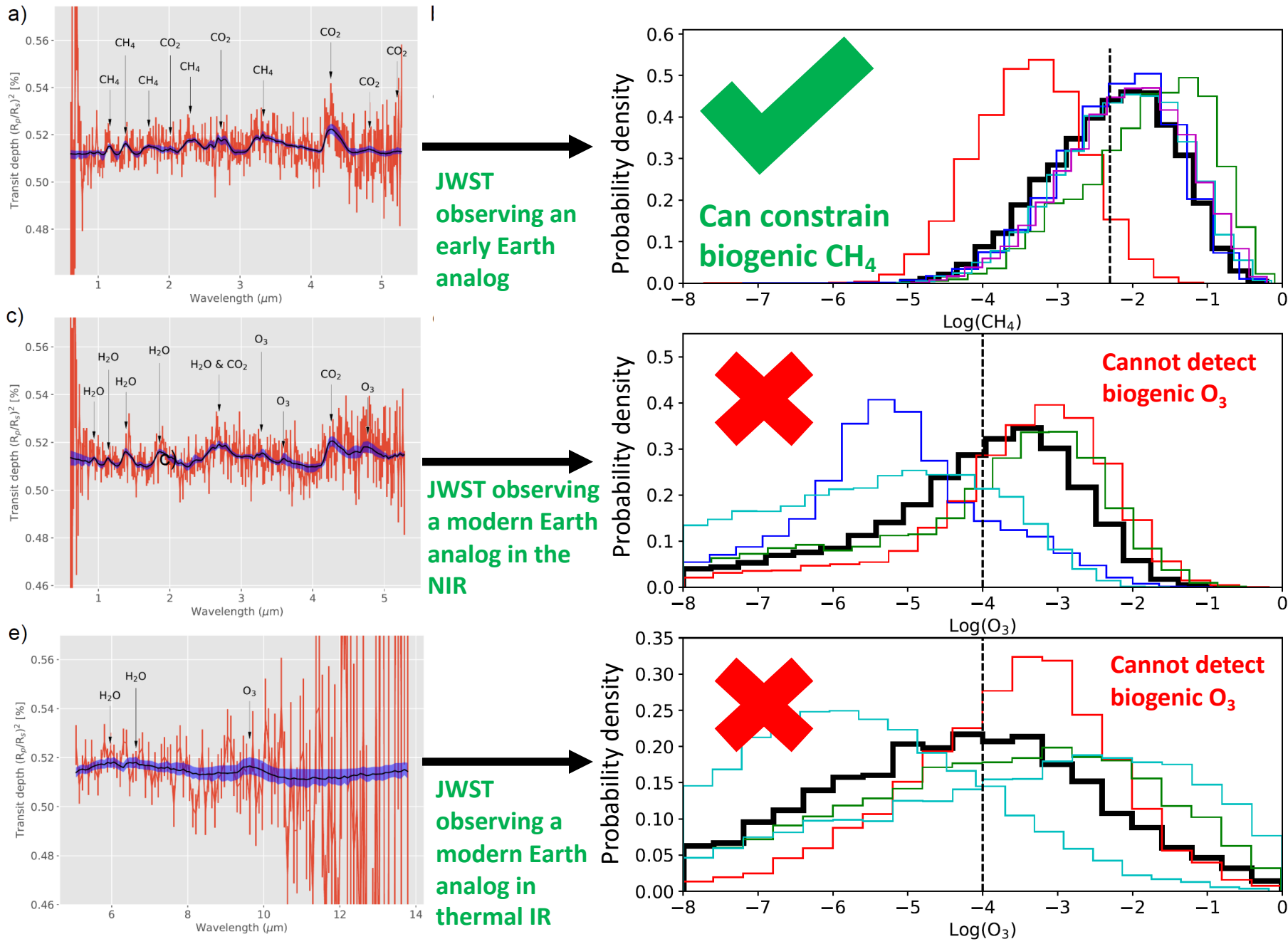


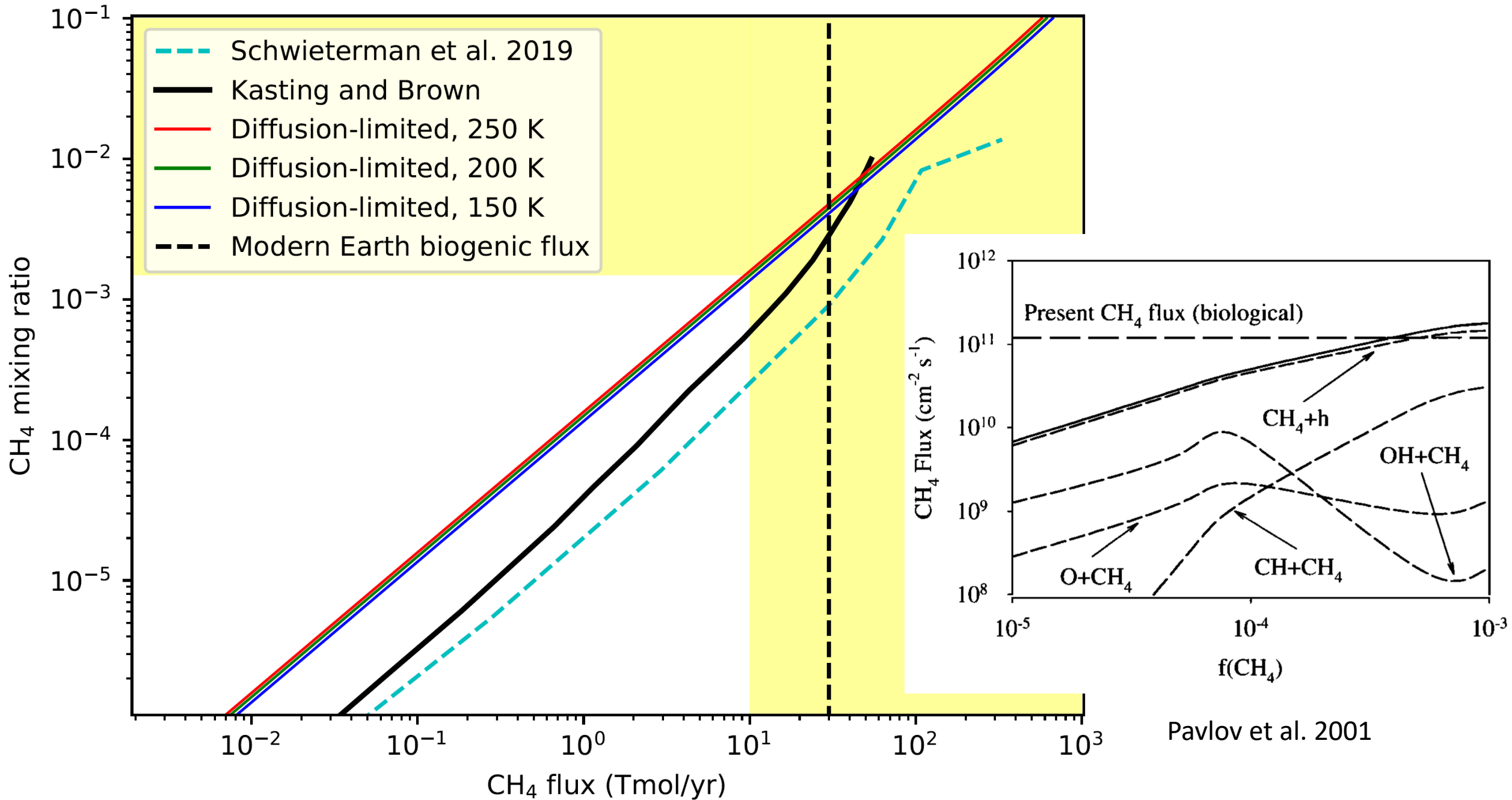
TRAPPIST-1e, NIRSpec prism on JWST, 10 transits



J. Krissansen-Totton, R. Garland, P. Irwin, D. C. Catling (2018) Detectability of biosignatures in anoxic atmospheres with the James Webb Space Telescope: A TRAPPIST-1e case study, *The Astronomical Journal*.

Is CH₄+CO₂ biosignature easier to observe than O₃?





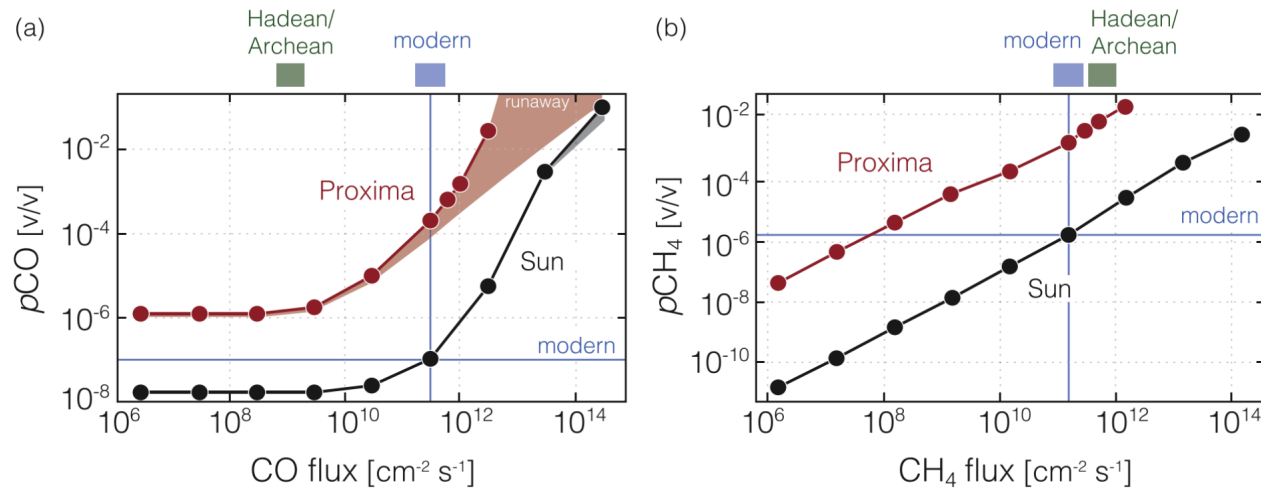
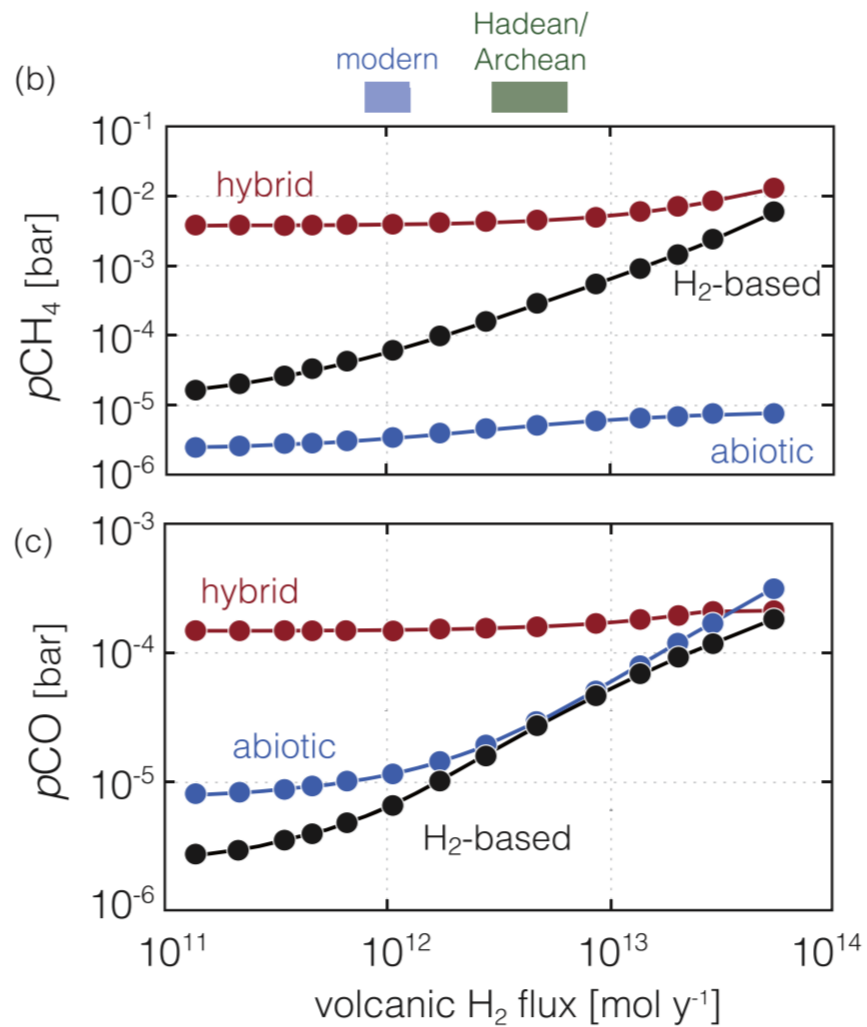
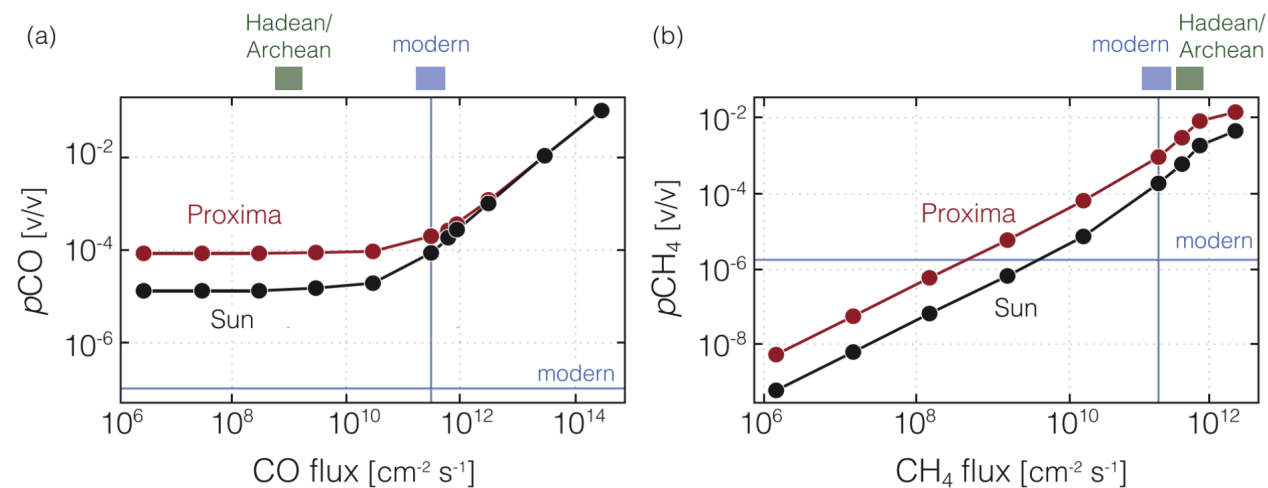


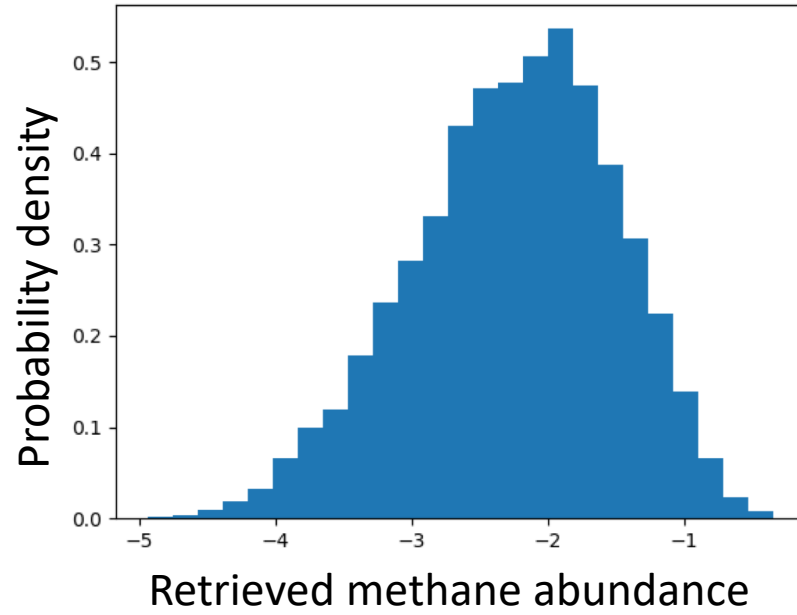
Figure 2. Atmospheric mixing ratios of CO (a) and CH_4 (b) as a function of surface molecular fluxes for oxygen-rich, modern Earth-like atmospheres. Results are shown for both the Sun (black) and the M-star Proxima Centauri (red). Horizontal line shows mixing ratios for each gas in the modern atmosphere, while modern and estimated Hadean/Archean fluxes of CO and CH_4 are shown above each panel. The shaded ranges in (a) depict results for models assuming only an upward molecular flux of CO (upper end) and assuming both an upward molecular flux and the maximum possible deposition velocity of $1.2 \times 10^{-4} \text{ cm s}^{-1}$ (lower end). Note that for the highest surface fluxes around Proxima Centauri the atmosphere enters a “CO runaway.” Atmospheric $p\text{CO}_2$ is fixed at 360 ppm ($\sim 3.6 \times 10^{-4}$ bar) in all calculations.



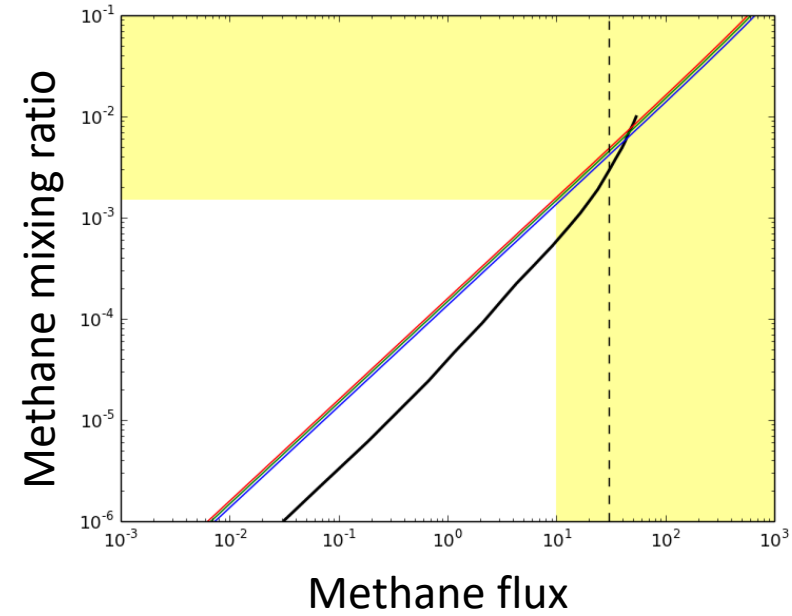
Archean sensitivity test

		Available Gibbs free energy (J/mol)	
		<u>Matlab</u>	ASPEN
Temperature, K	273.15	369	359
	288.15	234	232
	298.15	220	204
<u>Ocean alkalinity</u> , mmol/kg (pH)	4 (pH=5.4)	363	359
	40 (pH=6.4)	234	232
	200 (pH=7.1)	152	181
Ocean salinity, relative to modern	0.1	199	198
	1	234	232
	10	249	223
Ocean volume, relative to modern	0.1	89	89
	0.5	188	187
	1	234	232
	2	279	277
	10	405	401
	50	701	682
<u>Atmospheric pressure</u> , <u>atm</u>	0.1	125	141
	0.5	220	213
	1	234	232
	2	271	262
	10	366	354
Low pressure (0.5 bar) and variable bulk abundances	17% N ₂ , 80% CO ₂	198	191
	50% N ₂ , 47% CO ₂	220	213
	77% N ₂ , 20% CO ₂	224	214
Low N ₂ abundance	2% N ₂ , 95% CO ₂	151	150
	50% N ₂ , 47% CO ₂	234	232

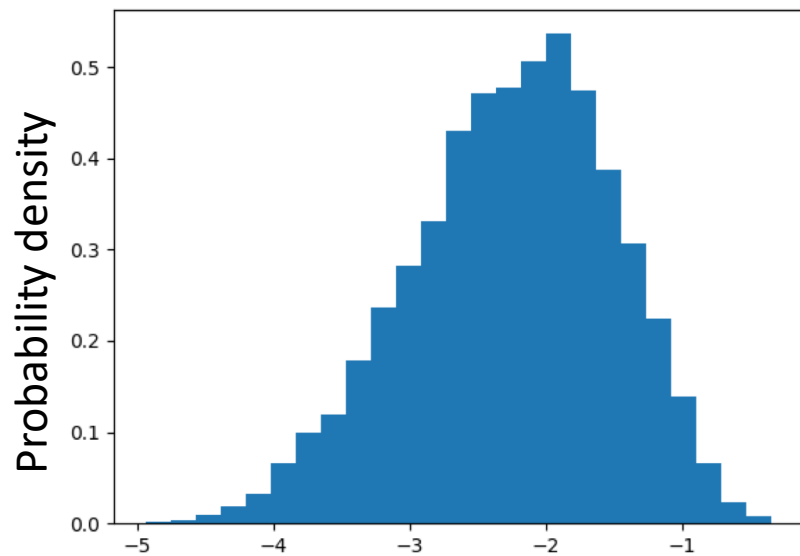
Getting from methane abundances to Pr(life)



+

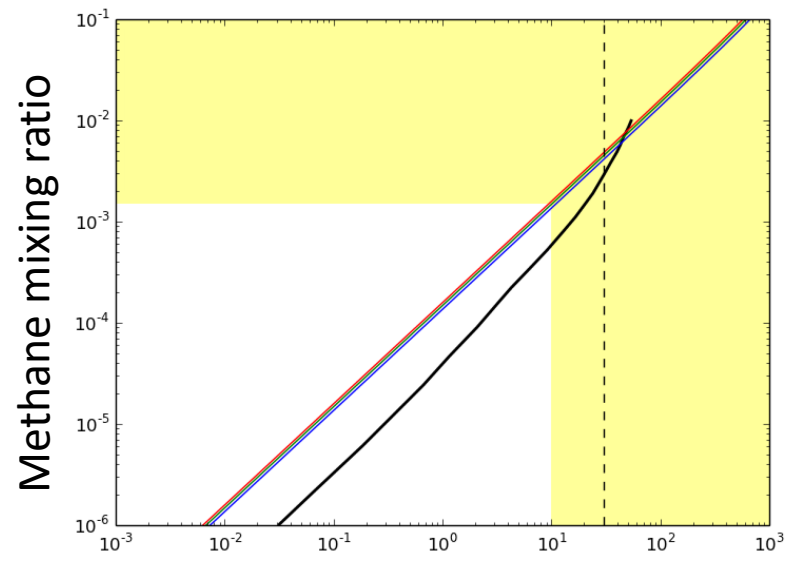


Getting from methane abundances to Pr(life)

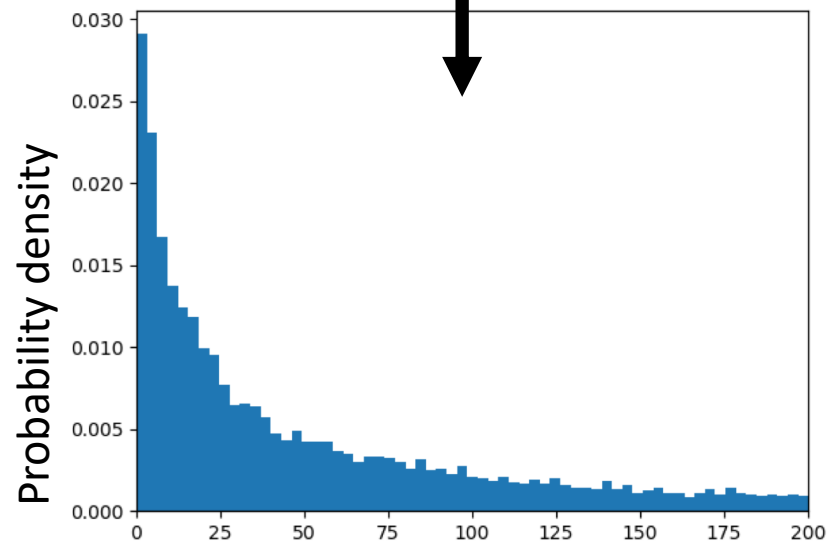


Retrieved methane abundance

+

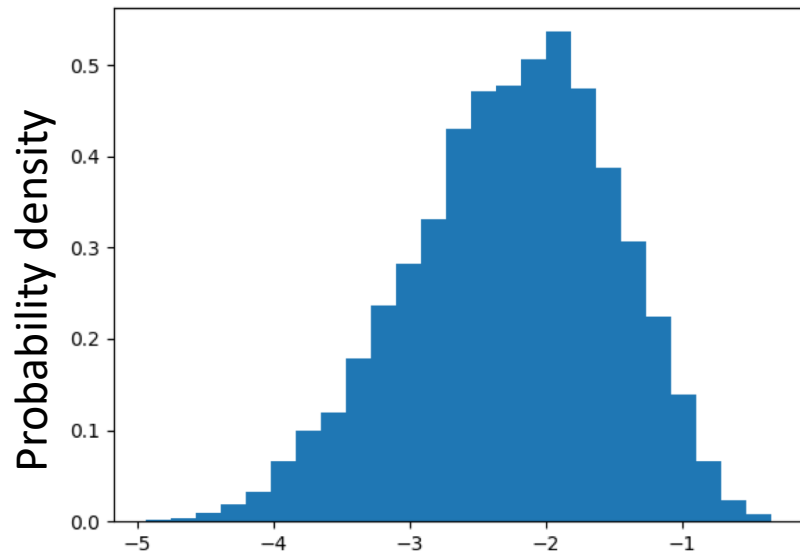


Methane flux



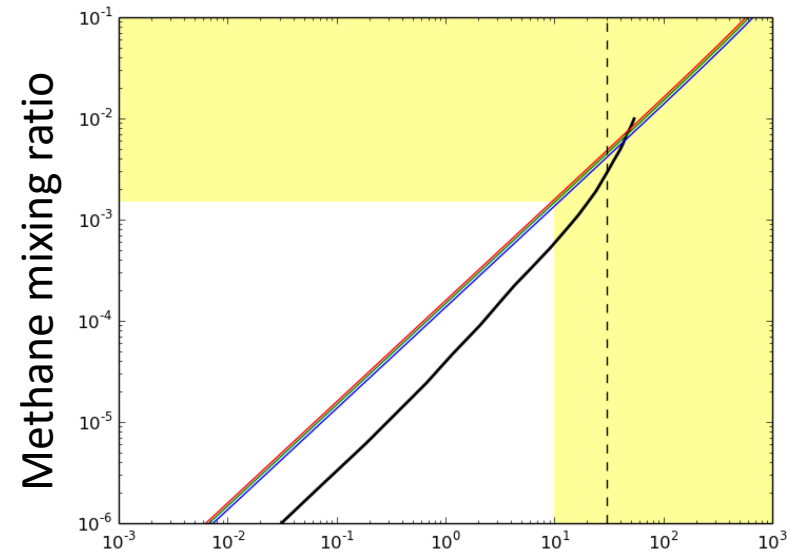
Required methane flux

Getting from methane abundances to Pr(life)

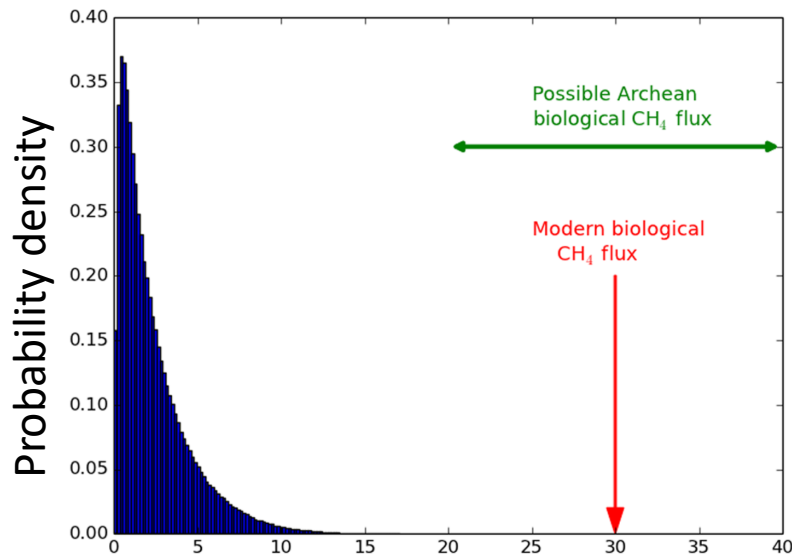


Retrieved methane abundance

+

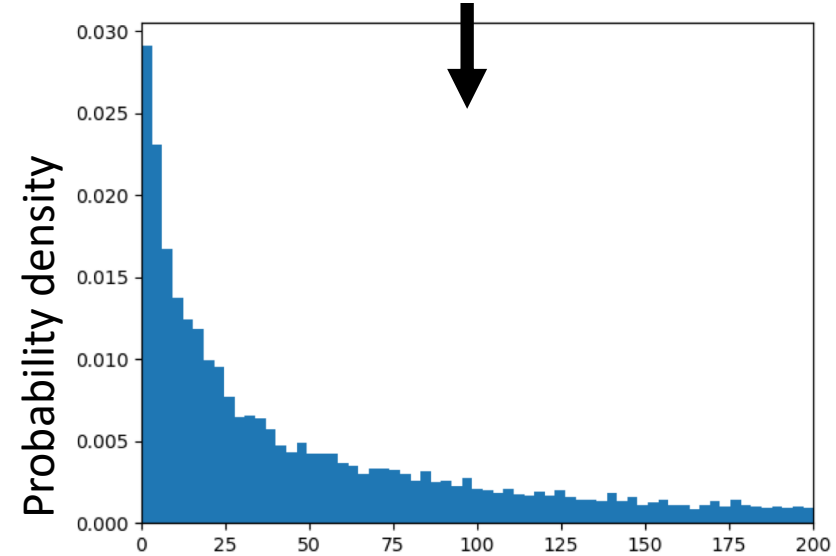


Methane flux



Methane flux

+

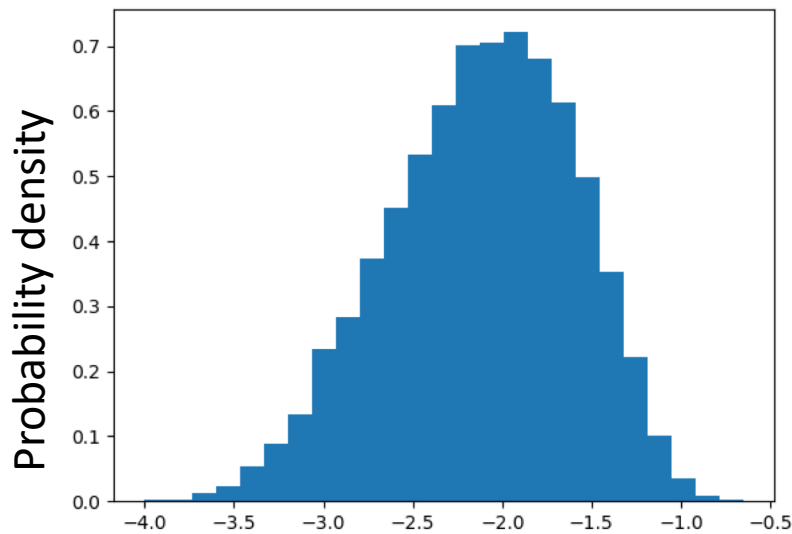


Required methane flux



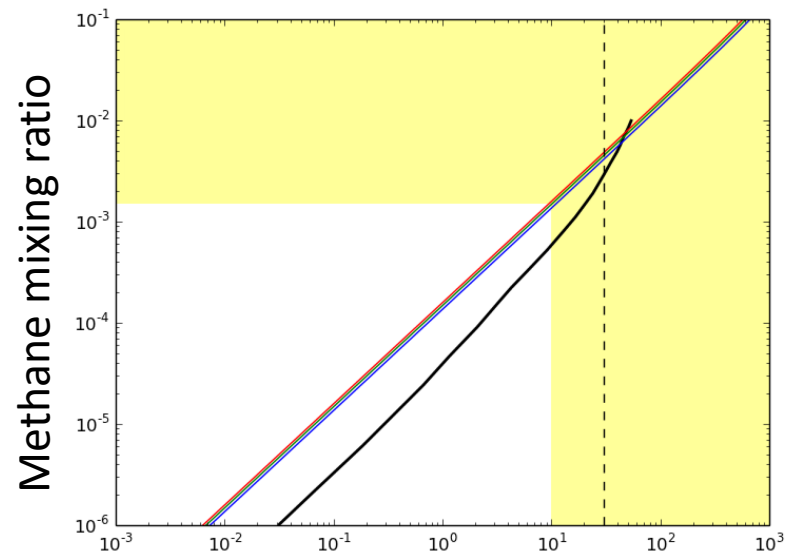
Pr(explain with known abiotic process) = 9% Typical 10 orbit case

Getting from methane abundances to Pr(life)

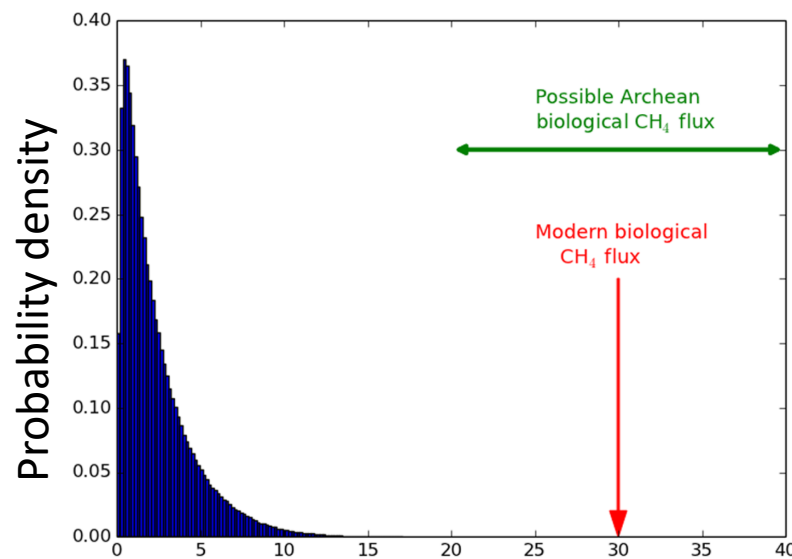


Retrieved methane abundance

+

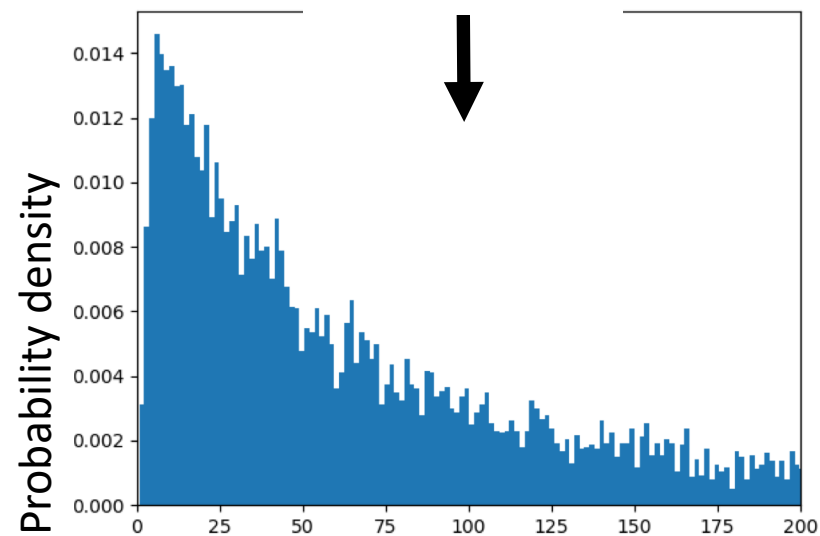


Methane flux



Methane flux

+



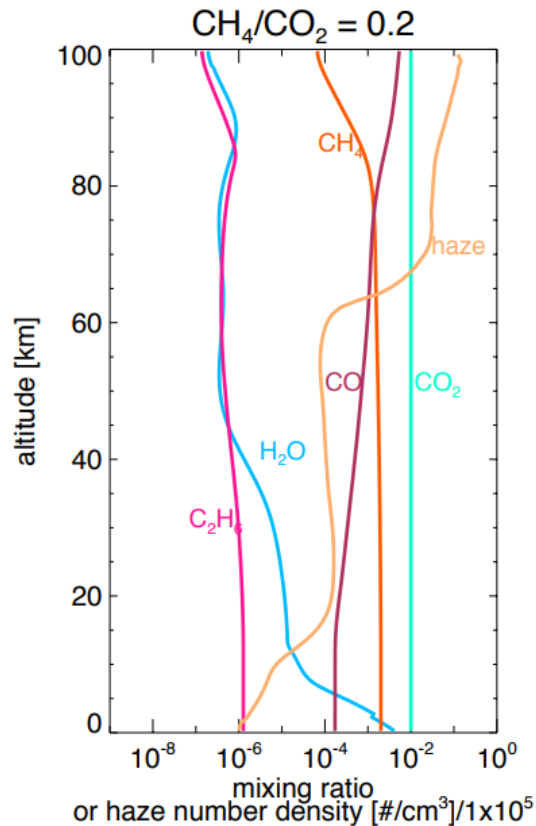
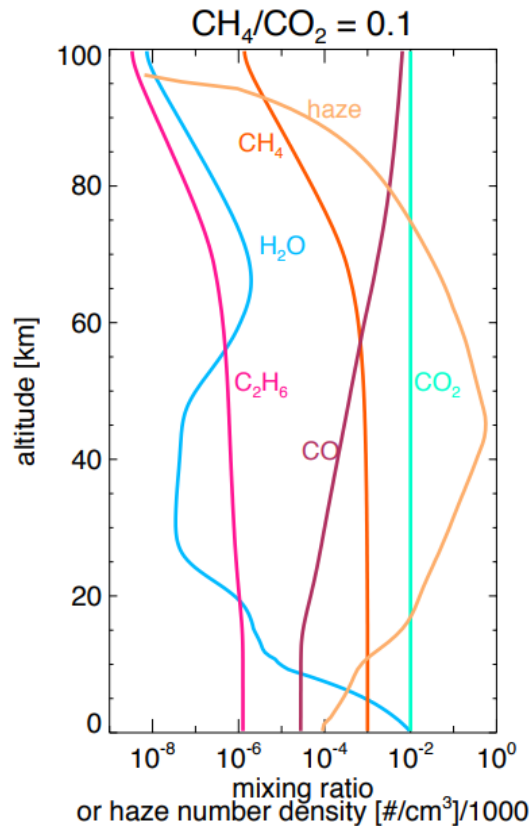
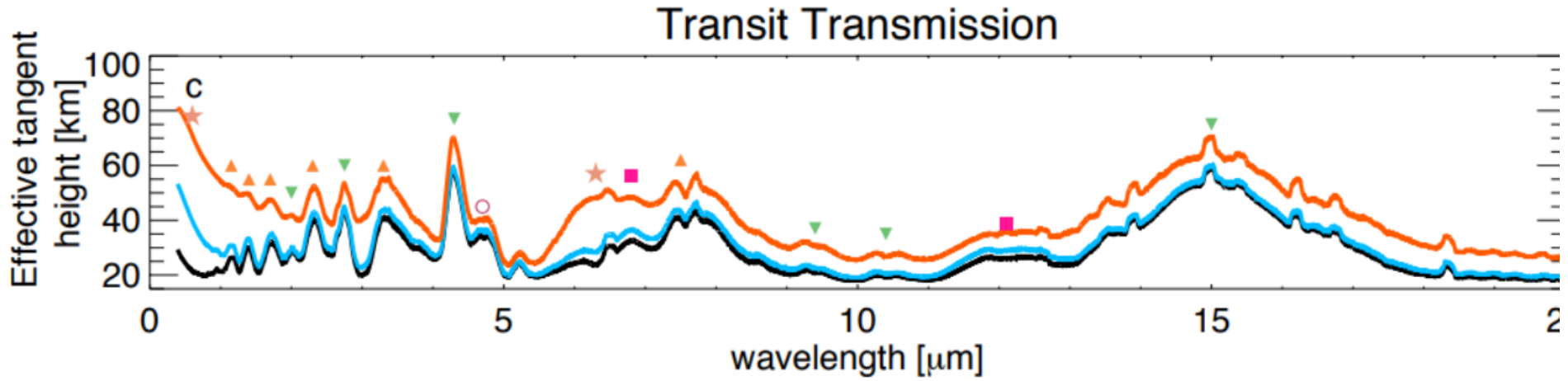
Required methane flux



Pr(explain with known abiotic process) = 2%

50 orbit case

Haze and vertical profiles



Arney et al. 2016 *Astrobiology*

Nested Sampling Retrieval algorithm

$$\Pr(\Theta|\mathbf{D}, H) = \frac{\Pr(\mathbf{D}|\Theta, H) \Pr(\Theta|H)}{\Pr(\mathbf{D}|H)}$$

Feroz and Hobson 2008, *MNRAS*

where $\Pr(\Theta|\mathbf{D}, H) \equiv P(\Theta)$ is the posterior probability distribution of the parameters, $\Pr(\mathbf{D}|\Theta, H) \equiv L(\Theta)$ is the likelihood, $\Pr(\Theta|H) \equiv \pi(\Theta)$ is the prior, and $\Pr(\mathbf{D}|H) \equiv \mathcal{Z}$ is the Bayesian evidence.

$$0 < X_M < \dots < X_2 < X_1 < X_0 = 1,$$

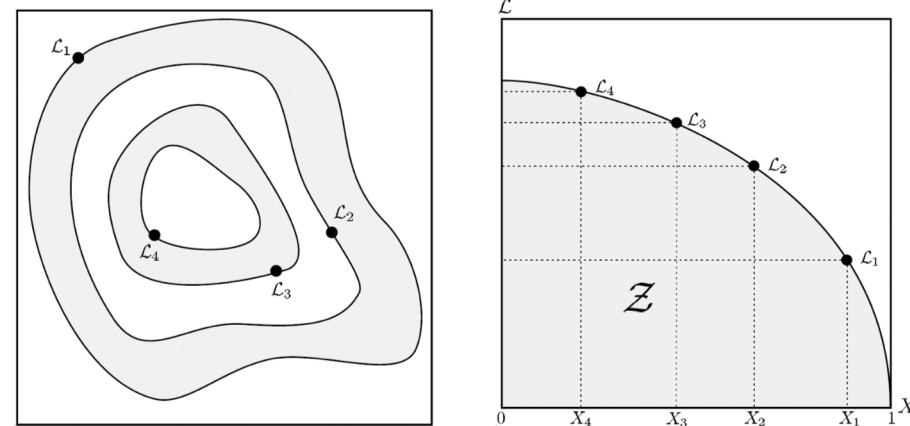
$$\mathcal{Z} = \int \mathcal{L}(\Theta) \pi(\Theta) d^D \Theta$$

Transform this horrible integral into a 1D integral over prior volume:

$$X(\lambda) = \int_{L(\Theta) > \lambda} \pi(\Theta) d^D \Theta,$$

$$\mathcal{Z} = \int_0^1 L(X) dX.$$

Can then calculate Bayesian evidence as summation over likelihood contours:

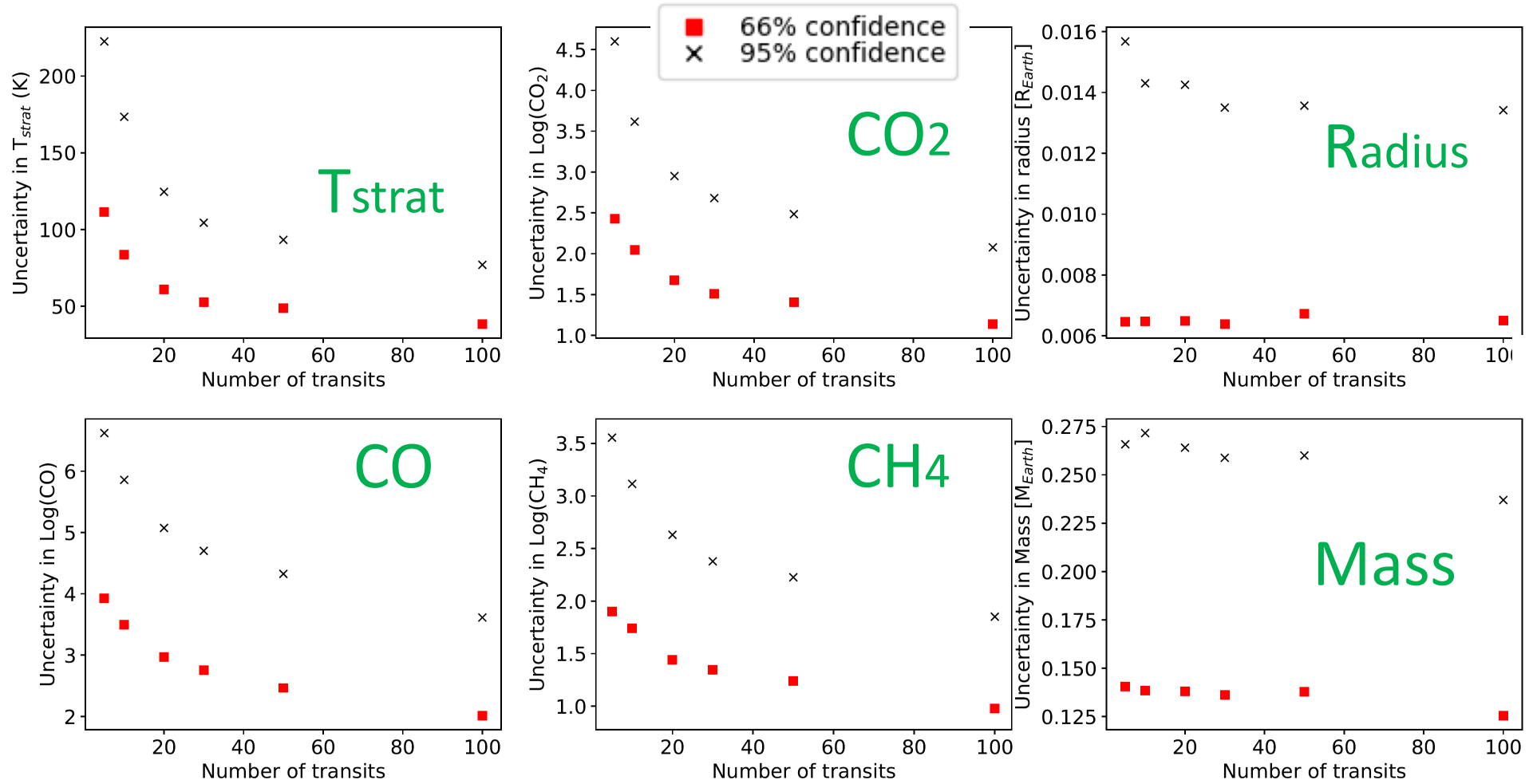


$$\mathcal{Z} = \sum_{i=1}^M L_i w_i.$$

Posteriors can then be easily calculated: $p_i = \frac{L_i w_i}{\mathcal{Z}}$.

Nested Sampling Produces same posteriors as MCMC

Does observing more transits help?



What about clouds?

Archean Earth-like Trappist-1e with Earth-like, grey clouds (10 transits with NIRSpec)

